

RADIOCARBON CHRONOLOGY OF CENTRAL ALASKA: TECHNOLOGICAL CONTINUITY AND ECONOMIC CHANGE

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ABSTRACT. This research presents the first comprehensive radiocarbon chronology for central Alaska, encompassing the late Pleistocene and Holocene archaeological record. Dated component distributions, comprised of 274 ^{14}C dates from 160 components, indicate changing land-use strategies and subsistence economies, reflecting primarily lowland exploitation of bison, wapiti, and birds prior to 6000 cal BP, followed by increasing caribou and fish exploitation and use of upland areas. Microblade technology is conserved from the earliest components to ~1000 cal BP, and this continuity is not reflected in current cultural history sequences. Using component abundance as a proxy for population, initial colonization is associated with climate amelioration after ~14,000 cal BP, and population declines are associated with the Younger Dryas (13,000–12,000 cal BP) and initial establishment of widespread spruce forests (10,000–9000 cal BP).

INTRODUCTION

The archaeology of central Alaska, defined here as the Tanana, Susitna, and Copper River basins, encompasses an important record (Figure 1). This region is arguably the longest continuously inhabited area in the Western Hemisphere (Holmes et al. 1996; this paper). Archaeological data has the potential to contribute significantly to ongoing debates about the colonization of the New World and late Pleistocene extinctions (Hofman and Todd 2001; Grayson and Meltzer 2002; Waguespack and Surovell 2003; Shapiro et al. 2004). Ethnoarchaeological work on subsistence, settlement, and land-use strategies of hunter-gatherers in high-latitude environments have proven important in broader anthropological theory building (Amsden 1977; Binford 1977, 1978, 1980, 1991; Enloe 1993). However, synthetic work to date has typically been restricted to a few well-known sites (Sheppard et al. 1991), restricted to a limited time frame (Hamilton and Goebel 1999; Mason et al. 2001; Bever 2006) or a limited geographic area (Dixon et al. 1985).

Most current interpretations of prehistory are derived from cultural historical frameworks, which are more descriptive than explanative. These cultural sequences are based primarily on presence/absence of specific lithic tool types and technologies, rather than on differences in subsistence and land-use strategies, site structure, and organization (e.g. Cook and McKennan 1970; Cook 1975; Bacon 1977; Dixon 1985; Powers and Hoffecker 1989; West 1996b). There are limitations to these conceptual approaches as applied to assemblages in this region. These archaeological constructs are descriptive and employ normative concepts of culture, offering relatively few avenues for testing hypotheses for cultural change or adaptation, and can mask patterning in assemblage variability (Binford 1983). Cultural historical interpretations typically rely on relatively few excavated sites, increasing the potential effects of palimpsests on identifying discrete depositional or activity sets (Schiffer 1976; Carr 1985). Dry Creek Component 2 (C2) is a clear example: it is used as an exemplar of the microblade-bearing Denali complex (Powers et al. 1983; Dixon 1985; Hamilton and Goebel 1999), yet only 36% of the spatial clusters contain microblade technology (Potter 2005; Bever 2006). Many of the culturally “diagnostic” artifact types/classes are not restricted in time. The data presented here demonstrates long-term technological continuity that requires a re-evaluation of current cultural constructs and alternate approaches to explaining interassemblage variability.

In the last 30 yr, numerous cultural resource management and academic investigations have resulted in a great increase in empirical data, particularly radiocarbon-dated components, which have yet to be fully evaluated. Most of these data have never been synthesized on a regional basis, nor used as proxies to evaluate population trends. These dated components represent a useful data set for

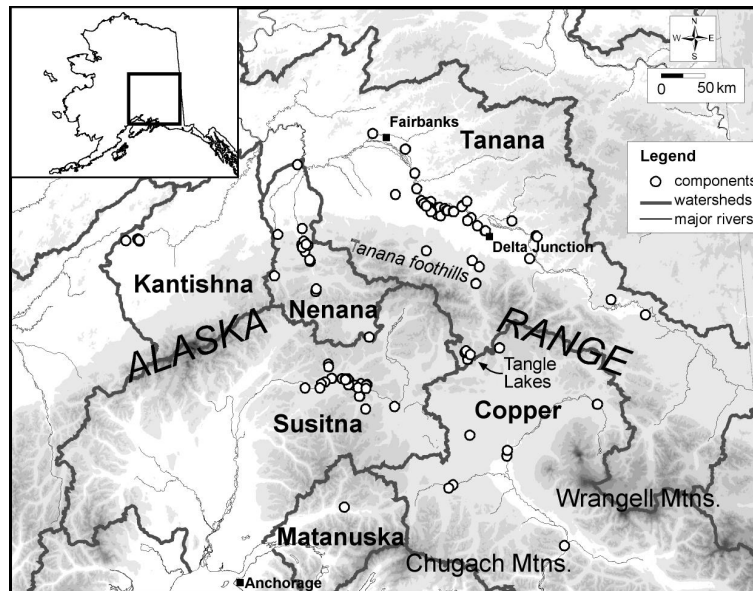


Figure 1 Central Alaska study area, showing dated component locations (elevation shaded in 500-m asl increments). The 2006–2007 surveys described in the text occurred between Fairbanks and Delta Junction south of the Tanana River.

estimating baseline data on technology and economy, as a first step for analyzing interassemblage variability.

This article addresses the limitations in previous intersite analyses by synthesizing a comprehensive record of ^{14}C -dated components and identifying major patterns of technological and economic change. Ambiguities and avenues for further inquiry are evaluated, and implications for cultural history are described. Detailed analysis patterning among lithic technology, faunal remains, residential and logistical mobility, and land use in the Holocene are presented in Potter (2008). The cultural transformation at ~1000 cal BP is explored through competing models of cultural change, population replacement, and taphonomic bias in Potter (forthcoming).

