

Texture-based classification of cloud and ice-cap surface features

W. G. REES AND I-I LIN

Scott Polar Research Institute, University of Cambridge, Lensfield Road, Cambridge CB2 1ER, U.K.

ABSTRACT. It is well known that the interpretation of high resolution (<100 m) visible and near infrared (e.g. Landsat) imagery of large ice masses is hindered by the uniform reflectivity of snow, ice and cloud surfaces. Such interpretation is at present largely performed manually, but there is a good prospect that it could be automated by the incorporation of image texture. This paper describes preliminary work towards the identification of the most appropriate texture technique, or combination of techniques, and assesses the likely performance of such methods.

Different textures are identified with different types of surface cover, and the use of these differences to classify images is investigated. Specifically, we compare a traditional texture measure, the Grey Level Co-occurrence Matrix (GLCM), with a modification of a relatively new technique, fractional Brownian motion (FBM). These two methods are applied to three Landsat MSS images of the Nordaustlandet ice cap, Svalbard. The classification accuracy, computation time and memory required, advantages and limitations of the two methods are compared. The GLCM technique appears to be able to distinguish three groups of image classes, namely dry snow, wet snow, and melt features, ablation areas or cloud cover. The FBM technique is computationally more efficient, and though it performs in general less well than the GLCM technique it gives better discrimination of cloud cover.

INTRODUCTION

The potential of Landsat-type images with high spatial resolution (<100 m) to reveal surface features on terrestrial ice masses has long been recognised (e.g. Swithinbank, 1988). Previous work has mostly been directed towards spectral classification combined with visual interpretation (e.g. Orheim and Luchitta, 1987; Dowdeswell and McIntyre, 1987). However, such analysis is hindered by the uniformity of reflectance of snow and cloud surfaces, which gives rise to a very limited range of tonal variation in such images. Texture analysis, which quantifies the spatial variation of image tone, has the potential to provide important complementary data (Welch and others, 1989, 1990; Key, 1990) and has already shown the value of texture analysis for identifying the presence (this is particularly difficult over snow and ice masses), and perhaps type, of cloud cover. It is not surprising that there should be considerable interest in the use of texture methods for discriminating between these image classes. Apart from the negative reason that it is difficult to effect a discrimination using only radiometric data, there is a physical justification for expecting the use of image texture to succeed, since bare ice, snow and cloud have distinct morphologies.

TECHNIQUES FOR THE QUANTIFICATION OF IMAGE TEXTURE

It is not easy to give simple definition of image texture

(Mather, 1987; Rees, 1990), since in principle it describes all spatial dependencies of the image. In practice, a finite number of texture parameters is defined. Typically these include the concepts of image roughness, homogeneity, linearity, contrast etc. However, the definition of these parameters often involves a subjective element so that it is possible to define a number of parameters, giving different numerical values, all of which are intended to quantify the same texture quality. Many different texture measures have been developed; the most widely used method is based on the Grey Level Co-occurrence Matrix (GLCM) (Haralick and others, 1973; Haralick, 1979). In this research, we compare the performance of this method with that of a new technique, based on the concept of fractals (Feder, 1987; Lin and Rees, 1992), in classifying three Landsat multispectral scanner (MSS) images of the Nordaustlandet ice cap, Svalbard.

Grey Level Co-occurrence Matrix

The GLCM technique has been described in detail by Haralick and others (1973). Here we provide a summary of the important aspects.

Texture properties are defined for a two-dimensional window in the image, and calculated as various properties of the GLCM measured within the window. If the image contains n grey levels (e.g. $n = 256$ for 8-bit data), the GLCM is an n by n square matrix \mathbf{S} in which the element $S_{ij}(\mathbf{v})$ is the frequency with which grey level i is separated from grey level j by a vector \mathbf{v} . The matrix \mathbf{S} is thus a function of the chosen displacement vector \mathbf{v} .

