A new glacier inventory of Iran

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ABSTRACT. A new glacier inventory of Iran, compiled according to GLIMS guidelines through the use of photogrammetry and remote sensing supported by fieldwork, provides the first comprehensive study of its mountain glaciers. The glaciers are found in five main areas: two in the higher elevations of the Alborz mountain range (Damavand and Takhte-Soleiman regions), two on the Zardkuh and Oshtorankuh mountain chain in the Zagros mountain range and one in the Sabalan Mountains in northwest Iran. Several important glacier attributes, including minimum and maximum height of ice, area and maximum length and width, together with glacier extent, were successfully extracted using aerial and satellite imagery. Thereafter a comprehensive glacier database was established in a GIS environment.

INTRODUCTION

Glacier changes are key indicators for ongoing global climate change. They influence hydrologic regimes and may trigger natural hazards. In addition, they are unique sources of fresh water and an important economic factor for hydropower production (Paul, 2003).

In recent decades, a spectacular retreat has been observed on most of the monitored glaciers in the world. If this evolution is confirmed in coming years, it will have important consequences in terms of water resources, economic development and risk management in the surrounding areas. Up to now, only a few of the world's mountain glaciers have been monitored, mostly by ground measurements, which often provide information only once or twice a year at a few points (Trouvé and others, 2007).

Due to the difficulty of reaching high-altitude glaciers in risky mountainous areas, even with the advent of the differential global positioning system (DGPS) it remains difficult and time-consuming to perform regular ground-based surveys of glaciers (Berthier and others, 2005). The acquisition and processing of remotely sensed data should provide more information and improve glacier monitoring.

The use of satellite data for updating or generating glacier inventories has been discussed for many years. Furthermore, the combined use of satellite imagery and digital elevation models (DEMs) in a Geographic Information System (GIS) environment is a valuable method of efficient operational data processing (Paul, 2003).

In Iran, monitoring glaciers is especially important, since they supply a considerable portion of the drinking water. Glaciers in Iran are found in five main areas (Fig. 1): two are located in the higher elevations of the Alborz mountain range in northern Iran (one in the vicinity of Takhte–Soleiman and the other on Mount Damavand); the others are found on the Sabalan Mountains in northwestern Iran, and on the Zardkuh and Oshtorankuh Mountains in the Zagros mountain range of southwestern Iran.

Based on the geological structure of the underlying bedrock of the glaciers, we divided Iran's glaciers into two categories: (1) granite glaciers and (2) karstic glaciers. These two types of glaciers have different glacial features (e.g. the thickness of ice in karstic glaciers is less than that of granite glaciers). Moreover, special features such as crevasses and moraines are not found on karstic glaciers, whereas sinkholes are observed on their surface. The glaciers in Takhte–Soleiman and Sabalan regions and around Damavand summit have the granitic structure, whereas Oshtorankuh and Zardkuh glaciers are karstic.

Unfortunately, the glaciers in Iran were not thoroughly investigated until recently. The little work that was conducted by Europeans and Iranians was mostly ground-based (Ferrigno, 1991).

Here we present the results of the very first glacier inventory in Iran through the use of photogrammetry and remote sensing. Several important glaciers attributes including minimum and maximum height of ice, area and maximum length and width, along with glacier extent, have been successfully extracted using aerial and satellite imagery. These results have also been verified by extensive fieldwork. A comprehensive glacier database was established in a GIS environment, with respect to attributes defined in the Global Land Ice Measurements from Space (GLIMS) project.

DATA

Aerial and satellite image data

Airborne photography has been a major source of information on glaciers, as it is able to provide high-resolution data for the computation of high-resolution digital terrain models (DTMs) and orthoimages with accurate positions, obtained using GPS control points. However, due to the high cost of each campaign, few aerial photographs are available, which leads us to make use of another important source of information, that acquired by optical sensors on remotesensing satellites. Multitemporal satellite imagery, with its increasing resolution, large spatial coverage and spectral information, can be used to identify specific glacial features for change detection and risk assessment.

In this work, a variety of satellite images (Table 1) acquired by different sensors, along with aerial photographs, were utilized to compile the Iran Glacier Inventory (IGI).