Technology and Chronology in Central Alaska

Cultural constructs in this region have been typically created on the basis of co-occurring sets of features at the level of attribute, type, and class. Thus, presence or absence of a particular type of projectile point (e.g. Chindadn, Kavik, side-notched) or class of lithics (e.g. flake burin), or even technology (e.g. microcore and blade) has been used to demarcate cultural entities, in the form of traditions, phases, or complexes, based on the proclivities of the originator. Cultural sequences vary somewhat, but most follow a basic pattern: an early period (>13,000 cal BP) marked by considerable technological variability (e.g. Chindadn complex, Nenana complex, East Beringian tradition, Northern Paleoindian tradition); followed by an early Holocene period dominated by microblade technology (e.g. Denali complex); followed by a mid-Holocene period associated with new technology (notched bifaces, notched cobbles, tabular cores) and possible continuation of microblade technology (e.g. Northern Archaic tradition, Tuktuk phase, Denali phase); followed by the late Holocene Athabascan tradition, associated with increased organic technology and presence of housepits and cache pits (e.g. Cook and McKennan 1970; Dixon 1985; Holmes 2001).

Ambiguities have been identified that affect these basic constructs and sequences, such as potential late occurrence of microblade and wedge-core technology (Cook 1969; Shinkwin 1979; Dixon 1985; Bowers 1999) and potential co-occurrence with notched bifaces, hallmarks of the Northern Archaic tradition (Bacon 1977; Cook and Gillispie 1986; Holmes 1986; Clark 2001; Ackerman 2004). The transition from early to middle Holocene complexes is also not well understood, and an occupation hiatus has been posited between 7700–6200 cal BP (or longer) (West 1996b:552; Potter 2004a). The ¹⁴C chronology developed here is used to evaluate these ambiguities.

Previous Radiocarbon Syntheses

This article represents the first comprehensive archaeological ¹⁴C chronology for this region. Only a few well-known sites dominate the cultural historical literature for the region (Hamilton and Goebel 1999; Holmes 2001). Previous intersite comparisons focus primarily on late Pleistocene/early Holocene archaeology with little attention to the middle and late Holocene period. The most influential compendium of ¹⁴C data were numerous papers in West (1996a), but these were limited to components older than ~7000 BP. These data were used widely in late Pleistocene/early Holocene reviews, focused primarily on the peopling of the Americas (West 1996b; Dilley 1998; Hamilton and Goebel 1999; Dixon 1999, 2001; Yesner 2001; Bever 2006). Mason et al. (2001) expanded parts of this sample, focusing on Denali complex sites in Alaska and Yukon Territory, but again only considered components older than ~7000 BP. A comprehensive ¹⁴C database was compiled and calibrated for the Copper River basin and surrounding highlands (Potter 1997), encompassing only a part of this study area.

METHODS

Database Development, Variables, and Data Limitations

This ¹⁴C database was compiled from published articles, books, cultural resource reports, and theses, and does not include ongoing investigations where the results have not yet been published. The database includes 274 ¹⁴C dates from 160 components at 113 sites (see Appendix). Component delineation followed the original investigators, with exceptions noted below. ¹⁴C dates on cultural features were preferred over stratigraphic dates (e.g. Broken Mammoth Cultural Zone [CZ] 4). Bone apatite dates were not considered due to susceptibility of contamination, and soil organic (bulk sampled) dates were only considered if charcoal dates were unavailable. Dates determined to be discordant by the original investigator were not used (with a few exceptions, see below). Multiple dates on single stratigraphic contexts were averaged following Ward and Wilson (1978) using the CALIB v 5.0 program (Stuiver and Reimer 1993), providing a single age estimate. The age estimates were calibrated using CALIB v 5.0 with the IntCal04 terrestrial calibration curve (Reimer et al. 2004). Components were grouped into 1000-calendar yr intervals by the median of each date range to mitigate the lack of precision of single age estimators.

Variables gathered from the primary literature included lithic assemblage characteristics and associated fauna. To counter sample size effects, microblade technology, notched bifaces, and fauna, taxa were denoted as presence/absence. Space-averaging may be affected by environmental differences among subregions (Lyman 2003). Environmental variability in the study area is primarily affected by elevation, with the Tanana-Kuskokwim and Copper River lowlands currently dominated by boreal forests, contrasted with the foothills of the Alaska Range and Talkeetna Mountains dominated by moist and alpine tundra and dwarf and tall shrubs (Warhaftig 1965; Gallant et al. 1995). This dichotomy is evaluated by assigning values of “upland” and “lowland” to components in these 2 environments (the break is around 500 m asl). Five major subregions are distinguished within the study area; 2 primarily lowland areas currently dominated by boreal forest (Tanana and Copper

River basins, areas below 500 and 1000 m asl, respectively), and 3 primarily upland areas dominated by moist and alpine tundra (Upper Nenana and Upper Susitna River valleys and the Tanana foothills/Tangle Lakes area) (Figure 1). Sites within the lower Nenana and Kantishna basins are included in the Tanana basin subregion. Sites within the Matanuska basin are included in the Upper Susitna subregion. Absolute elevation (in m asl) was derived from the 15-min digital elevation models (DEM) for Alaska (US Geological Survey 1979).

Components were grouped by time periods derived from transitions among cultural constructs within cultural historical sequences in order to assess broad levels of economic and technological change. Late Pleistocene (14,000–12,000 cal BP) comprises early complexes like Chindadn and Nenana, and is associated with glacial conditions ($n = 11$ components). Early Holocene (12,000–6000 cal BP) comprises the Denali complex, and is associated with the expansion of the boreal forest ($n = 51$ components). Middle Holocene (6000–1000 cal BP) comprises the Northern Archaic tradition and Late Denali complex ($n = 76$ components). Late Holocene (<1000 cal BP) comprises the Athabascan tradition ($n = 22$ components).

