

MIE SCATTERING AND THE MARTIAN ATMOSPHERE

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Abstract. It has been suggested that the discrepancy between radio occultation determinations of the Martian atmospheric surface pressure (3.8 to 7 mb) and those deduced from optical polarization measurements and a simple Rayleigh atmosphere model (about 10 mb) are the results of sub-micron sized aerosols in the Martian atmosphere. Based on observed viewing angle dependence of the polarization of the Martian disk in the visual range, a Mie scattering analysis has been made utilizing the measured complex index of refraction of limonite and bulk solid CO₂. The results of this study indicate that limonite aerosols alone are unsatisfactory to explain the viewing angle observations, whereas solid CO₂ (and H₂O ice) aerosol spheres, having a dominant particle radius range between 0.28 and 0.35 μ , could bring planetary and laboratory observations into compatibility. It is suggested, further, that solid CO₂ aerosols could explain limb brightening in the blue spectral range. Various distributions of solid CO₂ and H₂O Mie particles with radii up to 0.35 μ show an opposition effect. However, the role of these aerosols in explaining the Mars opposition observations is very dependent on the optical properties of the underlying Mars surface material.

Measurements of the Martian atmospheric surface pressure by the radio frequency occultation determinations of Mariners IV, VI, and VII yield values between 3.8 and 7 mb (Kliore *et al.*, 1966; Kliore, 1969). The spectroscopic data have been in reasonably good agreement, yielding a value of the order of 10 mb (Kaplan *et al.*, 1964). Still another technique, polarization measurement, has been yielding pressure values that are too large in comparison to the previously noted methods (Dollfus and Focas, 1966). Part of the problem with polarization lies in the nature of the laboratory specimen used for surface simulation as well as in the planetary model used for interpretation of laboratory polarization data compared to earth-based observations of the planet. Polarization is an extremely sensitive indicator of atmospheric as well as surface optical effects (Dollfus and Focas, 1966).

It cannot be emphasized too strongly that average gross features are involved in both telescopic observations and surface simulations, and as large a scale as possible of laboratory surface simulation is necessary (Egan, 1967). Further, simultaneous consideration of many aspects of surface simulation, e.g., photometry and polarimetry and thermal effects, etc., is required (Egan, 1969).

The polarimetric results of Dollfus have been examined and refined analysis yields an average surface pressure of about 20 mb (O'Leary and Pollack, 1969), a factor of 3 or more greater than the 4 to 7 mb value range deduced from Mariner spacecraft data. Our analyses based on further refinements of our large scale surface simulation experimental data have indicated a total surface pressure of about 10 mb for a simple Rayleigh atmosphere model (Foreman, 1969). Thus, an aerosol contribution equivalent to the optical properties of the order of 5 mb of atmosphere appears indicated. Various investigators (e.g., Kuiper (1964), Rea and O'Leary (1965), Coulson (1969)) have calculated the Mie scattering effects of terrestrial, water ice, and silica aerosols.

O'Leary (1967) has questioned the significance of any surface simulation attempt in view of the numerous possibilities for aerosol effects. Based on Mars observations and various optical models, our opinion is that the best approach still lies in a laboratory simulation as close as possible to Mars data and with an allowance being made for a generally small atmospheric perturbing effect.

More extensive Mie scattering calculations for estimating the Mars aerosol effect have become possible this past year with the determination of the complex index of refraction of Venango County (Pennsylvania) limonite (Egan and Becker, 1969) and of bulk solid CO_2 (Egan and Spagnolo, 1969) in the spectral range between 0.35 and 1.0μ . Both substances are potential candidate constituents of Mars aerosols. It is found that the refractive index of solid CO_2 is close to that of H_2O ice. Therefore, computations for dry ice may be considered as also approximately valid for H_2O ice particles.

This report summarizes a particular aspect of polarization that appears to permit an inference of the nature of aerosols in the Martian atmosphere; a Mie scattering computer program designed and run by Mrs. Jaylee Mead had supplied us with the single scatterer data with which this study has been concerned.