Data pre-processing

In the pre-processing stage, selection of the most suitable satellite scenes, as well as geometric correction, were the most important tasks. For data selection, some limitations existed in terms of cloud cover at high altitudes and snow

Table 1. Optical satellite imagery used in the Iran Glacier Inventory 2008. ETM+: Enhanced Thematic Mapper Plus; IRS: Indian Remotesensing Satellite; SPOT: Système Probatoire pour l'Observation de la Terre; ASTER: Advanced Spaceborne Thermal Emission and Reflection Radiometer

Satellite image	Date	Glacier region
QuickBird (panchromatic) QuickBird (multispectral) Landsat ETM+ Landsat ETM+ IRS P5 (panchromatic) SPOT5 (panchromatic) ASTER ASTER IRS (panchromatic)	September 2005 September 2005 September 2002 September 2002 September 2007 September 2004 August 2001 August 2001 September 2003	Takhte-Soleiman Takhte-Soleiman Damavand Takhte-Soleiman Takhte-Soleiman Takhte-Soleiman Oshtorankuh Zardkuh Zardkuh

cover, which had to be taken into consideration. Hence, images acquired at the end of the ablation period, around September, were chosen. Having selected suitable satellite imagery (level 1 data that had already been radiometrically corrected), geometric correction and georeferencing of images were carried out.

DEM generation

Several series of aerial photographs at scales of 1:40 000 and 1:50 000 were selected over the five glacier zones in Iran. These photographs were provided by Iran's National Cartographic Center. The $\sim\!23\,\text{cm}\times23\,\text{cm}$ negatives were scanned at a resolution of 14 μm .

Firstly, the digital images were oriented through photogrammetric processing. The required ground-control points (GCPs) were acquired using the 'bundle-block aerial triangulation' technique, which enables global restitution of a block of photographs and reduction of GCPs. Bundleblock adjustment is an iterative method based on the use of photographic coordinates as observations. Application of the central projection method enables the conversion of these observations into terrain coordinates in one step (Trouvé and others, 2007). This requires a block of photographs with at least 60% overlap and 20% side-lap (often there is more) (Trouvé and others, 2007). Moreover, sets of GPS points located on both sides of the glaciers in different study areas were required for the orientation process. The number of points used in each study area varied, depending on data availability. The matching procedure was carried out both automatically and manually. Wherever automatic correlation was problematic, due to factors such as snow-cover conditions or poor contrast, stereoscopic measurement was required. Eventually, DEMs and ortho-photographs of glacier zones were produced using Virtozou NT photogrammetric software. Later, orthoimage mosaics of the glacier areas were also generated.

GLACIER INVENTORY

Iran's glaciers have been the subject of many studies by several European and Iranian researchers since 1933 (e.g. Busk, 1933; Desio, 1934; Bobek, 1937; Harding, 1957; Péguy, 1959; Schweizer, 1970; Troll, 1973; Grunert and others, 1978; Ferrigno, 1991; Rezaei, 2003; Vaziri, 2003)

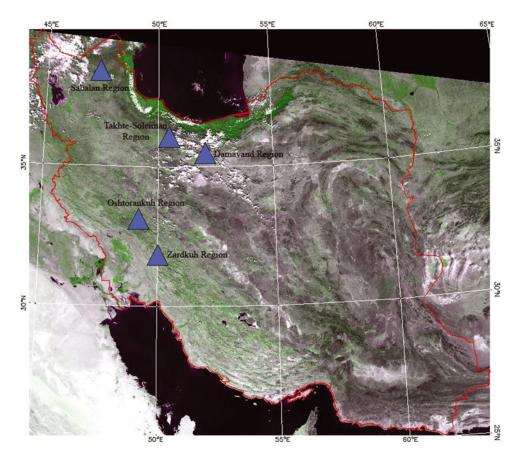


Fig. 1. Glacier regions in Iran (MODIS (moderate-resolution imaging spectroradiometer) image).

However, none of these represent a comprehensive inventory of the glaciers and their characteristics; for instance, one glacier region (Oshtorankuh) was not identified at all. Moreover, there was no accurate information on the total areal coverage of glaciers in Iran. A rough estimate was made to be $20\,\mathrm{km^2}$ (Ferrigno, 1991), based on estimates of $14\,\mathrm{km^2}$ for the glaciers of the Takhte–Soleiman region and 2, 2.4 and $1.5\,\mathrm{km^2}$ for Damavand, Sabalan and Zardkuh regions, respectively.

In this study, initially five main glacier regions in Iran were identified through the use of aerial and satellite imagery and extensive fieldwork (Fig. 1). The fieldwork was carried out by F. Vaziri, who has been studying the glaciers since 1983. His work (Vaziri, 2003) remains the most valid reference on Iran's glaciers at present.

