



Article

Marchettiite, (NH₄)C₅H₃N₄O₃, a new organic mineral from Mount Cervandone, Devero Valley, Western–Central Alps, Italy

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ABSTRACT

The new mineral marchettiite (IMA2017-066) is the natural equivalent of ammonium hydrogen urate. It has a simple molecular formula $C_5H_7N_5O_3$ and can be alternatively written as (NH₄) $C_5H_3N_4O_3$. Marchettiite was found in a cleft at Mount Cervandone, Devero Valley, Piedmont, Italy, where it occurs as aggregates of opaque pale pink to white, platy prismatic crystals. This mineral has a white streak, dull and opaque lustre, it is not fluorescent and has a hardness of 2–2.5 (Mohs' scale). The tenacity is brittle and crystals have a good cleavage parallel to {001}. The calculated density is 1.69 g/cm³. Marchettiite is biaxial (–) with 2V of 47.24°; the optical properties of marchettiite were determined by periodic-DFT methods providing the following values: $\alpha = 1.372$, $\beta = 1.681$ and $\gamma = 1.768$. No twinning was observed. Electron microprobe analyses gave the following chemical formula: $C_{4.99}H_{6.97}N_{4.91}O_{3.00}$. Although the small crystal size did not allow refinement of structural data by single-crystal diffraction, we were able to refine the structure by powder micro X-ray diffraction. Marchettiite has space group $P\bar{1}$ and the following unit-cell parameters: a = 3.6533(2) Å, b = 10.2046(7) Å, c = 10.5837(7) Å, $\alpha = 113.809(5)^\circ$, $\beta = 91.313(8)^\circ$, $\gamma = 92.44(1)^\circ$ and V = 360.312 Å³. The strongest lines in the powder diffraction pattern [d in Å (I)(hkl)] are: 9.784(50)(001); $8.663(80)(01\bar{1})$; 5.659(100)(011); $3.443(100)(10\bar{1})$; 3.241(70)(003) and $3.158(100)(1\bar{1}\bar{1})$. Marchettiite is named after Gianfranco Marchetti, the mineral collector who found this mineral.

Keywords: ammonium hydrogen urate, Devero valley, marchettiite, Mount Cervandone, new mineral

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Introduction

Mount Cervandone, in the Devero Valley, and the Swiss Wannigletscher (also named Cherbadung) are the type localities of rare earth element (REE) arsenates and arsenites including asbecasite (Graeser, 1966), cafarsite (Graeser, cervandonite-(Ce) (Armbruster et al., 1988), fetiasite (Graeser et al., 1994), gasparite-(Ce) (Graeser and Schwander, 1987) and paraniite-(Y) (Demartin et al., 1994). Mount Cervandone is also the type locality of deveroite-(Ce), a natural REE-oxalate (Guastoni et al., 2013). Marchettiite (Fig. 1) is the eighth type mineral and the seventh organic mineral found in the alpine fissures of Mount Cervandone and Wannigletscher after deveroite-(Ce) (Guastoni et al., 2013), humboldtine (Guastoni et al., 2015), lindbergite (Albertini and Meisser, 2012), moolooite (Guastoni, 2009), tinnunculite (Guastoni et al., 2020) and whewellite (Guastoni et al., 2015).

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In this study, we describe a new mineral named marchettiite, ideally $C_5H_7N_5O_3$, found at Mount Cervandone, Devero valley, Western-Central Alps, Italy. Marchettiite is an organic compound collected in 1995 in an alpine cleft by Gianfranco Marchetti, a local mineral collector. The mineral is stable in natural atmospheric conditions and at room temperature.

Marchettiite has no natural analogue, powder micro-X-ray diffraction and electron microprobe analysis highlighted that is the natural analogue of the synthetic ammonium hydrogen urate compound (Friedel *et al.*, 2006; Sekkoum *et al.*, 2016); also known as the anhydrous ammonium salt of uric acid. The simple molecular formula $C_5H_7N_5O_3$ can be written alternatively as $(NH_4)C_5H_3N_4O_3$ (Tettenhorst and Gerkin, 1999).