There are several limitations to these data, including cultural contexts, sampling, and taphonomic bias. Many components have associated stratigraphic dates or single dates on cultural features, both of which may reduce dating accuracy and precision. Over half of these components are dated through associated stratigraphic dates ($n = 99$, 62% of the total), 54 (34%) have dates associated with cultural features, and 7 (4%) have unknown/unreported associations. Do these data constitute a representative sample of cultural components within the region? The question is difficult to answer, given the current level of understanding and the relative lack of integrative intersite variability studies (Potter 2005). Site discovery is directly related to sampling effort, which has largely followed development in the region. However, several linear transects cross the study area, oriented both east-west and north-south, providing checks against this bias (e.g. Cook 1977; Aigner and Gannon 1981a,b; Bowers et al. 1995; Potter et al. 2002, 2007a,b). Surveys have resulted in discovery of components from every period of human occupation in North America, and while investigator bias for earlier sites (more common as research topics) may factor here, a recent survey through the mid-Tanana basin resulted in the discovery of 56 buried prehistoric components, 36 of them dated without bias for expected age (Potter et al. 2007b). This distribution generally matches the distribution of previously dated components, with some exceptions noted below.

Taphonomic bias favoring later components is difficult to evaluate without detailed geoarchaeological investigations. However, the exponential population curve peaking in the most recent interval predicted by Surovell and Brantingham (2007) in cases of taphonomic bias is not observed here (compare their Figure 2 with Figure 2, next page). For this reason, the dated components are used as proxies for paleodemography. However, due to the limitations discussed above, the patterns presented here should be seen as tentative pending further research on search image adequacy, stratified sampling, and regional geoarchaeology.

Problematic Sites

A number of sites presented problems in component delineation and age estimation, and for the sake of clarity and completeness, they are detailed here. Teklanika West possibly contains 3 components (Goebel et al. 1996), but the cultural material has not been demarcated on a stratigraphic basis (West 1996c), so I follow Mason et al. (2001) in only listing the earliest component (C1). The later Holocene dates from Donnelly Ridge (West 1967) and the Little Panguingue Creek hearth (Hofecker and Powers 1996) are tentatively accepted here given their acceptance by other archaeologists (e.g. Shinkwin 1979:161–2), and lack of evidence for contamination.

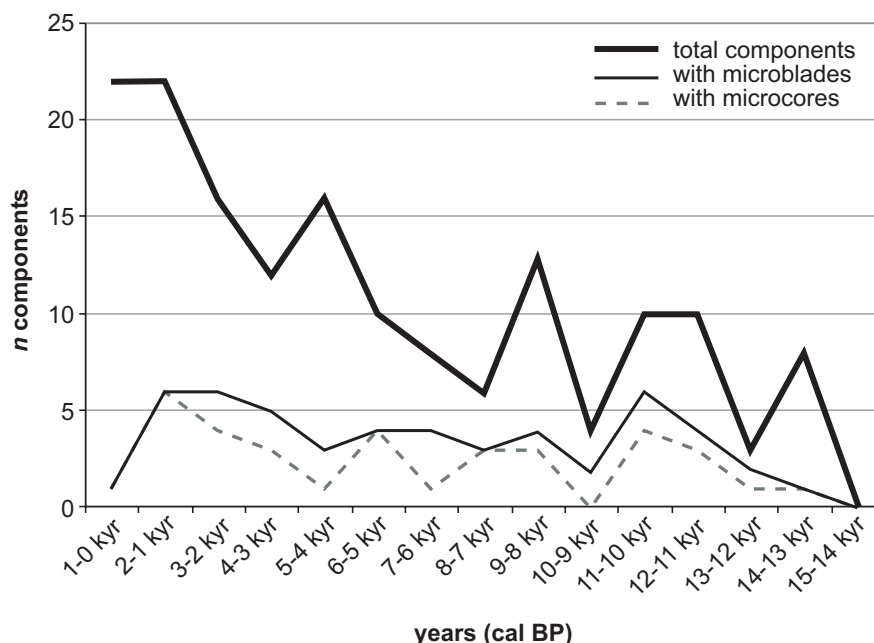


Figure 2 Distribution of component ages and associated technology per 1000-yr age interval (using median calibrated values).

Dry Creek C2 may be composed of multiple occupations (Mason et al. 2001), or Dry Creek C1 and C2 may be a single component with post-depositional disturbance (Thorson 2006). Here, Dry Creek C1 and C2 are considered 2 distinct components, following Powers et al. (1983). Dixon (1993:86) lists new ¹⁴C dates for Jay Creek Ridge (~9800 BP), but there is no report on context, so the dates in Dixon et al. (1985) are used here (~7000 BP), following Mason et al. (2001). The chronologically older dates suggest the Upper Susitna area was habitable by at least 12,000 cal BP. West et al. (1996b:388) note that Whitmore Ridge Component 2 occurs in the A2b horizon, with 3 stratigraphically associated dates, 2 of which overlap. Though West et al. (1996b:393) suggest the dates are too young, they are consistent with the overall stratigraphy and ¹⁴C chronology, and since no evidence for contamination is provided, the 2 overlapping dates are tentatively accepted here.

The ¹⁴C record at Healy Lake Village site has been extensively discussed (Cook 1969, 1996; Erlandson et al. 1991; Dilley 1998). For this synthesis, only dates derived from charcoal are used for Levels 1–5. Cook (1996) defined 3 cultural stages at Healy Lake Village: Athabascan (Levels 1–3); Transitional (Levels 4–5); and Chindadn (Levels 6–10). These distinctions are kept for this synthesis, except that Level 1 is separated from Levels 2–3, given a significantly younger hearth date in Level 1 and overlap of dates in Levels 2–3. Chindadn samples date between 11,400–8000 BP, with no correlation between depth and age. Though multiple occupations are likely, it is interpreted to be a single component dating to the average of charcoal dates, following Cook (1996).

RESULTS

Radiocarbon Chronology

Assuming component abundance reflects population size, the ¹⁴C-dated component distribution was used to estimate paleodemography. Figure 2 illustrates the absolute number of dated components per

1000 cal yr BP. The overall trend is a gradual increase in population to a peak of 2000–1000 cal BP. The decrease in the final 1000-yr period reflects the reliance on typological dating during the proto-historic period (e.g. trade beads). At this low resolution, there are peaks and dips in component abundance. This distribution may be affected by sampling bias, particularly disproportional focus on sites of a certain age or stratigraphic setting. For example, the North Alaska Range project that resulted in the discovery and testing of numerous sites in the Upper Nenana Valley was designed to locate intact late Pleistocene landforms (Powers et al. 1983; Hoffecker 1985).