In this inventory, the current size of Iran's glaciers is estimated to be $27 \, \mathrm{km^2}$. It is found that the glaciers in the Takhte–Soleiman region occupy a total area of $7.55 \, \mathrm{km^2}$, and glaciers in Damavand, Sabalan, Zardkuh and Oshtorankuh regions occupy areas of 3.4, 4.27, 7.1 and $4.8 \, \mathrm{km^2}$, respectively.

Related physical attributes were calculated through image-processing techniques, and the results were validated by ground survey. In the following we present a thorough description of each glacier region, along with its extracted attributes.

Takhte-Soleiman region

The largest glaciers in Iran exist in the western part of the Alborz mountain range, in the region of the Takhte–Soleiman massif (36°22′ N, 50°57′ E). The highest peak of this region, Alamkuh (4850 m a.s.l.), is the second highest peak in Iran. The region consists of 45 peaks >4000 m in elevation that are joined by high, narrow summit ridges. The glaciers are found mostly on the steep, northward-facing slopes of these peaks and ridges (Vaziri, 2003).

In this inventory, the glaciers in Takhte–Soleiman region were comprehensively studied through the use of aerial and satellite imagery. The largest glacier was found to occupy an area of $\sim 3.3 \, \mathrm{km^2}$. It is called Northern Glacier, previously known to researchers as Sarchal Glacier. It is formed from the confluence of three unequal segments (Fig. 2). The surface of the glacier tongue is covered with debris 0.5–1.5 m thick. According to Vaziri, each of these segments is considered to be a glacier, and 1, 2, 3 in Figure 3 were named Alamchal, Patakht and Takhte–Soleiman, respectively. The height of the snowline is estimated to be $\sim 4000-4150 \, \mathrm{m}$. The thickness of ice on this glacier varies from 50 m in the lower part of the



Fig. 2. Three-dimensional visualization of Northern Glacier with its digitized outline (QuickBird image 2005). (Takhte–Soleiman region.)

glacier to 80 m on the cirque (Fig. 4a) (Vaziri, 2003). Several crevasses are observed on the surface of this glacier (Fig. 4b).

In Figure 3, four other mountain glaciers can be noted in the Takhte–Soleiman region: Haft-Khan, Merjikesh, Khersan and Western Glaciers.

The DEM of this region was used to obtain three-dimensional (3-D) glacier attributes, such as minimum and maximum height of ice occurrence. Glacier area was provided within the GIS automatically for each glacier, and glacier length was obtained from digitized central flowlines. Further 3-D parameters (e.g. slope and aspect) were obtained from DEM-derived products.

Two 3-D DEM parameters are displayed in Figure 5 for the Takhte–Soleiman region. In Figure 5a, grey shades from black to white denote slope from high to low values. A mean slope for each glacier is also provided in Table 2. The aspect values in Figure 5b are coded as eight grey levels, representing eight aspect sectors.

In Table 2 a list of the glaciers in the Takhte–Soleiman region, along with their physical attributes extracted from

Table 2. Physical attributes of glaciers in Takhte–Soleiman region

Glacier		Orientation	Minimum height m	Maximum height m	Area km²	Maximum length km	Maximum width km	Mean slope °
Northern	Alamchal	Northeast	3630	4360	2.27	4.3	1.32	12.79
	Patakht	Northeast	3850	4130	0.47	1.9	0.36	13.27
	Takhte-Soleiman	Northeast	3740	4220	0.56	1.6	0.25	15.19
Western		Northwest	3900	4330	0.46	1.7	0.47	21.33
Haft-Khan		North	3740	4250	2.52	3.06	0.4	20.06
Khersan		Southeast	4065	4425	0.54	2	0.24	16.93
Merjikesh		Southeast	4070	4460	0.73	1.37	0.22	21.89

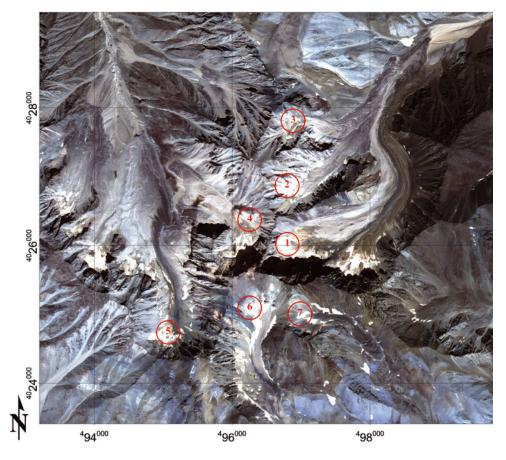


Fig. 3. QuickBird image of the Takhte–Soleiman region. 1. Alamchal Glacier; 2. Patakht Glacier; 3. Takhte–Soleiman Glacier; 4. Western Glacier; 5. Haft-Khan Glacier; 6. Khersan Glacier; 7. Merjikesh Glacier.