The new mineral and the name (symbol Mht) have been approved by the Commission on New Minerals, Nomenclature and Classification (CNMNC) of the International Mineralogical Association (IMA2017-066, Guastoni *et al.*, 2017). The holotype material is deposited in the mineralogical collection of the Museum of Mineralogy of the Department of Geosciences at the University of Padova (Italy), under the catalogue number MMP M171892. The name is for Gianfranco Marchetti (1943–2013), discoverer of the mineral and amateur mineralogist.

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Fig. 1. Optical microphotograph of marchettiite holotype (Museum of Mineralogy of the Department of Geosciences at the University of Padova, catalogue number MMP M171892), the long side of the picture is 6 mm (photo: Stefano Castelli).

Occurrence

Mount Cervandone is composed of fine-grained to very finegrained two-mica leucocratic gneisses belonging to the lower-Penninic, Mount Leone-Arbola' nappe (Dal Piaz, 1975; Steck, 2008). The Mount Leone nappe emplacement took place in a regime of ductile deformation and contains 3T-polytype phengite relicts overprinted by the regional amphibolite-facies metamorphism (Hammerschmidt and Frank, 1991). In this area, REE bearing-minerals include several arsenites, arsenates, phosphates, oxides and deveroite-(Ce), a natural REE-oxalate (Guastoni et al., 2006; Guastoni et al., 2010; Guastoni et al., 2013). REE-bearing minerals, with the exception of the organic compound deveroite-(Ce), are related to the presence of swarms of sub-horizontal dykes with pegmatitic textures, which crop out for hundreds of metres along the slopes of Mount Cervandone and Punta Marani. These dykes were involved with regional amphibolite-facies metamorphism and underwent ductile deformation developing folds or filling gaps related to foliation boudinage of the main schistosity. These dykes are composed of coarse grained vitreous colourless or smoky quartz, K-feldspar and silvery-greenish micas. Pegmatite dykes also contain accessory minerals including centimetric up to decimetric black tourmaline prisms and curved mica flakes embedded in vitreous 'smoky' quartz which exhibit boudinage-like textures. Occasionally, centimetric pale sky-blue beryl prismatic crystals were also found. Pegmatitic dykes are locally cut by Alpine-type fissures which contain a rich assemblage of typical Mount Cervandone minerals as reported above.

The formation of marchettiite requires the presence of water as the fluid medium. This natural anhydrous ammonium hydrogen urate is not related to any amphibolite-facies metamorphism or pegmatite dykes but was formed by circulation of meteoric waters at low temperatures enriched with soluble altered biological material (Hazen *et al.*, 2013).

Experimental methods

Electron microprobe analysis

Semi-quantitative electron microprobe analyses (EMPA) were performed on the same crystal fragment used for the micro-X-ray diffraction analysis at the Dipartimento di Scienze Geologiche e

Table 1. Semi-quantitative chemical data (by wt.%) for marchettiite (n = 15).

Constituent	Mean	Range	Probe standard
С	32.3	30.4-34.3	Calcite (Iceland spar)
N	37.2	34.5–39.5	PBN, pyrolitic boron nitride (Union Carbide Corp., USA)
*H	3.8	-	
*0	25.9	-	
Total	99.2		

Notes: *H, *O, calculated by stoichiometry (Friedel et al., 2006)

Geotecnologie of the University of Milano-Bicocca, Italy. The chemical composition was determined using a SEM Tescan VEGA TS 5136 XM equipped with energy-dispersive analysis (EDS) by means of an EDAX GENESIS 4000 XMS electron microprobe. Major and minor elements were determined at 20 kV accelerating voltage and 190 pA (in a Faraday cup), spot size = 250 nm, 100 s as counting time and working distance = 23 mm. For EDS semi-quantitative analyses the following lines were used: $CK\alpha$, $NK\alpha$. C and N were measured (in wt.%) on a crystal fragment that had been polished and Au coated using an Edwards Sputter Coater S150B, under an Argon atmosphere at 6.5 mbar, operating at 40 mA/1 kV, for 10 s in order to obtain an estimated 2.5 nm of Au coating.