Recent linear surveys in the mid-Tanana basin, yielding 36 components dated without bias for expected age (Potter et al. 2007a,b), offer data to evaluate this possibility. Figure 3 shows component percentages for each interval based on all components from the 2006–2007 surveys ($n = 36$ components) compared with all previously known data ($n = 124$ components compiled in Potter 2004a). The distributions are relatively similar except for relatively more sites in the early Holocene period (9000–5000 cal BP), and fewer sites in the mid to late Holocene (after 5000 cal BP). The early Holocene period is often interpreted as a time of transition, from earlier Beringian technology and subsistence to boreal forest adaptations associated with the Northern Archaic tradition (Anderson and Douglas 1968; Dixon 1985; Clark 1994). Analysis of previous intersite data (Potter 2004a) indicated a possible hiatus in occupation in central Alaska between 7700–6200 cal BP (only 5 components were previously known from this period, ~4% of the total). However, the new data set includes 6 components dating to this period (15% of the new components), demonstrating the presence of human occupation of the region throughout the mid-late Holocene (Figure 2). Technological data indicates many aspects of technology (e.g. microblades, wedge-shaped microcores) were conserved through this period, thus linking the early and later Holocene microblade industries.

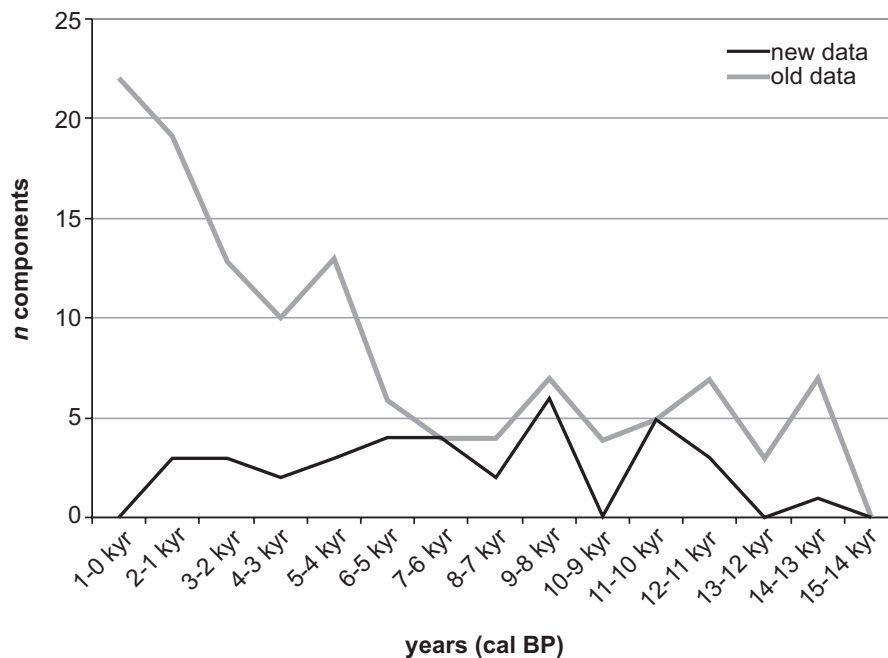


Figure 3 Comparison of component age distributions from recent surveys in the mid-Tanana basin (Potter et al. 2007a,b), labeled “new data,” and from previous surveys (Potter 2004a), labeled “old data.”

The general similarity of these distributions suggests that these data are representative enough for exploration of this distribution as a proxy for paleo-population. After an initial peak at 14,000–13,000 cal BP, representing the initial occupation of Alaska, the population dips for the next thousand-yr interval, rising again between 12,000–11,000 cal BP. The initial colonization event(s) have been discussed from many different perspectives (e.g. West 1981, 1996b; Powers and Hoffecker 1989; Hamilton and Goebel 1999; Dixon 2001; Holmes 2001; Yesner 2001; Bever 2006), and these dates support a correlation of early colonization with climate amelioration after ~14,000 cal BP, associated with the transition from herb tundra to shrub tundra (Ager and Brubaker 1985; Bigelow and Powers 2001).

The depopulation tracked at 13,000–12,000 cal BP correlates broadly to the Younger Dryas stadial, a cold period that is associated with glacial readvance (Bigelow and Edwards 2001). These data run counter to Bigelow and Powers (2001) and Mason et al. (2001), who noted no decrease in site occurrence during this period, but the latter examined only those sites associated with the Denali complex. This depopulation trend is, however, noted by Bever (2006:612), and importantly this pattern is reinforced here with a more widespread Alaska-wide context, indicating that the effects of the Younger Dryas may have been significant. Bever (2006:613) further notes that the diversity of the pre-Younger Dryas material contrasts with the single Denali complex technology after the Younger Dryas. The continuity of microcore and blade technology evidenced here could be explained by 1) continuity of regional populations using the same technology or 2) population replacement by microblade-using groups from outside the region (Siberia or northwest North America, see technology discussion below).

A second sharp population decrease is inferred for the 10,000–9000 cal BP interval, which might be correlated with the establishment of widespread spruce forests (*Picea* spp.) in the Tanana basin in the early Holocene. Kaufman et al. (2004:536) note the Holocene Thermal Maximum between 11,300–9100 cal BP in this region. West (1981:221–4) argues for an early Holocene peak and subsequent population crash as warming climate and growth of the boreal forest decimated Denali complex populations. The peak at 9000–8000 cal BP is harder to correlate with broader climate or vegetation changes. Mason et al. (2001) examined a selection of early Holocene Denali complex components in Alaska and the Yukon Territory, and their date distribution is relatively similar to the one presented here between 10,000–8000 cal BP. They interpret the spike at 9000–8000 cal BP as increased occupation associated with a cooling event at ~8200 cal BP (Klitgaard-Kristensen et al. 1998) due to increased abundance of caribou (Mason et al. 2001:539). However, faunal analysis using a much larger data set indicates that caribou hunting becomes more dominant after ~5000 cal BP (see below, Potter 2008).