Many texture parameters can be extracted from the GLCM. The three which are studied in this work are the angular second moment (ASM), the contrast (CON), and the dissimilarity (DIS)

$$\begin{aligned}
 \text{ASM} &= \sum_{i=1}^n \sum_{j=1}^n S_{ij}^2 \\
 \text{CON} &= \sum_{i=1}^n \sum_{j=1}^n (i - j)^2 S_{ij} \\
 \text{DIS} &= \sum_{i=1}^n \sum_{j=1}^n |i - j| S_{ij}
 \end{aligned}$$

The angular second moment is essentially a measure of the homogeneity of the image, since it reaches its maximum possible value if the pixels within the image, or within the window under study, all have the same (arbitrary) grey level. It achieves its minimum possible value if the grey levels are uniformly distributed over the possible range of values, and uncorelated on the scale \mathbf{v} used in the analysis. The contrast and dissimilarity, on the other hand, measure the tendency of neighbouring pixels to have different values of grey level, and both become zero for homogeneous images. The contrast parameter assigns a greater weight to large differences in grey level than does the dissimilarity parameter.

The angular second moment is unaffected the image calibration, but the other two paramters are changed if the calibration is changed. If the calibration is changed such that a grey level of x becomes a grey level kx , the contrast increases by a factor of k^2 and the dissimilarity by a factor of k .

In this research, the window size is chosen as 24 by 36 pixels (corresponding to an area about 2 km by 2 km on the ground), and the vector displacement \mathbf{v} is chosen as (1,0), so that only variations in the x-direction of the image (vertically oriented features) are measured. This choice of \mathbf{v} is clearly somewhat arbitrary, though it is quite common in the use of the GLCM for the study of image texture. In principle, one should use all possible values of \mathbf{v} within the image window. However, the amount of computation required to generate the GLCMs and then to extract the texture paramters would be extremely large, and our result show that taking only $|\mathbf{v}| = 1$ gives good discrimination of several different image classes. The restriction to vertically oriented features poses a potential problem, though for features which are likely to be statistically isotropic such as those studied in the present work the problem is likely to be more apparent than real. Again, the restriction can be justified pragmatically on the grounds that the approach yields generally good discrimination.

Fractional Brownian Motion

The fractional Brownian motion (FBM) method is based on the concept of fractals, which are geometric constructions with fractional dimensionality (as distinct from the more conventional integral dimensionality of Euclidean geometry, in which lines have dimension 1, surfaces have dimension 2, and volumes have dimension 3). The application of this idea to the study of ice sheets has been discussed in detail by Rees (1992), who showed

using Landsat imagery that the surface topography of an ice cap may be described as a surface of dimension 2.12. Since, for the purposes of the present paper, the concept of fractional dimensionality merely provides a convenient framework in which to describe image texture, we refer the reader to the references given above, and other references contained therein, and present here only the main parameters.

The seimvariogram $\gamma(\mathbf{v})$ of an image is defined, following (e.g.) Curran (1988), as

$$\gamma(\mathbf{v}) = \frac{\langle [I(x) - I(x + \mathbf{v})]^2 \rangle}{2}$$

where $I(x)$ is the grey level of the pixel whose coordinate is x , and the average represented by $\langle \rangle$ is performed over all the pixels lying within a window of the image. The windows were chosen to be identical to those used for the GLCM analysis. A semivariogram is formed by evaluating $\gamma(\mathbf{v})$ for vector displacements \mathbf{v} along a one-dimensional transect.

For surfaces which display true fractional Brownian motion,

$$\gamma \propto \mathbf{v}^{4-2D}$$

where D is a constant known as the fractal dimension. For real surfaces this relationship is expected to break down for some minimum and some maximum value of \mathbf{v} . The technique we adopt, described in greater detail in Lin and Rees (1992), is to plot the $\gamma(\mathbf{v})$ relationship as $\log(\gamma)$ against $\log(\mathbf{v})$ and to use an iterative procedure to define the optimum straight line through the points, and the slope of this line, from which D can be determined. We also define a parameter which we call the "shift", which is the extrapolated value of $\log(\gamma)$ when $\log(\mathbf{v}) = 0$. It thus corresponds to the image variance for a single pixel.