The mean analytical results are given (Table 1). The empirical formula, based on 3O atoms per formula unit (apfu) obtained from semi-quantitative EMPA (with H_2O and C_2O_3 calculated by stoichiometry, due to the very small quantity of material available) is: $C_{4.99}H_{6.97}N_{4.91}O_{3.00}$. The ideal formula is $C_5H_7N_5O_3$, or alternatively written as (NH₄)C₅H₃N₄O₃, which requires: C 32.44, H 3.81, N 37.83, O 25.92, total 100.00 wt.%.

Optical properties

The refractive indices were computed from the crystal structure using the periodic coupled-perturbed Kohn Sham method as implemented in *CRYSTAL17* (Dovesi *et al.*, 2018; Ferrero *et al.*, 2008) using the B3LYP/pob-TZVP_rev2 (Vilela Oliveira *et al.*, 2019) level of theory with the shrinking factor set to 8.

X-ray diffraction

Powder micro X-ray diffraction (PXRD) data were recorded on a Rigaku-Oxford Diffraction Supernova diffractometer equipped with a micro X-ray source (beam diameter = 0.12 mm and λ = Mo $K\alpha$) and a 200K Pilatus detector (Dectris). An aggregate of ~0.100 mm was glued on a borosilicate fibre. The sample-to-detector distance was 68 mm. The experimental PXRD profile was extracted with *Crysalis Pro* (Rigaku Oxford Diffraction), and further analysed with the programs *Highscore Plus* (Panalytical) and *GSAS 2* (Toby and Von Dreele, 2013). Although the small crystal sizes did not allow us to refine structural data, it was possible to obtain unit-cell parameters, which are in good agreement (Table 2) with those reported for synthetic anhydrous ammonium acid urate (Friedel *et al.*, 2006) and to perform a qualitative validation of the structural model.

The limited data quality did not allow for *ab initio* structure solution. Nevertheless, starting from the structural model proposed by Friedel *et al.* (2006), we have performed a rigid-body Rietveld refinement which converged to a meaningful structure. Our refined structure differs from the published one for the

Table 2. Unit-cell parameters from X-ray diffraction data for marchettiite in comparison with the synthetic analogue.

	a (Å)	b (Å)	c (Å)	α (°)	β (°)	γ (°)
Marchettiite Anhydrous ammonium acid urate (Friedel <i>et al.</i> , 2006)	3.6533(2)	10.2046(7)	10.5837(7)	113.809(5)	91.313(8)	92.44(1)
	3.650(3)	10.215(4)	10.597(5)	113.9(1)	91.1(1)	92.3(1)

Table 2. nafeasite. Only calculated lines with intensities > 1.5 are listed.

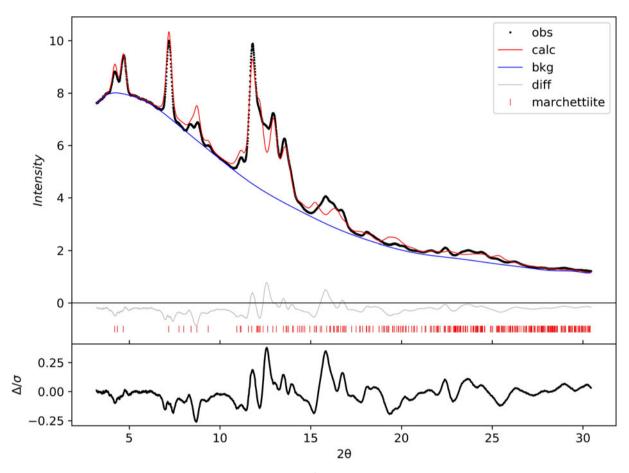


Fig. 2. Rigid-body Rietveld refinement of marchettiite against PXRD data ($\lambda = 0.71073 \text{ Å}; R_p = 0.032; R_{wp} = 0.045; R_{Brogg} = 0.130; R_{F2} = 0.148$).

absence of any excess $\mathrm{NH_3}$ in our structural model, in agreement with the chemical analysis. A second difference is in the intramolecular geometry, as we have maintained the planarity expected for the flat and rigid monodeprotonated urate anion. Though the fitting of the experimental pattern is not ideal, and the model cannot be reliably refined further without overparameterising, the fair agreement between observed and calculated pattern (Fig. 2) is further evidence supporting the identification of marchettiite as the natural analogue of ammonium hydrogen urate. Observed d values and intensities are given in Table 3, atomic coordinates in Table 4 and bond distances in Table 5. The crystallographic information file has been deposited with the Principal Editor of *Mineralogical Magazine* and is available as Supplementary material (see below). The crystal structure of marchettiite is showed in Fig. 3.