The mid-late Holocene (after 6000 cal BP) is generally characterized with increasing population, especially after 3000 cal BP. The greater relative abundance of sites could also be due to increased archaeological visibility or a biased search image. While no known significant climate or vegetation change is known for this period, new technology and artifact types enter the region (including side-notched biface forms, notched cobbles, and tabular microblade cores) (Cook and McKennan 1970; Dixon 1985). However, older technologies were also conserved and were used alongside the new forms (i.e. wedge-shaped microblade core forms). While the archaeological data could support partial population replacement (Dumond 1969; Workman 1978) or diffusion (Clark 1994), the component distribution may reflect an effective adaptation to the boreal forest (probably through a combination of new technology and new settlement and subsistence strategies).

Regional Chronologies

Different land-use strategies are apparent in the component distributions. Figure 4 illustrates component ages for the 3 upland and 2 lowland areas in the study area. The Tanana basin subregion was largely unglaciated during the late Pleistocene (Kaufman and Manley 2004), and the northern foothills of the Alaska Range (Upper Nenana and Tanana foothills/Tangle Lakes subregions) were first occupied during the late Pleistocene, between 13,000–12,000 cal BP. The Upper Susitna was first occupied by ~8000 cal BP (or ~11,400 cal BP if the early dates in Dixon [1993:86] are considered). The Copper basin was dominated by the glacier-dammed Lake Atna until ~11,600–9700 cal BP (Ferrians 1989:87), but the earliest known components date to ~2500 cal BP.

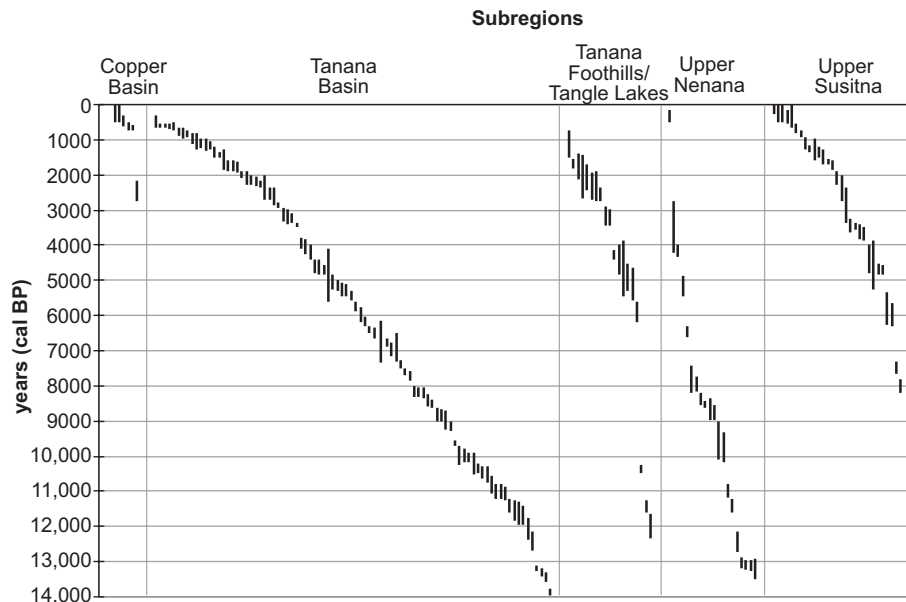


Figure 4 Component calibrated age ranges (2σ), ordered by median age estimator and subregion

The upland areas are differentially occupied through time, comprising 32% of early Holocene components, 53% of middle Holocene components, and 36% of late Holocene components (the late Pleistocene is not considered since some of the upland areas were not yet ice-free). Elevation data for each component confirms this pattern (Figure 5). Early Holocene components average 424 ± 187 m asl (median = 361) compared with 522 ± 243 m asl (median = 487) for middle Holocene and 462 ± 206 m asl (median = 487) for late Holocene components. Average elevation values are significantly different between early and middle Holocene components (t test for independent samples, $t = -2.42$, $df = 25$, $p = 0.017$), whereas middle and late Holocene components have similar distributions of sites relative to elevation ($t = 1.05$, $df = 96$, $p = 0.294$). Average elevation values are similar between the late Pleistocene and early Holocene occupations ($t = -0.89$, $df = 60$, $p = 0.377$), indicating similar land-use strategies in the early Holocene associated with lowland areas, even after upland areas were deglaciated.

While the Tanana subregion has a continuous record with no major breaks, there are several breaks in the upland subregions (Figure 4). These upland areas all have chronological gaps between 7700–6200 cal BP, and the fact that this is replicated in all 3 upland regions may indicate widespread

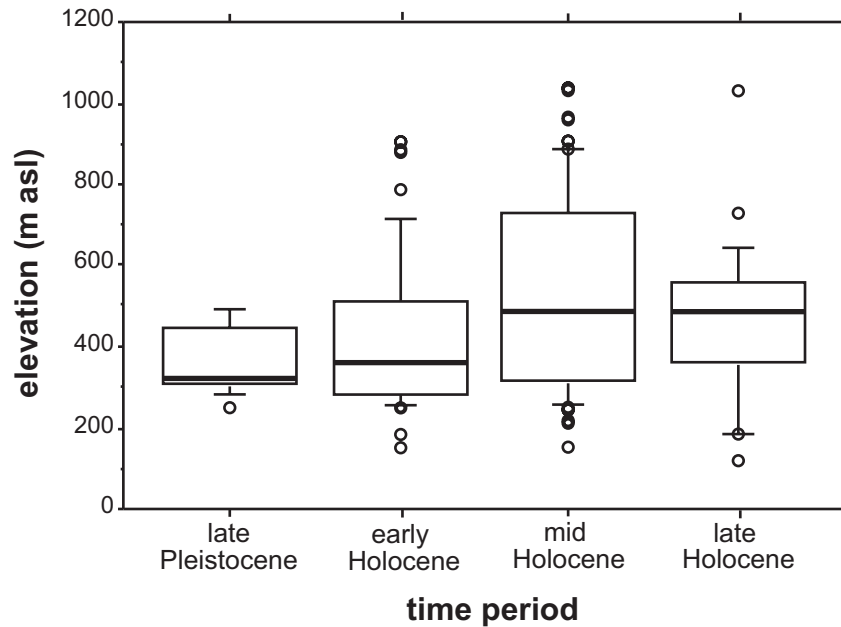


Figure 5 Box plots of elevation by time period

changing land-use strategies in the early to middle Holocene, rather than sampling bias. Consequently, the increased use of upland areas seen after ~6200 cal BP, along with new technologies associated with the Northern Archaic tradition, likely reflects new land-use strategies linked with seasonal caribou exploitation, and is consistent with increased numbers of components with caribou remains during this period (see below, Potter 2008). Shorter chronological gaps between 4000–3000 cal BP in the Tangle Lakes area may be artifacts of small sample size or changes in population or land-use strategies.