DATA SET

We have compared the performance of the GLCM and FBM texture methods by using them to study three Landsat MSS images of Nordaustlandet, Svalbard, for which image classifications were known. Three images were needed to give examples of all the image classes (discussed below) which we wished to study. Image 1 is band 1 (0.5 to 0.6 μm) of the Landsat 1 image from path 233, row 002, acquired on 25 March 1973. This image (which is illustrated in Lin and Rees, 1992) shows a dry snow surface. Image 2 is band 4 (0.8 to 1.0 μm) of the Landsat 2 image from path 233, row 002, acquired on 3 July 1976. This image is partly covered by thin cirrus cloud (Fig. 1). Image 3 is band 1 of the Landsat 2 image from path 234, row 002, acquired on 1 August 1981. This image shows various melt features, as well as areas of wet snow (Fig. 2). It would have been better to be able to use the same spectral band throughout our analysis, but (a common problem with 10 or 12 year old data on magnetic tape) some channels had become corrupted. However, the very high correlation between visible and near infrared bands in images of snow and cloud means that this is unlikely to have introduced any significant errors into our study of image texture. A very simple calibration was performed on each image, by scaling the

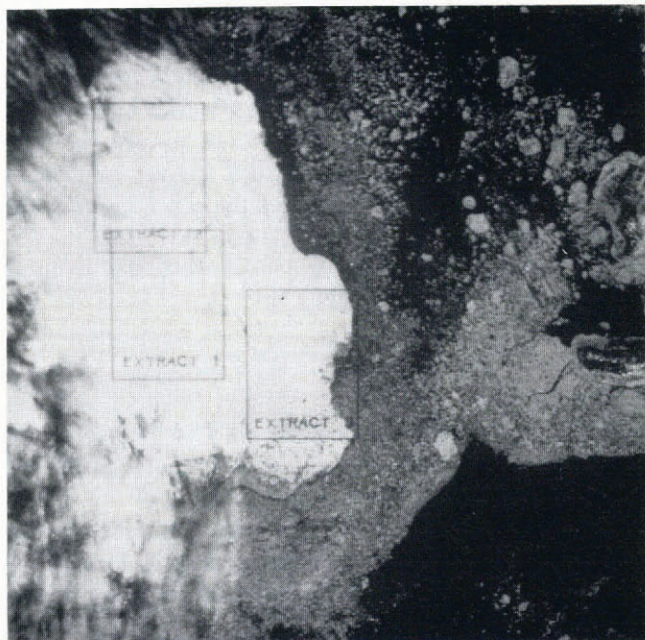


Fig. 1. Image 2 (Landsat 2, Band 4, July 1976). The image shows an area approximately 185 km square.

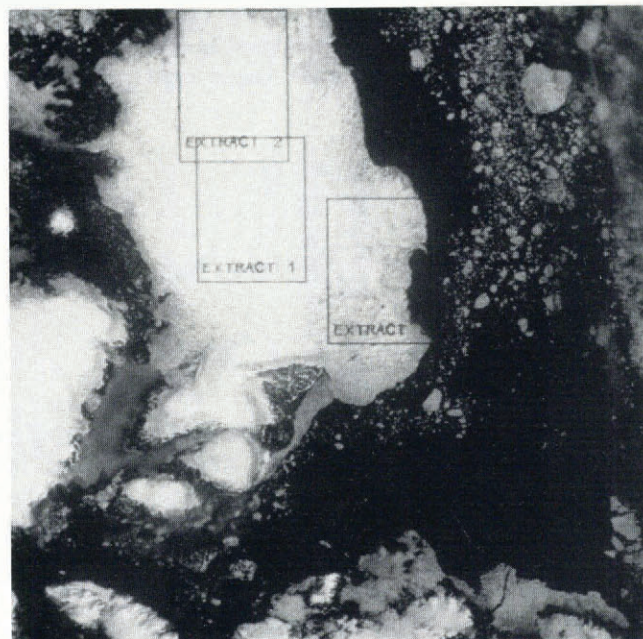


Fig. 2. Image 3 (Landsat 2, Band 1, August 1981). The image shows an area approximately 185 km square.

total range of grey levels to the range 0 to 63 (i.e. rescaling them to 6-bit data).