Micro-Raman spectroscopy

The micro-Raman spectrum of marchettiite was collected at the Department of Chemical Sciences of the University of Padova using a Raman DXR (ThermoScientific) with a laser wavelength of $532~{\rm cm}^{-1}$ (power 8 mW). The spectrum was collected in the range of $200{\text -}3500~{\rm cm}^{-1}$ with a spectral resolution of $1.9~{\rm cm}^{-1}$ and a spatial resolution of 1 μ m. The Raman spectrum is shown in Fig. 4a (which also reports the Raman band values) and compared with the synthetic commercial anhydrous ammonium hydrogen urate (Santos *et al.*, 1984) in Fig. 4b.

Infrared spectroscopy

The Fourier-transform infrared (FTIR) spectrum was acquired at the Department of Ingegneria Industriale of the University of Padova by FTIR-ATR technique using a Nicolet iS50 FTIR (Thermo Fisher Scientific) equipped with diamond crystal. The spectrum was collected in the range of 4000–650 cm⁻¹ with a resolution of 4 cm⁻¹. The FTIR spectrum with peak values of marchettiite is given in Fig. 5; literature data are summarised in Table 6 (Benramdane *et al.*, 2008 and Sekkoum *et al.*, 2016). From the FTIR results, it is evident that our sample of marchettiite is not pure material but is probably intermixed with whewellite.

Table 3. Powder X-ray diffraction data (d in Å) for marchettiite.*

I _{rel}	<i>d</i> meas. (Å)	I _{rel, calc.}	$d_{\mathrm{calc.}}$ (Å)	hkl
24	9.676	18	9.674	001
37	8.685	26	8.700	011
75	5.666	65	5.663	011
9	4.854	8	4.837	002
17	4.642	13	4.661	020
3	4.324	6	4.350	0 2 2
12	3.643	12	3.647	100
38	3.495	25	3.520	013
100	3.452	100	3.463	101
54	3.358	43	3.333	111
11	3.265	11	3.289	032
30	3.248	6	3.225	003
51	3.150	57	3.153	1 Ī Ī
13	3.116	11	3.107	030
49	3.010	27	3.018	211
17	2.940	10	2.954	1 2 0
7	2.643	3	2.612	$0.1 \bar{4}$
22	2.576	5	2.540	0 4 2
15	2.508	6	2.509	041
17	2.442	1	2.449	034
7	2.253	1	2.255	133
3	2.195	<1	2.175	$0.4 \bar{4}$
3	2.078	<1	2.082	0 4 1
3	2.026	<1	2.021	$14\bar{1}$
2	1.964	<1	1.959	1 3 2
3	1.893	<1	1.893	$0.5\bar{4}$
7	1.827	1	1.829	0 4 2
3	1.752	<1	1.750	036
6	1.725	<1	1.719	1 3 3
7	1.678	2	1.675	$2\bar{1}\bar{2}$
3	1.624	2	1.623	150
3	1.563	<1	1.562	155
2	1.476	<1	1.479	213
3	1.422	<1	1.426	2 0 4
2	1.317	<1	1.313	0 4 8
2	1.154	<1	1.157	1 2 7
2	1.069	<1	1.069	028