2005 QuickBird panchromatic imagery (spatial resolution 0.6 m) and the DEM, are presented. Two close-range photographs of the glaciers in the Takhte–Soleiman region are shown in Figure 4.

Sabalan region

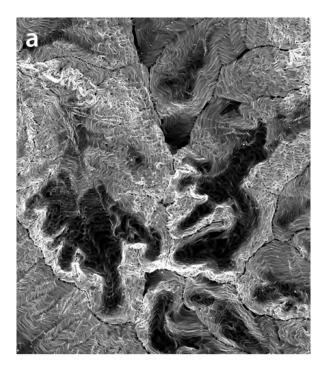
This glacier region lies in the northwestern part of the country and is located 85 km from the city of Ardebil in eastern Azerbaijan province. The area has a continental climate with hot, relatively dry summers and extremely cold, snowy winters. The precipitation, which falls almost exclusively as snow in the late autumn, winter and spring,

nourishes the glaciers on the highest peaks; however, during the arid summer, there is considerable ablation (Schweizer, 1970). The height of the snowline is estimated to be 4000 m. The highest peak in this region, Mount Sabalan (38°16′ N, 47°50′ E), is an inactive stratovolcano that rises to a height of 4811 m a.s.l. (Vaziri, 2003). It has a permanent crater lake formed at its summit, which is covered with ice from September through to June the next year. This freshwater lake occupies an area of 0.01 km² and is 15 m deep.

The glaciers on Sabalan are often characterized by a thick, continuous debris cover, and many glaciers appear to grade into or continue as rock glaciers (Ferrigno, 1991). As can be



Fig. 4. Close-range photographs of Alamchal Glacier (October 1995): (a) the glacier's cirque; (b) a crevasse on the surface of the glacier. (Photographs taken by F. Vaziri.)



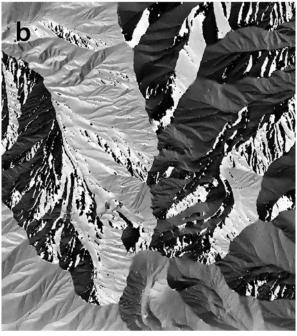


Fig. 5. Illustration of products from the DEM of the Takhte–Soleiman region: (a) slope, (b) aspect.

seen in Figure 6, penitents are found on the summit of Sabalan: they are an annual phenomenon of snow sublimation, related to solar radiation in an arid climate (Ferrigno, 1991). Although penitents form on the highest peaks in Iran, climate is more important than elevation to their formation. The incidence angle of the Sun, the intensity of the radiation, the dryness of the air in late summer and early autumn, and the time of maximum development, determine the occurrence, size and shape of the penitents (Ferrigno, 1991).

The main active glacier of this region, Northern Glacier (Fig. 7a), is a cirque glacier and includes three parts with special characteristics. The steep first part extends from the top of the glacier down to 4300 m a.s.l., where a large crevasse crosses the whole glacier. In the second part,

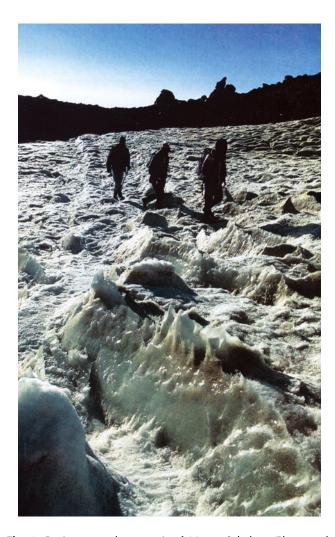


Fig. 6. Penitents on the summit of Mount Sabalan. (Photograph taken by F. Vaziri.)

spanning elevations from 4300 to 3900 m a.s.l., surface moraines of different sizes can be found, as well as linear crevasses in which water flows. The third part, continuing down to 3540 m in elevation, consists of coarse glacial sediments, including big cobblestones. In the northern part of Sabalan region, in Shirvan Valley, a big valley glacier exists which is a continuation of Northern Glacier. Five glaciers (Fig. 8) are identified in this region. Table 3 presents the related physical attributes, extracted from aerial photographs. Three rivers, namely, Anar-chai, Ghoori-chai and Darvish-chai, originate from the glaciers in the Sabalan region (Vaziri, 2003).