164 windows of 24 by 36 pixels were defined within three extracts of these images, shown in Figures 1 and 2 and selected to contain areas identified as dry snow, wet snow, melt areas (distinguishable from wet snow by the appearance of melt streams and melt ponds), ablation areas, and cloud cover. These identifications were performed by visual interpretation and by comparison with Dowdeswell 91984) and G.S. Hamilton (personal communication). The texture parameters defined above were calculated for each window.

RESULTS

GLCM technique

Values of the angular second moment (ASM), contrast (CON) and dissimilarity (DIS) were calculated for a total

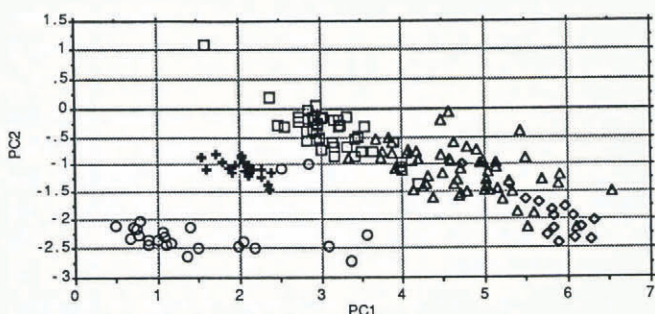


Fig. 3. Principal components derived from the GLCM analysis. The circles represent dry snow, the squares wet snow, the triangles melt areas, the lozenges ablation areas and the crosses cloud.

of 1643 windows in the three images, representing all five image classes. The results showed significant correlation between ASM, CON and DIS, and values spread over wide ranges, so a principal components analysis (e.g. Rees, 1990) was performed on the results in order to generate uncorrelated texture measures. Before performing this analysis, the data were transformed to their logarithms to produce distributions which were more nearly normal.

The three principal components were found to be (ignoring additive constants)

$$PC_1 = +0.720 \log(ASM) - 0.580 \log(CON) - 0.382 \log(DIS)$$

$$PC_2 = +0.687 \log(ASM) + 0.674 \log(CON) + 0.272 \log(DIS)$$

$$PC_3 = +0.100 \log(ASM) - 0.458 \log(CON) + 0.883 \log(DIS)$$

The third principal component contained less than 0.2% of the total variance and was ignored. Figure 3 shows the first two principal components, plotted as a cluster diagram. It can be seen from the figure that wet snow can be distinguished from other classes by the value of PC_1 , and that dry snow can be distinguished from other classes by the value of PC_2 . Other separations are less obvious, but we can quantify them as follows. We can define a simple index of separability S for two probability distributions $p_1(x)$ and $p_2(x)$ as

$$S = 1 - \frac{\int p_1(x)p_2(x)dx}{\sqrt{\int p_1^2(x)dx \int p_2^2(x)dx}}$$

such that $S = 0$ if the two distributions are identical, and $S = 1$ if they do not overlap. Separability indices are calculated for each of the ten possible pairs of image

classes, using first PC_1 and then PC_2 , and the higher value of S is chosen. A matrix of separability values can then be defined:

	Melt	Ablation	Wet	Cloud
Dry	0.91	1.00	1.00	0.92
Melt		0.70	0.94	0.74
Ablation		1.00	0.67	
Wet				0.99

If we assume that S must be greater than about 0.8 for an effective discrimination, we can see from this table that the GLCM technique provides poor discrimination between melt areas, ablation areas and cloud cover, but that other classifications and discriminations are unambiguous.

Fractal technique

Values of the fractal dimension D and the shift parameter were calculated for 432 transects of the three images, again chosen to represent each of the five image classes. Figure 4 shows the resulting histograms for each class. It can be seen from the figure that wet and dry snow surfaces have very similar, and high, values of D . Ablation areas show somewhat smaller values of D , and melt areas and cloud cover show very broad distributions. The shift parameter is lowest for cloud, intermediate for ablation areas, and highest for wet snow, dry snow and melt areas.