^{*}The strongest lines are given in bold

Results and discussion

Physical and optical properties

Marchettiite occurs as aggregates of opaque pale pink to white, elongated, platy, prismatic crystals (the actual morphology can be determined only by SEM). Indeed, marchettiite shows no monocrystals (as demonstrated by X-ray diffraction) and is always found in microcrystalline aggregates in a cleft of fine grained orthogneiss. Marchettiite shows no inclusions of, or evident intergrowths with, other minerals. The maximum size of the marchettiite aggregates is 3 mm (Figs 1 and 6). Marchettiite is associated with albite, K-feldspar, muscovite, quartz and whewellite. It has a white streak, a dull and opaque lustre, is not fluorescent, and neither pleochroism nor twinning were observed. Due to the extremely limited thickness of the crystal hardness is only estimated as 2-2½ on the Moh scale. Attempts of micro-indentation on a very thin crystal only provided unreliable data. The tenacity is brittle and the crystals have a good cleavage, parallel to {001}. Due to the small grain size, it was impossible to measure the density. The calculated density using the empirical formula and our unit-cell data is equal to 1.69 g/cm³, in good agreement with the synthetic analogue, which is 1.772 g/cm³ (Tettenhorst and Gerkin, 1999).

Optical properties were determined by using the periodic – DFT method (details are reported in the Experimental methods

Table 4. Fractional atomic coordinates and isotropic displacement parameters $(\mathring{\mathbb{A}}^2)$ for marchettiite. Hydrogen urate and ammonium cations were treated as rigid bodies, hence standard uncertainties are reported only for the reference points and not for the individual atomic coordinates. For hydrogen urate, the geometrical centre of the 6-membered ring, indicated as point 'A', was used as the reference point. For ammonium, N10 was used as reference point. A global isotropic displacement parameter was refined for each rigid body.

	x/a	y/b	z/c	Global U _{iso}
Hydroge	en urate			
Α	0.5115(16)	0.6984(12)	0.3711(9)	0.056(2)
N1	0.5579	0.5839	0.4087	
H1	0.5943	0.4979	0.4379	
C2	0.6749	0.7227	0.4983	
02	0.8183	0.7442	0.6096	
N3	0.6278	0.8367	0.4606	
C4	0.4645	0.8104	0.3329	
C5	0.3500	0.6796	0.2447	
C6	0.3935	0.5574	0.2808	
06	0.2899	0.4371	0.2020	
N7	0.1970	0.6786	0.1245	
H7	0.0835	0.5913	0.0366	
C8	0.2334	0.8237	0.1514	
08	0.1300	0.8711	0.0696	
N9	0.3984	0.9102	0.2796	
H9	0.4578	1.0221	0.3247	
Ammon	ium			
N10	0.139(7)	0.844(3)	0.849(3)	0.076(2)
H11	0.027	0.789	0.753	
H12	0.120	0.774	0.894	
H13	0.410	0.854	0.835	
H14	0.024	0.939	0.903	

 $\textbf{Table 5.} \ \ \text{Interatomic distances (Å) for marchettiite. The values reflect the unrefined intramolecular distances of the rigid bodies.}$

Hydrogen urate			
N1—C2	1.391	C5—C6	1.458
N1—C6	1.385	C4—C5	1.325
N3-C2	1.390	C2-O2	1.210
N3-C4	1.383	C6—O6	1.210
C4-N9	1.375	C8-O8	1.210
C5-N7	1.374		
C8-N7	1.390	H1-N1	1.05
C8-N9	1.391	H7—N7	1.06
Ammonium			
<n10—h></n10—h>	1.01		

below). Marchettiite is biaxial (-), with $\alpha = 1.372$, $\beta = 1.681$, $\gamma = 1.768$ and $2V = 47.243^{\circ}$. This relatively new approach for minerals was adopted as we realised that, at least at the scale of the FTIR analyses (see sections below), it seemed evident that marchettiite was intermixed with a probable whewellite impurity and, at the same time, considering the polycrystalline nature of marchettiite it was actually impossible to obtain any reliable optical properties experimentally.

The refractive indices of marchettiite are in satisfying agreement with those of tinnunculite ($C_5H_4N_4O_3\cdot 2H_2O$), of: $\alpha=1.503(3)$, $\beta=1.712(3)$, $\gamma=1.740(1)$, biaxial (–) and $2V_{obs}=40(10)^\circ$ (Pekov *et al.*, 2017) and its synthetic counterpart, uric acid dihydrate with: $\alpha=1.508(3)$, $\beta=1.691(3)$, $\gamma=1.728(3)$, biaxial (–) and $2V_{obs}=40.4^\circ(6)$ (Ringertz, 1965).