Polarization observations by Dollfus of the central light 'desert' regions of Mars, in the visual spectral range, exhibit a small but real variation of polarization with viewing angle (Figure 1) (Dollfus, 1966). This is of the order of 0.1% to 0.2% based

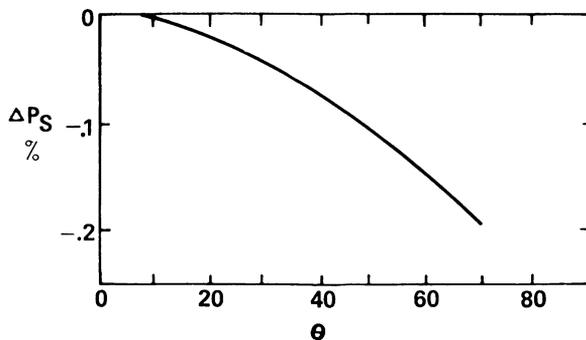


Fig. 1. Mars oblique polarization difference (Dollfus, 1965) $\Delta P_s = P_\theta^\circ - P_0^\circ$.

on the most updated information analyzed by Dollfus (Cann *et al.*, 1965). In the course of our simulation experiments, we also have discerned a comparable real variation of polarization with viewing angle (and wavelength dependent), however, of opposite sign from Dollfus' observations (Egan and Foreman, 1968). A possible explanation for this subtle paradox is the existence of Mie scatterers in the Martian atmosphere which could produce sufficient negative polarization to bring the particulate limonite surface simulation data into agreement with the planetary characteristics. Rayleigh

scattering of atmospheric molecules will only produce positive polarization, and, thus, cannot be expected to explain the problem. However, any atmospheric Rayleigh scattering polarization would have to be compensated for additionally by Mie scattering.

The change of Martian polarization with viewing angle shown in Figure 1 amounts to a few tenths of a percent difference in polarization between the center of the disk and the limb, for a particular phase angle, with the limb having the lower polarization.

Our laboratory simulation for the limonite particulate specimen that most closely matches the polarization-brightness characteristics of the bright regions of Mars (i.e., Venango County, Pennsylvania limonite (Egan, 1969)) indicates an oblique polarization difference between 60 degrees and 0 degrees viewing angles for three wavelengths as shown by Figure 2. The data points for a range of phase angles between 5 degrees

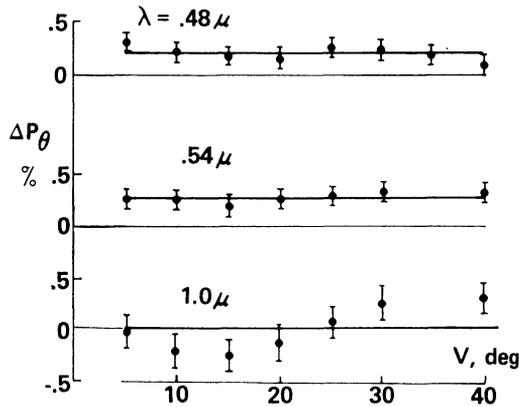


Fig. 2. Oblique polarization difference Venango limonite $\Delta P_{\theta} = P_{60^{\circ}} - P_{0^{\circ}}$.

and 40 degrees fall on either side of the rms average values shown. The data accuracy is higher for the 0.48μ and 0.54μ wavelengths because the photomultiplier sensor signal to noise ratio is greater than for the one used at 1.0μ . The approximate spectral variation of rms average values are shown in Figure 3. In the visual range (0.54μ), the limonite specimen will be in agreement with the Martian data if the aerosol component produces a decreasing polarization (i.e., negative ΔP) with increasing viewing angle, at any particular phase angle. (This trend also should hold true for all phase angles.)

Mie scattering by individual particles is dependent on the phase angle, V , not viewing angle, θ , except when the scatterers exist in a layer above a planet's surface. In this latter situation, as the viewing angle increases, the atmospheric and aerosol layer optical path length increases; the polarization contribution of a widespread aerosol layer can become relatively large with viewing toward the limb.

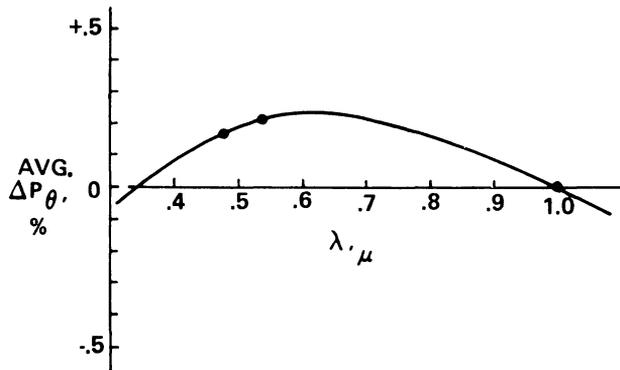


Fig. 3. Average oblique polarization difference – Venango limonite.