Separability statistics have also been calculated for the 5 classes on the basis of the distributions of D and the shift parameter, again taking the higher calculated value of S for each pair of classes, with results as follows:

	Melt	Ablation	Wet	Cloud
Dry	0.53	0.90	0.02	0.88
Melt		0.33	0.57	0.94
Ablation			0.98	0.84
Wet				0.95

The analysis shows that the use of the fractal dimension D alone is a rather poor discriminant of image class, although with the inclusion of shift data the

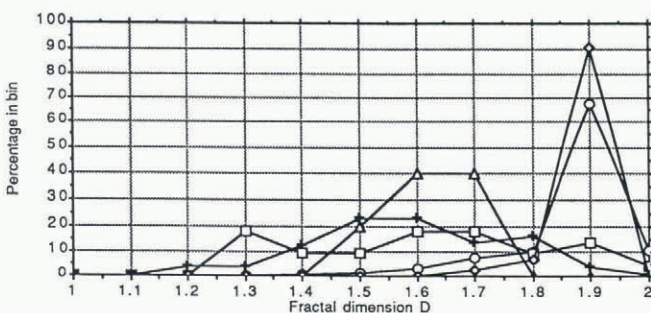


Fig. 4. Histograms of the values of (a) fractal dimension D and (b) the shift parameter, from the fractional Brownian motion analysis. The symbols are as in Figure 3.

number of ambiguous discriminations is reduced to four, namely dry snow from wet snow and melt areas, and melt areas from ablation areas and wet snow. If the results of the FBM and GLCM separability analyses are combined, we see that the number of ambiguous discriminations can be reduced to one, namely that between melt areas and ablation areas.

DISCUSSION

Our analysis shows that the GLCM technique is, in general, able to make a more effective discrimination of image classes than is the FBM technique, although the latter achieves a greater discrimination between cloud and melt areas, and between cloud and ablation areas. However, it is less profitable to think of the techniques as rivals, but rather as complementary methods which, when combined, achieve a high level of discrimination in virtually all cases.

It is useful to compare the computational efficiencies of the two texture algorithms. For a window of size $A \times B$ pixels ($A \approx B$), the number of operations required to calculate the GLCM S_{ij} is approximately AB as long as the modulus of the vector \mathbf{v} is much less than A or B . If the total number of possible grey levels is n , the number of computations required to calculate each of the three texture parameters defined in section 2 is n^2 . Thus the total number of operations required to generate the set of three texture parameters is approximately $AB + 3n^2$. Taking $A = 24$, $B = 36$ and $n = 256$ gives approximately 2×10^5 operations. However, in practice, the range of grey levels present in Landsat images of snow, ice or cloud is very restricted, and it is more realistic to take $n \approx 10$. In this case, the number of operations required to generate the GLCM parameters is closer to 10^3 .

For a transect of length A pixels, the number of operations required to generate a semivariogram with m points is approximately mA , and the number of operations required to perform the iterative curve fitting and the extraction of the texture parameters defined in section 2 is approximately $3m^2$. The total number of operations required is about $mA + 3m^2$. Taking $A = 30$ and $m = 12$ (typical values for the size of window employed in this work) gives a total of about 10^3 operations, so that the FBM and GLCM techniques should in practice require similar amounts of computation. This was confirmed experimentally during the course of this work. However, we stress that, in principle, the calculation of GLCM-based texture measures can involve up to 100 times as much computation as the FBM measures, and that the complete classification of an image by this method could become prohibitive, requiring up to about 10^9 operations.

We emphasise the preliminary nature of the work described in this paper. It is open to criticism on the grounds that we have used data from several images and two spectral bands, rather than just one of each, which raises the possibility that differential calibration and sun angle effects may affect our results. It must also be stressed that the use of texture analysis in image classification should not be relied upon until the physical mechanisms responsible for generating the image texture are clearly

understood. At present this can be done only in a very broad and qualitative way by recognizing that different surface morphologies will give rise to different image textures, but we are currently attempting to put this understanding on a quantitative basis. For example, the high value of the fractal dimension D measured for wet and dry snow areas can be understood since in these areas variations in image brightness are likely to be caused predominantly by topographic variation (Rees and Dowdeswell, 1988), and this is small in the areas studied (Dowdeswell and McIntyre, 1987). Lower values of D for melt areas probably reflect real variations in the surface structure at scales of less than the image window size of 2 km. If an understanding at this level can be quantified and extended to include the GLCM parameters, it should allow questions such as the optimum spatial resolution and window size, and the optimum spectral band (or combination of bands) to be determined. However, we believe that the general conclusion from this work will continue to stand, namely that the use of texture techniques (probably a combination of GLCM and fractal techniques) has the potential to provide unambiguous classifications of images of snow, ice and cloud, and that this classification can ultimately be automated.