DISCUSSION

Technological Continuity

These ¹⁴C-dated component age estimates provide baseline data useful for a variety of purposes, including assessing technological and subsistence change and estimating paleo-population change. Continuity of microblade technology is readily apparent, as demonstrated in Figure 2. This not only includes microblades as products of this technology, but also particular core forms, including wedge-shaped microcores with associated core tablets rejuvenation (Holmes 2001, 2004; Yesner and Pearson 2002; see also Bowers 1999). Microblade technology generally parallels overall component abundance throughout the Holocene until ~1000 cal BP.

In the earliest period (14,000–13,000 cal BP), there is 1 microblade-bearing component (Swan Point CZ4) and 7 non-microblade-bearing components. However, only 5 of the latter have more than 14 m² excavated—and of these, only Walker Road C1 contains more than 39 retouched pieces. Powers and Hoffecker (1989) posited a non-microblade Nenana complex on the basis of some of these sites, but since then, Chindadn points (diagnostic to that complex) have been found associated with microblades at Swan Point CZ3 (Holmes 2008) and Broken Mammoth CZ3 (Krasinski 2005:32). This supports the contention that Chindadn points and microblades associated at Healy Lake (Cook

1996) are not from mixed contexts as has been suggested (Hoffecker et al. 1993). One may argue that the relatively low frequencies of microblades in Swan Point CZ3 ($n = 37$) and Broken Mammoth CZ3 ($n = 44$) (Krasinski 2005; Holmes 2008) suggest stratigraphic mixing; however, at Gerstle River C3, several contemporaneous and spatially discrete lithic clusters contain microblades at varying frequencies (between 1 and 242, comprising 2% to 29% of total debitage; Potter 2005). This variation could be due to many factors, including activity area differences or the tendency for microblades to be deposited in small discrete loci, easily missed depending on sampling strategies. The point here is that microblades are present in sites assigned to the both Nenana and Denali complexes. This is consistent with the hypothesis that Nenana and Denali complexes represent different portions of a single late Pleistocene technological tradition (West 1996b; Holmes 2001).

After 13,000 cal BP, these data suggest that microblades represent a conservative technology that was well suited to coping with climatic oscillations at the Pleistocene/Holocene transition as well as the expanding boreal forest. Increased upland exploitation in the mid-Holocene seems related to subsistence/settlement strategy changes given the opportunity for exploitation in earlier periods. This coincides with the introduction of cultural material like notched bifaces and notched pebbles, associated with the Northern Archaic tradition (Anderson and Douglas 1968; Dixon 1985). These data show that along with this new technology, early Holocene and middle Holocene populations used the landscape in different ways, partially reflected in increased upland use, but also in faunal assemblage differences (see next section).

Economic Change

Of the 160 components in this study, 87 contain faunal remains (54% of total). Of these 87 components, 26 have fragmented, burned, and calcined fragments that are not analyzed further. The remaining 62 components (39% of total) provide a record for subsistence economies. Table 1 summarizes the variability in faunal presence/absence among components for each time period. Only those taxa present at >5 components are included. Large and small mammal categories follow from the original investigators. The late Pleistocene period is somewhat skewed by the Broken Mammoth CZ4 assemblage, which contains almost all of the listed taxa (Yesner 1996), and the generally small sample size from the Late Pleistocene ($n = 5$) should be considered.

While these data are very coarse grained (Σ NIISP is not used), significant patterning is evident. Bison and wapiti occurrence within archaeological assemblages decrease through time, whereas caribou and to a lesser extent moose increases. The sharpest break is between the early and middle Holocene (6000 cal BP). Most small and medium mammals appear in relatively more assemblages in the middle and late Holocene, particularly hare and canids. Fish are also more common in the later Holocene, but interestingly, birds are more common in the Late Pleistocene. These patterns indicate changes in subsistence economies consistent with the land-use patterns noted above. Both of these data sets indicate a shift from a broad subsistence base using primarily lowland areas incorporating bison and wapiti in the late Pleistocene and early Holocene to acquisition of more seasonally abundant game (caribou, fish) in the middle and late Holocene. The growing importance of caribou is reflected in increasing use of upland areas like the Upper Susitna Valley, where faunal assemblages are dominated by caribou remains.

Cultural History

Given these component age distributions, a modification of current cultural historical sequences may be in order. While late Pleistocene material may reflect considerable diversity, the East Beringian tradition proposed by Holmes (2001) may adequately encompass this variability, given that

Table 1 Faunal patterning by time period; cells represent percentages of total components per time period with at least 1 specimen of each taxonomic classification.