Damavand region

At present, all of the glaciers in the eastern Alborz mountain range are located on Damavand ($35^{\circ}57'$ N, $52^{\circ}06'$ E). Mount Damavand (5609 m a.s.l.), a dormant stratovolcano located ~ 70 km northeast of Tehran, is the highest peak in Iran (Vaziri, 2003). The summit crater is ~ 160 m wide and 20 m deep and has a small lake in it which is frozen most of the year.

Despite the low mean annual temperature and high level of precipitation in comparison with other glacial regions, due to the steep slope of the conical-shaped mountain and the heat induced by volcanic activities, glacial erosional features (e.g. glacial cirques) have not developed. However,

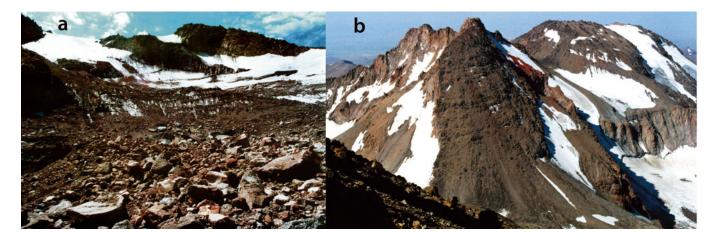


Fig. 7. Close-range photograph of Sabalan glaciers (July 1995). (a) Northern Glacier. (b) Glaciers around Herame-kasra summit: 1. Herame-kasra Glacier; 2. Herame-kasra Southeastern Glacier; 3. Southern Glacier. (Photographs taken by. F. Vaziri.)

Table 3. Physical attributes of glaciers in Sabalan region

Glacier	Orientation	Minimum height m	Maximum height m	Area km²	Maximum length km	Maximum width km
Northern	North	3660	4765	0.82	1.7	0.63
Northwestern	Northwest	3800	4550	0.62	2.1	0.3
Herame-kasra (Fig. 7b)	North	3600	4460	0.93	2.9	0.5
Herame-kasra Southeastern (Fig. 7b)	Southeast	4000	4300	0.3	1.25	0.12
Southern (Fig. 7b)	Southeast	3700	4330	1.6	3.4	0.6

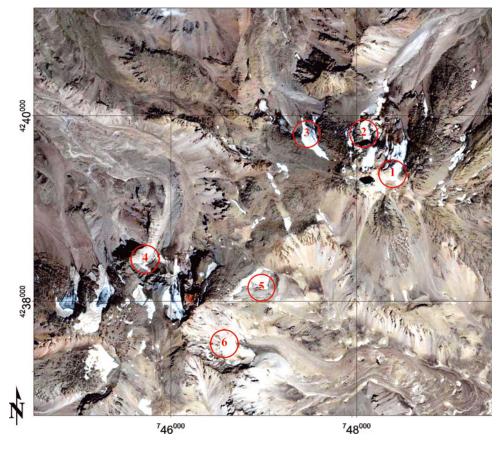


Fig. 8. QuickBird image of Sabalan region (Google Earth). 1. Sabalan summit crater; 2. Northern Glacier; 3. Northwestern Glacier; 4. Herame-kasra Glacier; 5. Herame-kasra Southeastern Glacier; 6. Southern Glacier.

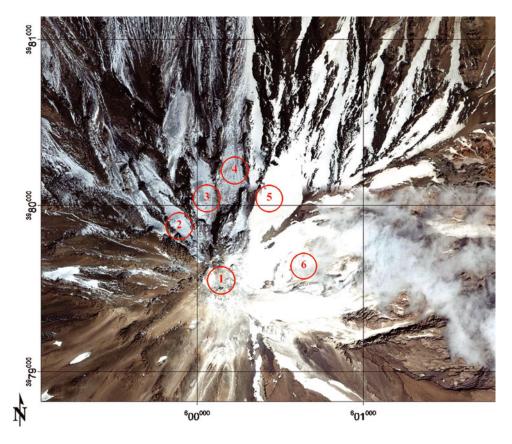


Fig. 9. QuickBird image of Mount Damavand (Google Earth). 1. Damavand summit crater; 2. Siuleh Glacier; 3. Dubi-sel Glacier; 4. Speleh Glacier; 5. Khurtabsar Glacier; 6. Yakhar Valley Glacier.