CONCLUSIONS

We have applied two texture measures to three temporally separated Landsat MSS images of the Nordaustlandet ice cap. The results show that different image classes have distinct textures, and confirm that texture analysis provides powerful additional information for the classification of such images. The results show that the use of the three GLCM texture parameters (angular second moment, contrast and dissimilarity) gives in most cases greater discrimination of surface classes than the use of the fractional Brownian motion technique. However, the two techniques appear to have different strengths and weaknesses, and their combination yields a high level of discrimination of all the image classes studied.

ACKNOWLEDGEMENTS

We gratefully acknowledge the technical assistance of Drs B. J. Devereux and J. A. Dowdeswell, and of Messrs O. Dikshit, M. R. Gorman, T. Mayo and G. Hamilton. The comments of Drs R. Bindschadler and R. A. Massom, and

of an anonymous referee, have greatly improved the presentation of this work.

REFERENCES

- Curran, P. J. 1988. The semivariogram in remote sensing. *Remote Sensing Environ.*, **24**, 493–507.
- Dowdeswell, J. A. 1984. *Remote sensing studies of Svalbard glaciers*. (Ph.D. thesis, University of Cambridge.)
- Dowdeswell, J. A. and N. F. McIntyre. 1987. The surface topography of large ice masses from Landsat imagery. *J. Glaciol.*, **33**, 16–23.
- Feder, J. 1987. *Fractals*. New York, Plenum Press.
- Haralick, R. M., K. Shanmugam and L., Dinstein. 1973. Textural features for image classification. *IEEE Trans. systems, management & cybernetics*, **SMC3**, 610–621.
- Haralick, R. M. 1979. Statistical and structural approach to texture. *Proc. IEEE*, **67**, 786–804.
- Key, J. 1990. Cloud cover analysis with Arctic Advanced Very High Resolution Radiometer data 2. Classification with spectral and textural measures. *J. Geophys. Res.*, **95**, 7661–7675.
- Lin, I. and W. G. Rees. In press. A new fractal texture classification of cloud and ice cap surface features from Landsat imagery. *Proceedings IGARSS '92*.
- Mather, P. M. 1987. *Computer processing of remotely-sensed images*. Chichester, etc., John Wiley.
- Orheim, O. and B. K. Luchitta. 1987. Snow and ice studies by Thematic Mapper and Multispectral Scanner Landsat images. *Ann. Glaciol.*, **9**, 109–118.
- Rees, W. G. 1990. *Physical principles of remote sensing*. Cambridge, Cambridge University Press.
- Rees, W. G. 1992. Measurement of the fractal dimension of ice-sheet surfaces using Landsat data. *Int. J. Remote Sensing*, **13**, 663–671.
- Rees, W. G. and J. A. Dowdeswell. 1988. Topographic effects on light scattering from snow. *Proc. IGARSS'88 Symp. ESA SP-284* 161–164.
- Swithinbank, C. 1988. Landsat images of Antarctica. *U.S. Geol. Surv. Prof. Pap.* 1386B.
- Welch, R. M., M. S. Navar and S. K. Sengupta. 1989. The effect of spatial resolution upon texture based cloud field classifications. *J. Geophys. Res.*, **94**, 14,767–14,781.
- Welch, R. M., K-S Kuo and S. K. Sengupta. 1990. Cloud and surface textural features in polar regions. *IEEE Trans. Geosci. Remote Sensing*, **28**, 520–528.
- The accuracy of references in the text and in this list is the responsibility of the authors, to whom queries should be addressed.*