Marchettiite has no natural non-biological analogue, as mentioned above, however there is a synthetic analogue (Friedel et al., 2006; Sekkoum et al., 2016). Other related natural organic

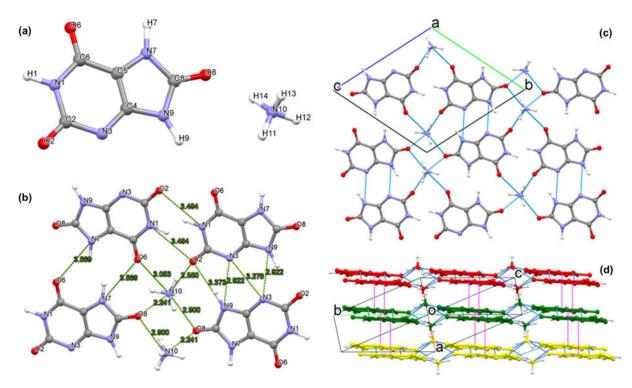


Fig. 3. The structure of marchettiite (colour code: C grey, N blue, O red and H white). (a) hydrogen urate anion and ammonium cation with atom labelling scheme. (b) The approximative distances between hydrogen bond donors and acceptors in the structure are indicated in green. (c) The hydrogen bonds (in blue) form a 2D network. (d) The hydrogen-bonded planes (distinguished with different colours) stack in the third dimension via $\pi \cdots \pi$ interactions (in magenta) among the anions. The structure is plotted using *Vesta 3* software (Momma and Izumi, 2011).

species approved by the IMA–CNMNC include: acetamide, CH₃CONH₂ (Srebrodol'skiy, 1975); guanine, C₅H₃(NH₂)N₄O (Bridge, 1974); kladnoite, C₆H₄(CO)₂NH (Jehlička *et al.*, 2007); tinnunculite, C₅H₄N₄O₃·2H₂O (Pekov *et al.*, 2017); urea, CO (NH₂)₂ (Bridge, 1973) and uricite, C₅H₄N₄O₃ (Bridge, 1974).

X-ray diffraction and marchettiite crystal structure

As noted above, the mineral marchettiite is the natural equivalent of the compound ammonium hydrogen urate, also known as the ammonium salt of uric acid (Sekkoum *et al.*, 2016).

The space group of marchettiite, in analogy with its synthetic counterpart, is $P\bar{1}$ and the unit-cell parameters refined by PXRD are: a=3.6533(2) Å, b=10.2046(7) Å, c=10.5837(7) Å, $\alpha=113.809(5)^\circ$, $\beta=91.313(8)^\circ$, $\gamma=92.44(1)^\circ$, V=360.312 Å and Z=2. The list of d spacings obtained by PXRD for marchettiite from the Le Bail refinement are reported in Table 3. The synthetic counterpart of marchettiite by Friedel et al. (2006) has the unit-cell parameters: a=3.650 Å, b=10.215 Å, c=10.597 Å, $\alpha=113.9^\circ$, $\beta=91.1^\circ$, $\gamma=92.3^\circ$ and V=360.8 Å The natural and synthetic compounds show an excellent agreement in all single-cell parameters and volume and this definitively supports the assumption that they have a nearly identical composition.

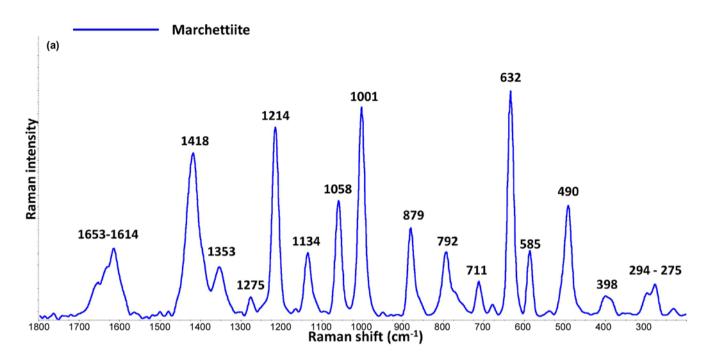
The crystal structure is composed by the flat and rigid hydrogen urate anions and ammonium cations (Fig. 4a), and features two ionic pairs per unit cell, related by the crystallographic inversion centre. Though the limited precision of the model does not allow a detailed discussion of the interactions, we can infer that hydrogen urate anions can interact side-by-side via pairs of