The results of the Mie scattering calculations* (simple spheres and single scattering) for Venango County limonite and bulk solid CO_2 at 0.54μ and 1.0μ are summarized in Figure 4. In this illustration is shown the variation of the $\sum_{r_0}^r B_r P_r$ product term with largest particle radius, r , for a flat (i.e., equal number) distribution of particle sizes between r_0 and r (r_0 is 0.005μ radius and r can be examined up to a maximum of 0.5μ). ($\sum_{r_0}^r B_r P_r$ represents the relative polarization of Mie aerosols, and is the numerically integrated product of relative brightness and polarization for integrating increments of particle radii, Δr , of 0.005μ . The summation is for all particle sizes up to the maximum, r , considered.) This product term is proportional to the polarization contribution of Mie aerosols to the total polarization characteristic of the planet, which also includes the atmosphere and the surface contributions. The curves are shown for a phase angle of 40 degrees.

Additional results were computed for various non-flat distributions (e.g., Gaussian, step rise-exponential decay) and were found to produce virtually no variation in qualitative results although some detailed quantitative differences were evident.

It is apparent that the Venango County limonite produces a negative polarization term at 0.54μ and zero polarization at 1.0μ (as required by Figure 3) almost uniquely at a dominant particle radius of about 0.27μ . However, additional data, not presented, but generated by our computer study, shows that at other phase angles between 10 degrees and 30 degrees, this dominant particle size yields positive polarization at $\lambda = 0.54 \mu$. Therefore, we can conclude that limonite is not a good choice, in itself, for Martian aerosols, at least under Mie scattering theory constraints.

For solid CO_2 , the situation appears different. In the particle size range between about 0.28 and 0.35μ radius, the viewing angle variation of polarization difference, ΔP , is of the correct sign and magnitude at the two wavelengths. (It should be noted

* Provided by Mrs. Jaylee Mead in private communications, July–October, 1969.

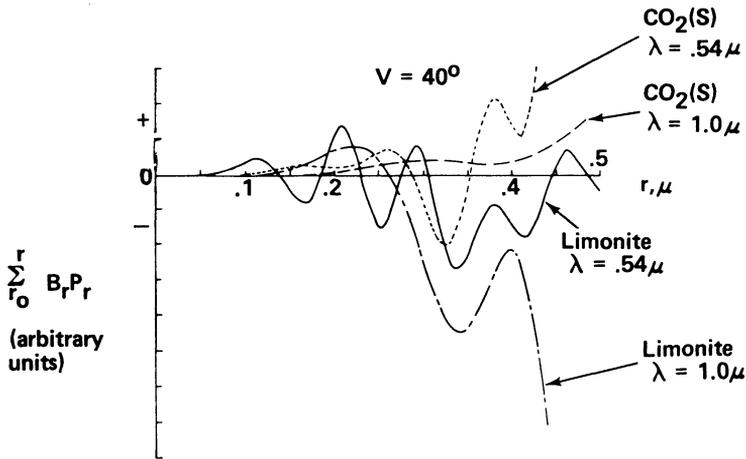


Fig. 4. Integrated product of brightness and polarization vs. Mie scatterer size.

that for $\lambda = 1.0 \mu$, we must subtract the summation ordinate value for all particle radii sizes smaller than 0.28μ from the value at larger sizes and this results in a virtually zero polarization term for dominant particle radii between 0.28μ and 0.35μ , as required.)

If we examine the photometric properties of solid CO_2 and limonite as a function of Mie particle radii at a 0 degrees phase angle, for instance, Figure 5, it is evident that the solid CO_2 particles are generally an order of magnitude brighter in the blue (0.365μ) than in the red (0.740μ), and increases in brightness with particle size (up to at least $r = 0.5 \mu$). Limonite aerosols, however, generally do not exhibit a strong color differentiation, and, in many size ranges are of the same order of brightness. Thus, if solid CO_2 aerosols of submicron size exist at the limbs of Mars, they could be expected to contribute to the commonly observed limb brightening in the blue portion of the spectrum.