| Taxa (number of components where they occur) | Late Holocene (0–1000 cal BP) <i>n</i> = 16 | Middle Holocene (1000–6000 cal BP) <i>n</i> = 26 | Early Holocene (6000–12,000 cal BP) <i>n</i> = 14 | Late Pleistocene (12,000–14,000 cal BP) <i>n</i> = 5 |
|--|--|---|--|---|
| General Size Classes | | | | |
| L mammal (<i>n</i> = 59) | 94% | 88% | 100% | 100% |
| S mammal (<i>n</i> = 28) | 56% | 50% | 29% | 40% |
| Ungulates | | | | |
| Caribou (<i>n</i> = 29) | 63% | 58% | 21% | 20% |
| Moose (<i>n</i> = 15) | 31% | 23% | 21% | 20% |
| Bison (<i>n</i> = 8) | 0% | 4% | 36% | 40% |
| Wapiti (<i>n</i> = 7) | 0% | 0% | 29% | 60% |
| Sheep (<i>n</i> = 7) | 13% | 0% | 21% | 40% |
| Other Mammals | | | | |
| Hare (<i>n</i> = 15) | 38% | 23% | 14% | 20% |
| Beaver (<i>n</i> = 11) | 13% | 27% | 7% | 20% |
| Canid (<i>n</i> = 9) | 31% | 8% | 7% | 20% |
| Bear (<i>n</i> = 6) | 13% | 12% | 7% | 0% |
| Birds and Fish | | | | |
| Birds (<i>n</i> = 13) | 19% | 15% | 21% | 60% |
| Fish (<i>n</i> = 13) | 31% | 23% | 7% | 20% |

material diagnostic to both Nenana and Denali complexes are found intermixed. Between ~12,000–6000 cal BP, the archaeology is dominated by cultural material assigned to the Denali complex (or Paleoarctic tradition). The Northern Archaic tradition, typically dated between 6000–3500 cal BP in this region (Dixon 1985), should be extended to ~1000 cal BP on the basis of continuity of lithic types and basic settlement patterns. A distinct Late Denali complex (Holmes 1977; Dixon 1985) is unnecessary, given the continuity of microblade technology. The well-known transformation in settlement, site structure, and technology associated with the Athabascan tradition (Shinkwin 1977, 1979; Clark 1981; Dixon 1985) occurred in this region between ~1300–800 cal BP, tentatively dated to ~1000 cal BP. This transition is explored in detail in Potter (forthcoming).

These data demonstrate continuity in certain technological elements along with economic and settlement system changes. Cultural historical constructs as currently developed on the basis of lithic typology alone may not be adequate to explain this cultural change. Rather, these cultural changes appear to relate more to variation in settlement, mobility, and subsistence systems. Potential avenues for exploring this diverse record involve analyzing site location, site structure, and organization, along with more traditional data sets like lithic typology and faunal remains. In this context, identification and description of recurring depositional and activity sets will be useful. Understanding how tools and toolkits were used as part of adaptive systems, incorporating settlement strategies and subsistence economies within a logistical and residential mobility system will result in more robust explanations of cultural change in this region.

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APPENDIX

Table A1 Radiocarbon-dated component date list (see Methods section for details).