Fig. 10. Yakhar Valley Glacier (July 2003). (Photograph taken by F. Vaziri.)

there are at least five main glaciers around Damavand summit which fall within the category of hanging glaciers (Fig. 9). There are four glaciers on the north face of Damavand, called Siuleh, Dubi-sel, Speleh and Khurtabsar. On the eastern side of Mount Damavand, in the massive Yakhar Valley, there is a large glacier, Yakhar Valley Glacier, which is the main source of Talkhrood River (Fig. 10). There are also several other glaciers on the south and southeast faces of Damavand which are relatively small, such as Kafar-Dareh Valley Glacier and Sardagh Glacier. Some of Damavand's glaciers cannot qualify as permanent glaciers, since they may disappear during a hot, dry summer.

The average snowline in this region is \sim 4400 m a.s.l. Due to the steep slopes of these glaciers, 26–40%, the thickness of ice does not exceed 20 m. The glaciers are covered with large penitents, 50–80 cm or more in height, and series of crevasses (Vaziri, 2003), though there is not much debris cover. Table 4 lists the main glaciers around Damavand summit, along with related glacial attributes.

Oshtorankuh region

In this glacier region, research was only conducted recently. Here we present the first results on the glaciers in the Oshtorankuh region.

Oshtorankuh region, with its highest peak, Sanboran summit, which rises to a height of ~4050 m a.s.l., is located 33°20′ N, 49°18′ E, 38 km from Doroud city in Lorestan province. There are several glaciers in this region that are formed on the north- and northeast-facing slopes, since the prevailing winds from the south and the intense solar radiation prevent snow accumulation on south-facing slopes (Vaziri, 2003). These glaciers provide a high



Fig. 11. Close-range photograph of Chal-kabud Glacier (July 1999). (Photograph taken by F. Vaziri.)

percentage of the drinking water in the surrounding urban regions and are the origins of several rivers; for instance, Chal-mishan, Chal-kabud (Fig. 11) and Chal-boran glaciers are the main sources of the Ab-mishan, Alvand Valley and Mirzayi rivers (Vaziri, 2003).

Due to the karstic nature of the glaciers' bedrock, several sinkholes are observed on the surface of the glaciers (Vaziri, 2003). Mechanisms of formation may include the gradual dissolution of the bedrock, in this case limestone, by percolating snowmelt water. Generally, in karstic glaciers, the ice thickness is affected by water circulation, but it also depends on annual precipitation.

The glacial attributes of the glaciers in the Oshtorankuh region (Fig. 12) were initially extracted from Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) imagery, but they were updated by processing aerial photographs. The results were also verified by extensive fieldwork. A list of the glaciers and their attributes is presented in Table 5.

Table 4. Physical attributes of glaciers around Damavand summit

Glacier	Orientation	Minimum height m	Maximum height m	Area km²	Maximum length km	Maximum width m
Siuleh	North	4000	5400	1	2.2	600
Dubi-sel	North	4100	5330	0.4	2	260
Speleh	North	4050	5400	0.5	2	400
Khurtabsar	North	3900	5070	0.5	1.7	300
Yakhar Valley	East	4000	5600	1	1.7	600

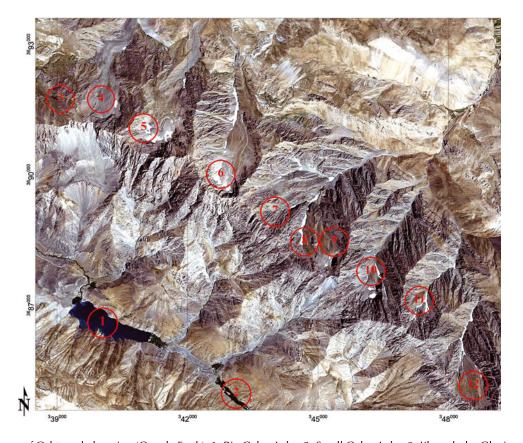


Fig. 12. SPOT image of Oshtorankuh region (Google Earth). 1. Big Gahar Lake; 2. Small Gahar Lake; 3. Kherseh dar Glacier; 4. Chal-mishan Glacier; 5. Chal-kabud Glacier; 6. Chal-boran Glacier; 7. Chal-fialeh-sun Glacier; 8. Shah-Takht Glacier; 9. Chal-piaro Glacier; 10. Chal-Homayun Glacier; 11. Chal-az-Nader Glacier; 12. Kul-Shakul Glacier.