In Figure 6, a further examination of solid CO_2 aerosols for other phase angles in the visible region shows that the particle size range between 0.28μ to 0.35μ remains compatible in sign, although varying, to some degree, in magnitude for the smaller phase angles.

Continuing to explore the photometric properties of solid CO_2 aerosols, in association with a Venango County limonite simulated surface, as a function of wavelength, one can see from Figure 7: (1) a 1–2 mm particulate limonite surface, at a wavelength of $0.54 \mu^*$ has an opposition effect close to that observed by O'Leary (1967) for Mars but above that inferred from de Vaucouleurs (1964); (2) solid CO_2 aerosols spheres, with a flat size distribution between 0.005μ radius and approximately 0.33μ , show an opposition effect particularly at the shorter wavelengths. It should be noted that in

* Currently, data available are for this wavelength only.

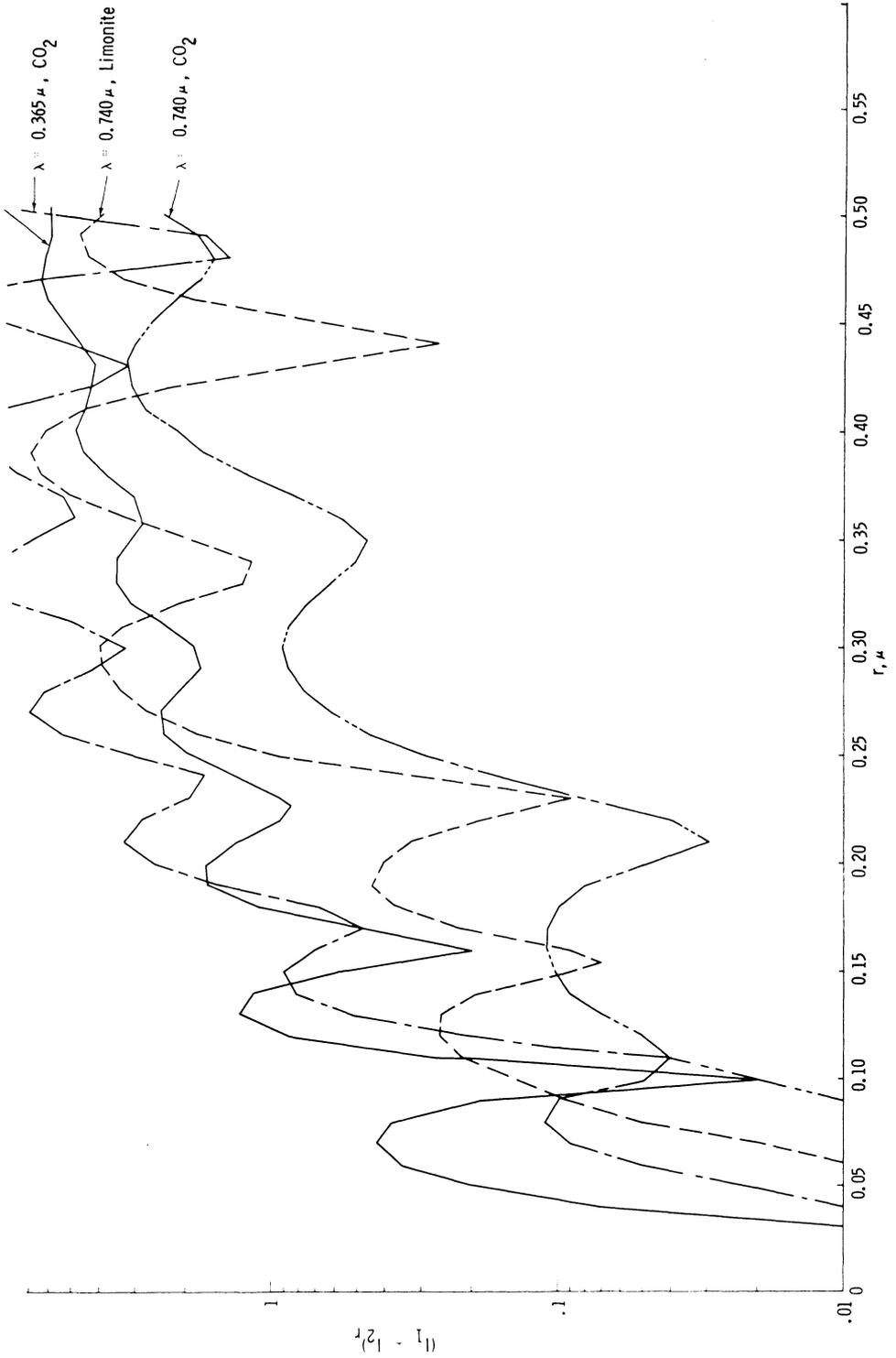


Fig. 5. Relative brightness of Mie particles ($V=0^\circ$).

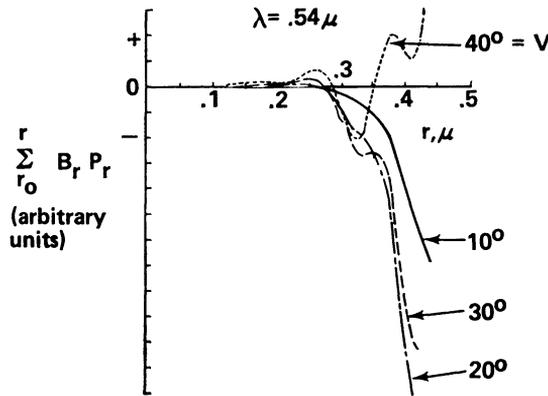


Fig. 6. Integrated product of brightness and polarization vs solid CO₂ particle size and phase angle.

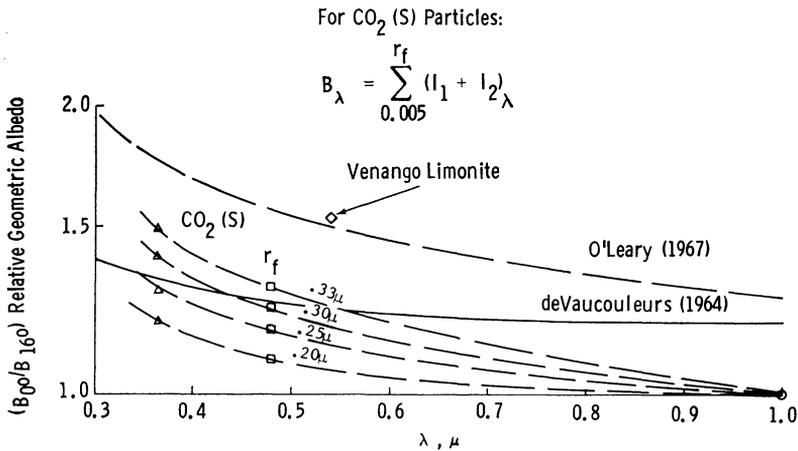


Fig. 7. Opposition effect: Mars observations compared to 1–2 mm limonite specimens and Mie calculations for solid CO₂.

the solid CO₂ particle size range inferred by our previous polarization considerations (i.e., ~0.300 to 0.325 μ radius), the opposition effect may be as large as 1.80 for λ=0.365 μ. In numerical computations of Mie scattering photometric characteristics, it is important to use small radii increments (Δr ≈ 0.005 μ) for the particles so that the fine details in the results will not be obscured. In addition, it is advisable to examine phase angles greater than the 16 degrees reference angle, usually employed for opposition analyses, for possible secondary maxima in the opposition effect curve; we have found no such violation of the opposition effect in our analyses of computer results for the ~0.30 μ particle radius size deduced for solid CO₂.

It should be noted that if other wavelength determinations with limonite surface specimens similarly follow the trend (Figure 7) of the Mars opposition observations, then the need for aerosol particles to explain the opposition phenomena are considerably lessened. Additional studies of opposition effect are discussed in the other conference paper by Mead (1970). Her results are based on a particular and different set of assumptions, but her conclusions do not conflict with those of this paper which are based on a limonite surface simulation model that closely matches the Martian observations.

