

CORRESPONDENCE

LOWER ORDOVICIAN TRILOBITES

SIR,—Since the March–April number of this Magazine was issued my attention has been drawn to a recent paper by F. Rasetti in the *American Journal of Science* (vol. 243, January, 1945) on the genera *Loganopeltis* and *Loganopeltoides* with a description and figure (p. 40, pl. 1, fig. 7) of the pygidium of his new species *Loganopeltis depressa* from the Lower Ordovician of Quebec. It seems to be almost identical with the specimen from the Shangort beds of Tourmakeady which was referred with considerable doubt to the genus *Cybelopsis* (*Geol. Mag.*, 82, 1945, p. 60), so that the generic reference may be altered to *Loganopeltis* with confidence.

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ON THE NORMAL FAULTING OF RIFT VALLEY STRUCTURES

SIRS,—In his recent very interesting paper on the above subject (*Geol. Mag.*, lxxxii, 1945, 37–44) Mr. H. G. Busk urges, rightly I believe, that normal faulting has played a greater part in rift tectonics than is often conceded. There are, however, certain of his statements on which I should like to comment.

On p. 42 he states that the rift valley movements were initiated on a broad peneplain; this peneplain is held to be of vast extent, and to include the great peneplain of the Northern Frontier District of Kenya.

While it is true that a great peneplain, usually referred to as the Miocene peneplain or the main peneplain, did extend over that part of Africa occupied by the Rift Zone, insufficient allowance has been made for the very large number of residuals and peneplain remnants that rest upon its surface in many areas. These remnants sometimes give rise to strong relief, and they are in places bounded by scarps eroded along ancient faults; such scarps (fault-line scarps) have sometimes been confused with Rift Valley fault-scarps. In the southern Rift zone at least there occur fault-line troughs, such as the Luangwa Valley, that have the form of the Rift Valleys, but long antedate them. Before progress can be made in the formulation of Rift theories based on scarps and plateaux, it is first necessary clearly to distinguish the true Rift scarps from the numerous other scarps of different age and origin.

The nature of a very large number of the Rift faults does not appear to be essentially different from that of the pre-Rift (post-Karoo) faults, which have always been accepted as normal faults, apart from some rare local disturbances.

Mr. Busk regards the great peneplain of the Northern Frontier District of Kenya as part of the main peneplain; it belongs, however, to a much younger cycle, which is separated from the main peneplain by scarps, up to 1,000 feet in height, that are usually erosion scarps, but sometimes possibly fault-line scarps. The main peneplain is represented by the southern Abyssinia plateau and the Uganda (Karamoja) peneplain.