hydrogen bonds on each side, thus forming H-bonded ribbons. Ammonium cations occupy interstitial sites between neighbouring ribbons and are within favourable distances (\leq 3.0 Å) from the oxygen acceptor atoms (O2, O6 and O8) of four different anions (Fig. 4b). Due to ammonium's tetrahedral geometry, it cannot interact with all four acceptors simultaneously and it is probably disordered around its average position (we could not experimentally locate the positions of hydrogen atoms). All in all, the possible hydrogen bonds give rise to a waved 2D network (Fig. 4c). The crystal packing is completed by a parallel stacking of hydrogen urate moieties, into columns extending along the crystallographic a axis (Fig. 4d). The average stacking distance of 3.340(2) Å is indicative of a π ··· π interaction among the anions.

Because of the abundance of H-bond donors and acceptors available (3 donors and 4 acceptors on acid urate; 4 donors on ammonium) and the subtle asymmetry of the acid urate anion, it is possible that disorder and stacking faults occur in the structure. This could explain the 'low quality' and low resolution, observed in the PXRD patterns, not only for our natural marchettiite sample, but also for the synthetic ammonium hydrogen urate (Friedel *et al.*, 2006).

FTIR and Raman spectroscopy

A satisfying match between marchettiite and the synthetic analogue is given by the FTIR data (Benramdane *et al.*, 2008 and Sekkoum *et al.*, 2016) (Table 6). However, by comparison with the literature data, it appears evident that marchettiite, at least for the sample analysed by FTIR, is not a pure material but is intermixed with whewellite. Indeed, the bands at \sim 3175, 2891,



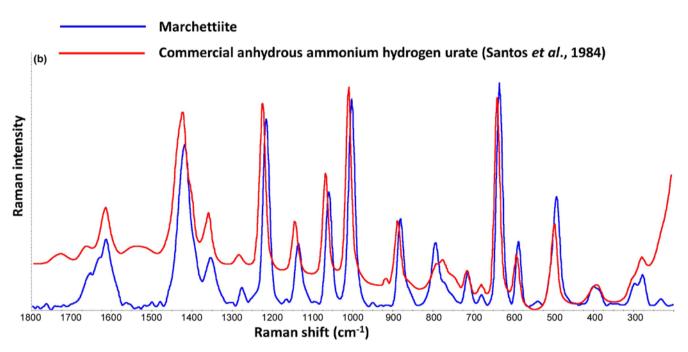


Fig. 4. Raman spectrum of marchettiite (a) and comparison with the synthetic analogue by Santos et al. (1984) in (b).

1540 and 768 cm⁻¹ are present for our marchettiite spectrum and the spectrum by Sekkoum *et al.* (2016) (see their table 3), which also contains whewellite.

The infrared characteristic broad bands of marchettiite in the range 3200–2600 cm⁻¹ correspond to N–H stretching of heterocyclic and ammonium groups. Other strong and sharp bands are: carbonyl (C=O) of urea groups at 1634–1640 cm⁻¹; conjugated carbonyl band at 1590–1596 cm⁻¹; valence stretching of =C–O at 1541 cm⁻¹; three bands around 1428–1432, 1388–1383 and 1341–1345 cm⁻¹ corresponding to N–H deformation; stretching vibration of O=C–H at 1274 cm⁻¹; other vibrations between

1139 and 1000 cm⁻¹ are attributed to the C-N stretching mode (Sekkoum *et al.*, 2016).