Figure 8 presents the variation of vapor pressure of solid CO_2 with temperature, and an indication of the equilibrium temperatures between the solid and a pure gaseous CO_2 atmosphere. For an atmospheric surface pressure of 7 mb, near the

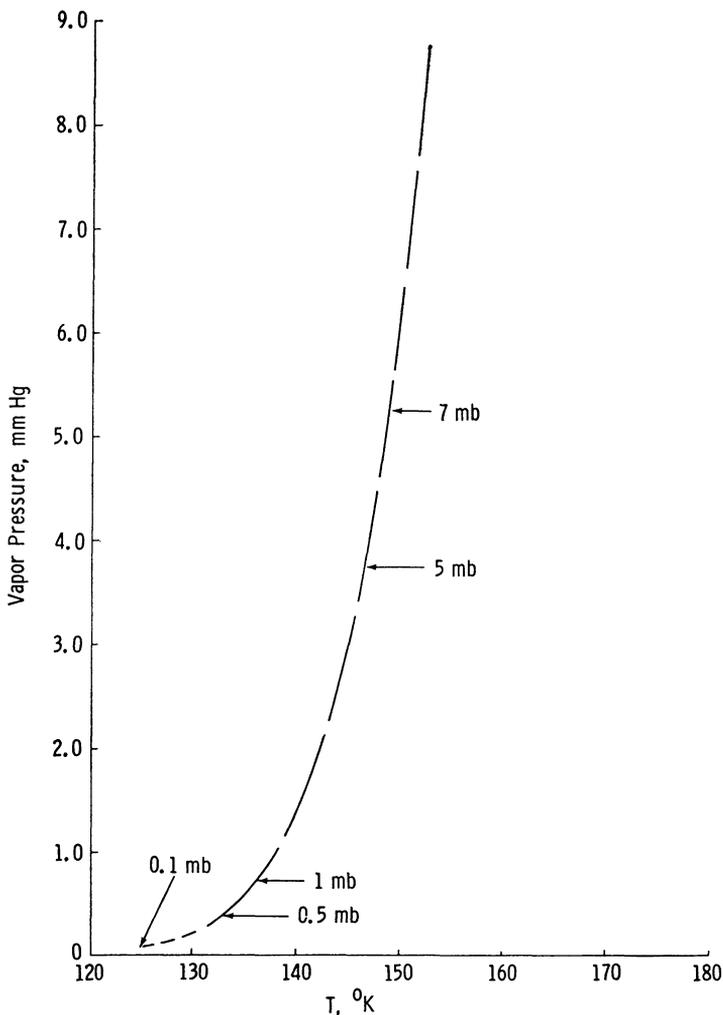


Fig. 8. Vapor pressure of solid CO_2 (NBS Circular No. 564, revised).

Kliore *et al.* inferred surface pressure (Kliore *et al.*, 1966; Kliore, 1969), the equilibrium temperature is almost 149 K. Radiation-dominated models of the Martian atmospheric dynamics have been studied by Gierasch and Goody (1967/1968) and earth type convective models by Leovy and Mintz (1966).

Profile models of the Martian atmosphere have temperature distributions which depend upon the assumed condition of surface temperature (Gierasch and Goody, 1967/1968; Leovy and Mintz, 1966; Johnson, 1965). Except for polar and high winter latitudes, the near-surface temperature generally does not seem to approach a low enough value to result in dry ice production. However at altitudes where the atmospheric pressure is of the order of a magnitude lower or less (i.e., at 10 km or above) atmospheric temperatures may become low enough because of atmospheric radiative and dynamic processes to coincide with the temperature at which solid CO₂ will form, as indicated by the phase curve. The 1969 Mariner spacecraft photographs (Leighton *et al.*, 1969) show that clouds do occur at altitudes of approximately 15 to 40 km above the Martian surface and it is reasonable to consider these clouds as consisting of solid CO₂ as well as H₂O ice.

The mean aerosol particle size of the order of 0.30 μ radius that is inferred by our results, could be considered also to be representative of a distribution of aerosol particle sizes between, say, 0.28 to 0.35 μ . It might be noted that this size range is quite unlike our terrestrial experience (Junge, 1967).

However, the total environment of Mars apparently is quite different from that of the Earth, and perhaps these results are not as unreasonable as they first may appear. Certainly, the limitations of the Mie scattering theory constraints must be recognized. In addition, our assumptions about a simple quiescent aerosol volume distribution are rightly questionable. Nevertheless, this study has improved the perspective concerning Mars, and hopefully enables one to discern what remains significant in the abundant speculation about the Martian atmosphere that pervades the literature.

Acknowledgments

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