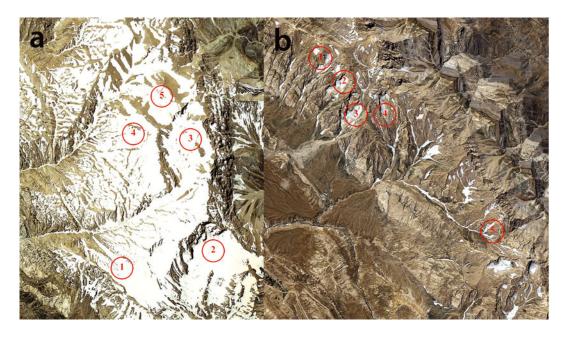


Fig. 13. SPOT image of Zardkuh region (Google Earth). (a) Glaciers around Joft-zarde and Zardkuh summits: 1. Chal-mishan Glacier; 2. Joft-zarde Glacier; 3. Khersan Glacier; 4. Pursunan Glacier; 5. Zardkuh Southeastern Glacier. (b) Glaciers around Sirdan summit: 1. Kuhrang Glacier; 2. Kuhrang Glacier; 3. Kuhrang Glacier; 4. Kuhrang Glacier; 5. Haft-tanan (Iluk) Glacier. Note that north is down in these images.

Zardkuh region

The Zardkuh mountain chain lies along the central Zagros mountain range in the province of ChaharMahal–Bakhtiari. The highest peak of this mountain chain, Zardkuh (Shah-Shahidan) summit (4548 m a.s.l.) is located at 32°22′ N, 50°04′ E.

There are several glaciers in this region, mostly cirque glaciers, from which two well-known rivers, Kuhrang and Bazoft (Karun river branches), originate. The greatest glacier concentration in this region is observed in two main areas: (1) around Joft-zarde and Shah-Shahidan (Zardkuh) summits (Fig. 13a) and (2) around Sirdan summit (Fig. 13b) (Rezaei, 2003). Since the underlying bedrock of these glaciers is karstic, several sinkholes are observed on the surfaces of the glaciers (Fig. 14a and b).

The attributes of these glaciers were initially extracted from aerial photographs, Indian Remote-sensing Satellite (IRS) and ASTER imagery. The results are also supported by fieldwork. The list of the Zardkuh region glaciers and their attributes is presented in Table 6.

GIS database

In this work we developed a GIS-based glacier database of Iran glaciers, carried out according to GLIMS standards. A GIS enables the automatic processing, integration, analysis and display of spatially oriented datasets (e.g. raster, vector, image), as well as storage and management of object-specific attributes, making it a preferred tool for conducting glacier inventories (Paul, 2003). GISs can contribute significantly to the monitoring and investigation of glaciers, since all the data are georeferenced and, consequently, facilitate management and control procedures.

Obtaining glacier attributes for a large number of glaciers by traditional methods is laborious and prone to errors (Paul, 2003). The use of GIS technology offers four essential improvements: (1) storage of two-dimensional glacier outlines in a vector layer for tracing geometry changes; (2) storage of corresponding glacier attributes in a relational database; (3) combination with a DEM for retrieval of 3-D glacier attributes and topographic indices; and (4) programming allows fully automatic processing of all data (Paul, 2003).

Table 5. Physical attributes of glaciers in the Oshtorankuh region

Glacier	Orientation	Minimum height	Maximum height	Area	Maximum length	Maximum width	
		m	m	km ²	km	m	
Kherseh dar	North	3380	3740	0.3	1	600	
Chal-mishan	Northeast	3180	3770	0.6	1	680	
Chal-kabud	Northeast	3600	3850	0.7	1	800	
Chal-boran	Northeast	3200	3700	1	1.8	600	
Chal-fialeh-sun	Northeast	3460	3600	0.2	0.9	500	
Shah-Takht	North	3100	3370	0.3	1.7	450	
Chal-piaro	North	3270	3380	0.4	1.4	400	
Chal-Homayun	Northeast	2860	3360	0.4	1.2	500	
Chal-az-Nader	Northeast	2800	3500	0.3	1.3	450	
Kul-Shakul	North	2880	3650	0.6	2	700	