In terms of Raman spectroscopy, we can state that the marchettiite spectrum and that of its synthetic analogue (Santos *et al.*, 1984) are nearly identical as shown in Fig. 5b. On the basis of Daudon *et al.* (1983), the peak at 1418 cm^{-1} corresponds to C = N stretching, the peak at 1215 and 1058 cm^{-1} are related to C-O stretching and O-H bending, whereas the peak at 1001 cm^{-1} is probably due to the ring vibration. The two peaks at 880 and 792 cm^{-1} correspond to the N-H out of plane bending and ring vibration, respectively, whereas the peaks at 633 and

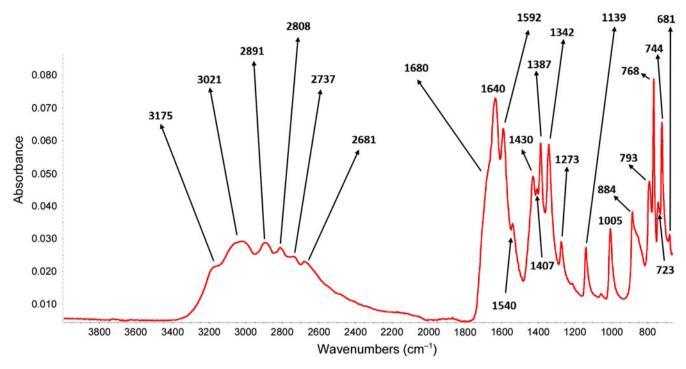


Fig. 5. Infrared spectrum of marchettiite.

Table 6. FTIR data for marchettiite in comparison with urate hydrogen ammonium from two sources. Bold notes the bands common to all three samples.

Manala attiita	Urate hydrogen ammonium +	Urate hydrogen
Marchettiite	whewellite	ammonium
This work	Sekkoum et al. (2016)	Benramdane et al. (2008)
	3431	
3175	3186	
3021	3071	3012
2891	2901	
2808	2803	2828
2737	2738	
2681	2678	2675
1680	1695	1673
1640	1640	
1592	1596	1589
1540	1541	
1430	1432	1435
1387	1383	1401
1342	1345	1349
1273	1274	1304
1139	1137	1123
1005	1001	1000
884	881	878
793	799	782
768	766	
744		740
723	722	
681		

585 cm⁻¹ are due to the ring breathing and ring vibration, respectively; additionally, the peak at 490 cm⁻¹ is a ring vibration.

Environment of formation of marchettiite

Considering that the analogue of marchettiite is known to be present in urinary stones, it is reasonable to think that the circulation

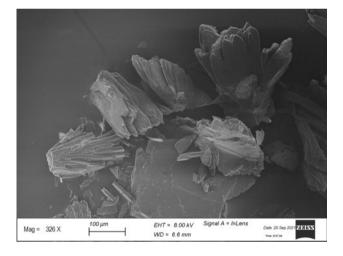


Fig. 6. Back-scatter electron image of marchettiite (\sim 850 μm wide). Holotype specimen, catalogue number MMP M171892.

of meteoric waters at temperatures between 20–30°C may have contributed to the formation of this new mineral at Mount Cervandone. Such an assertion is backed up by the fact that seven organic minerals have now been found at Mount Cervandone, and all of them were collected in a cleft associated with classic Alpine fissure minerals such as quartz, albite and muscovite. Deveroite-(Ce) is among these organic compounds: it is the natural hydrated REE-bearing analogue of the La, Ce and Nd synthetic oxalate compound (Ollendorf and Weigel, 1969; Gammons and Wood, 2000) probably formed by circulation of meteoric waters at low temperatures, between 20 and 30°C, that may occur on the rock surfaces during summer at 2500 m of altitude in the Italian Alps. Tinnunculite is another anhydrous uric acid (Pekov *et al.*, 2017) also described recently

at Mount Cervandone (Guastoni *et al.*, 2020); it is undoubtedly a natural organic mineral formed under the action of atmospheric agents in a geological environment with bird excrement mineralised on the surface of cliffs and rock fragments.

Nevertheless, the IMA-CNMNC considers that if a mineral derives from animal or organic waste it may be approved as a new mineral if the processes responsible for the crystallisation of the new compound excludes any human influence (Nickel and Grice, 1998). Marchettiite was formed from the transformation of animal excrement of birds or marmots at temperatures between 20 and 30°C – temperatures that may occur from the circulation of surface waters on the rock surfaces during the summer season on Mount Cervandone.

Further studies will clarify whether more organic samples are present in this area and what the potential is for such a deposit to exist. Natural organic phases have been previously overlooked because Mount Cervandone is renowned internationally as a mineral deposit for inorganic mineral phases formed during Alpine geological processes.

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