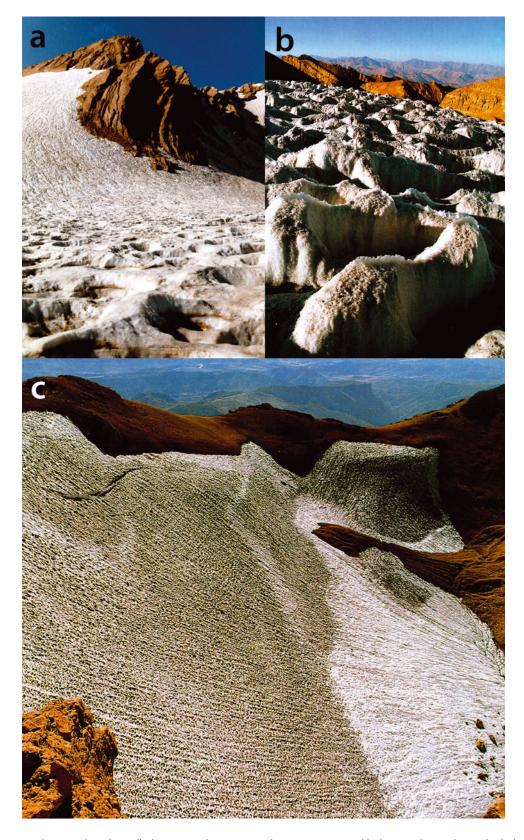


Fig. 14. Close-range photographs of Zardkuh region glaciers (October 1998): (a) sinkholes on the surface of Chal-mishan Glacier; (b) sinkholes on the surface of Zardkuh Southeastern Glacier; (c) Zardkuh Southeastern Glacier. (Photograph taken by F. Vaziri.)

Having performed the analysis on various data sources, such as aerial and satellite imagery, and also historic maps, output with a compatible format (ESRI shapefile) was imported into a GIS environment. The database of the glacier inventory includes topographic information and specific glaciological attributes, as well as the metadata. Different

spatial data layers, including digital boundaries of glaciers, DEMs, drainage basin limits, basin areas and georeferenced satellite imagery of different time intervals, constitute our GIS database. Attribute data were included, in accordance with GLIMS standards. In order to guarantee sustainable utilization of these resources for a wide range of scientific and

Table 6. Physical attributes of glaciers in the Zardkuh region

	Glacier	Orientation	Minimum height	Maximum height	Area	Maximum length	Maximum width
			m	m	km^2	km	m
Glaciers around Joft-zarde	Chal-mishan	North	3000	3800	0.9	2	1000
and Shah-Shahidan	Joft-zarde	Northeast	2900	3900	1.3	2	700
summits	Khersan	Northeast	3620	4060	1.8	2.3	700
	Pursunan	Northeast	3650	4000	0.4	1	400
	Zardkuh	Noutheast	3600	4100	0.6	1	400
	Southeastern Glacier						
Glaciers around Sirdan	Kuhrang (1)	Northeast	3300	3840	0.4	1	400
summit	Kuhrang (2)	Northeast	3000	3720	0.5	1.6	300
	Kuhrang (3)	Northeast	3090	3640	0.4	1.2	450
	Kuhrang (4)	Nnorth	3400	3770	0.4	1.2	500
Haft-tanan (Iluk) Glacier		Northeast	3540	3800	0.4	1.2	300

administrative applications, it is necessary to provide detailed standardized descriptions of the various datasets (Braun and others, 2001). In particular, these metadata records must comprise background information on data accuracy, and data sources should be compatible with a future data model for Iran glaciers (Braun and others, 2001). Hence, we created the metadata in the course of the work.

CONCLUSION

In this work we present the results of a first comprehensive inventory of Iran's glaciers. Five main glacier zones were identified and studied, using aerial and satellite imagery; related glacial attributes were extracted, with high accuracy, using different image-processing techniques. This work is significant; the first thorough study of Iran's glaciers using remote-sensing technology.

The glaciers are found in the higher elevations of the Alborz Mountains in the vicinity of Takhte–Soleiman region in the west, on Mount Damavand in the east, on the Sabalan Mountains in the northwest and on Zardkuh and Oshtorankuh Mountains in the Zagros mountain range in the southwest. Based on their geological structure, Iran's glaciers can be divided into two groups: (1) granite glaciers and (2) karstic glaciers. These two types of glaciers have different glacial features (e.g. the thickness of ice in karstic glaciers is less than that of granite glaciers). Moreover, special features, such as crevasses and moraines, are not found on karstic glaciers. The glaciers in the Takhte–Soleiman and Sabalan regions and around Damavand summit have the granite structure, whereas the Oshtorankuh and Zardkuh glaciers are karstic.

All the products from applying image-processing techniques on aerial and satellite imagery, including digital boundaries of glaciers, DEMs, drainage basin limits, basin areas and georeferenced satellite imagery, were incorporated into a GIS-based glacier inventory database.

A subject of ongoing research is to derive metric information on glacier flow using optical and synthetic aperture radar imagery. Future work will represent the results of surface displacements of Alamchal Glacier over a 2 year interval using high-resolution satellite imagery. We will use the results to update the Iran glacier database, providing more information about glacier dynamics and local climate change.

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