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2-σ range) | Reference |
|------------------------|----------------------------|-------------|-----------------------|-----------------------------|--------------------------|---------------------------------|
| Swan Point CZ4 | 11,660 ± 70 | Beta-56667 | c | stratigraphic | 13,693–13,344 | Holmes et al. 1996; Holmes 2004 |
| | 11,660 ± 60 | Beta-71372 | c | stratigraphic | 13,680–13,360 | |
| | 11,770 ± 140 | AA-19322 | o | organic residue on artifact | 13,901–13,320 | |
| | 12,040 ± 40 | unreported | c | hearth | 14,004–13,788 | |
| | 12,060 ± 70 | CAMS-17045 | iv | ivory collagen | 14,069–13,768 | |
| | 12,110 ± 50 | unreported | o | hearth residue | 14,090–13,823 | |
| | 12,360 ± 60 | unreported | o | hearth residue | 14,748–14,067 | |
| average | 12,003 ± 22 | | | 13,956–13,775 | | |
| Mead CZ4 | 11,560 ± 80 | CAMS-5198 | c | stratigraphic | 13,614–13,253 | Dilley 1998 |
| | 11,600 ± 60 | CAMS-4877 | c | stratigraphic | 13,618–13,302 | |
| Broken Mammoth CZ4 | 11,587 ± 48 | | | | 13,581–13,301 | Holmes 1996 |
| | 11,420 ± 70 | CAMS-5358 | c | hearth | 13,411–13,152 | |
| average | 11,510 ± 120 | WSU-4262 | c | hearth | 13,647–13,154 | |
| | 11,443 ± 60 | | | | 13,415–13,195 | |
| Owl Ridge C1 | 11,340 ± 150 | Beta-11209 | c | stratigraphic | 13,483–12,919 | Phippen 1988 |
| Little Delta Dune C1 | 11,300 ± 40 | Beta-232394 | c | hearth –split | 13,262–13,105 | Potter et al. 2007b |
| | 11,250 ± 60 | AA-76863 | c | hearth –split | 13,263–13,020 | |
| | 11,420 ± 60 | Beta-233316 | c | hearth | 13,401–13,173 | |
| average | 11,320 ± 30 | | | | 13,269–13,124 | |
| Walker Road C1 | 11,010 ± 230 | AA-1683 | c | hearth | 13,377–12,397 | Goebel et al. 1996 |
| | 11,170 ± 180 | AA-1681 | c | hearth | 13,363–12,834 | |
| | 11,300 ± 120 | AA-2264 | c | hearth | 13,379–12,948 | |
| average | 11,220 ± 92 | | | | 13,263–12,938 | |
| Moose Creek C1 | 11,190 ± 60 | Beta-96627 | c | hearth | 13,209–12,952 | Pearson 1999 |
| Dry Creek C1 | 11,120 ± 85 | SI-2880 | c | stratigraphic | 13,183–12,893 | Powers et al. 1983 |
| Moose Creek C2 | 10,500 ± 60 | Beta-106040 | c | hearth | 12,714–12,160 | Pearson 1999 |
| Mead CZ3 | 10,410 ± 80 | CAMS-5197 | c | stratigraphic | 12,652–12,045 | Dilley 1998 |
| | 10,460 ± 110 | CAMS-4876 | c | stratigraphic | 12,769–12,045 | |
| | 10,760 ± 170 | WSU-4425 | c | stratigraphic | 13,069–12,175 | |
| average | 10,469 ± 60 | | | | 12,671–12,142 | |
| Broken Mammoth CZ3 | 10,290 ± 70 | CAMS-5357 | c | hearth | 12,386–11,769 | Holmes 1996 |
| Phipps site | 10,230 ± 70 | Beta-63672 | c | stratigraphic | 12,350–11,624 | West et al. 1996a |
| Swan Point CZ3 | 10,010 ± 90 | Beta-190578 | c | hearth | 11,956–11,242 | Holmes et al. 1996 |
| | 10,025 ± 60 | Beta-170458 | c | hearth | 11,805–11,268 | |
| | 10,230 ± 80 | Beta-56666 | c | hearth | 12,366–11,616 | |
| average | 10,079 ± 42 | | | | 11,957–11,396 | |
| XBD-308 | 10,050 ± 70 | Beta-219659 | c | stratigraphic | 11,958–11,286 | Potter et al. 2007a |
| XBD-338 C2 | 10,000 ± 80 | Beta-232397 | c | stratigraphic | 11,819–11,240 | Potter et al. 2007b |
| Whitmore Ridge C1 | 9600 ± 140 | Beta-64578 | so | stratigraphic | 11,249–10,525 | West et al. 1996b |
| | 9830 ± 60 | Beta-70240 | so | stratigraphic | 11,394–11,161 | |
| | 9890 ± 70 | Beta-62222 | so | stratigraphic | 11,609–11,194 | |
| | 10,270 ± 70 | Beta-77268 | so | stratigraphic | 12,377–11,761 | |
| average | 9953 ± 37 | | | | 11,603–11,249 | |
| Little Delta River #3 | 9920 ± 60 | Beta-12331 | c | stratigraphic | 11,610–11,216 | Higgs et al. 1999 |
| Panguingue Creek C1 | 10,180 ± 130 | AA-1686 | c | stratigraphic | 12,379–11,318 | Powers and Maxwell 1986 |
| | 9836 ± 62 | Gx-17457 | c | stratigraphic | 11,401–11,140 | |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2- σ range) | Reference |
|---|-------------------------------|-------------|-----------------------|----------------------|-------------------------------------|---|
| | 9850 \pm 140 | Beta-55101 | c | stratigraphic | 11,818–10,777 | |
| average | 9893 \pm 52 | | | | 11,599–11,202 | |
| Gerstle River C1 | 9740 \pm 50 | Beta-133751 | c | stratigraphic | 11,247–10,883 | Potter 2005 |
| Little Delta Dune C2 | 9670 \pm 40 | Beta-232393 | c | stratigraphic | 11,203–10,796 | Potter et al. 2007b |
| Dry Creek C2 | 7895 \pm 105 | SI-2328 | c | stratigraphic | 9009–8459 | Powers et al. 1983; Bigelow and Powers 1994 |
| | 8915 \pm 70 | AA-11730 | c | stratigraphic | 10,233–9776 | |
| | 9340 \pm 195 | SI-2329 | c | stratigraphic | 11,186–10,178 | |
| | 9340 \pm 95 | SI-11733 | c | stratigraphic | 11,056–10,248 | |
| | 9690 \pm 75 | AA-11732 | c | stratigraphic | 11,236–10,778 | |
| | 10,060 \pm 85 | AA-11727 | c | stratigraphic | 11,971–11,273 | |
| | 10,540 \pm 70 | AA-11731 | c | stratigraphic | 12,792–12,238 | |
| | 10,615 \pm 100 | AA-11728 | c | stratigraphic | 12,845–12,239 | |
| | 10,690 \pm 250 | SI-1561 | c | hearth | 13,116–11,822 | |
| average | 9657 \pm 31 | | | | 11,191–10,801 | |
| Little Delta Dune C3 | 9650 \pm 60 | Beta-220218 | c | hearth | 11,200–10,775 | Potter et al. 2007b |
| Gerstle River C2 | 9400 \pm 50 | Beta-183110 | c | hearth | 11,057–10,571 | Potter 2005 |
| | 9510 \pm 40 | Beta-134098 | c | hearth | 11,075–10,609 | |
| average | 9449 \pm 41 | | | | 11,057–10,571 | |
| XBD-303 | 9340 \pm 80 | Beta-219658 | c | stratigraphic | 10,740–10,281 | Potter et al. 2007a |
| XBD-312 | 9290 \pm 50 | Beta-220214 | c | stratigraphic | 10,647–10,285 | Potter et al. 2007a |
| Sparks Point | 9060 \pm 425 | UGa-941 | so | stratigraphic | 11,335–9033 | West et al. 1996c |
| | 9110 \pm 80 | Beta-64577 | so | stratigraphic | 10,515–9967 | |
| | 9200 \pm 60 | Beta-62773 | so? | stratigraphic | 10,514–10,237 | |
| average | 9166 \pm 47 | | | | 10,486–10,233 | |
| Healy Lake Village Chindadn (levels 6–10) | 8655 \pm 280 | Gx-2171 | so? | stratigraphic | 10,406–9010 | Cook 1996 |
| | 8680 \pm 240 | Gx-2170 | c | stratigraphic | 10,287–9092 | |
| | 8990 \pm 60 | Beta-76070 | c | stratigraphic | 10,245–9916 | |
| | 9245 \pm 213 | AU-1 | c | hearth | 11,162–9892 | |
| | 9895 \pm 210 | Gx-2174 | c | hearth | 12,111–10,703 | |
| | 10,040 \pm 210 | SI-739 | c | stratigraphic | 12,567–10,874 | |
| | 10,434 \pm 279 | AU-3 | c | stratigraphic | 12,869–11,342 | |
| | 10,500 \pm 280 | Gx-1944 | c | stratigraphic | 12,925–11,394 | |
| average | 9142 \pm 51 | | | | 10,485–10,221 | |
| Chugwater C2 | 9460 \pm 130 | Beta-19498 | c? | stratigraphic? | 11,176–10,304 | Maitland 1986; Lively 1996 |
| | 8960 \pm 130 | Beta-18509 | c? | stratigraphic? | 10,403–9631 | |
| average | 9075 \pm 92 | | | | 10,500–9918 | |
| Gerstle River C3 | 8820 \pm 50 | Beta-183109 | c | hearth | 10,156–9686 | Potter 2001, 2005 |
| | 8830 \pm 50 | Beta-181678 | c | hearth | 10,156–9698 | |
| | 8860 \pm 70 | Beta-133750 | c | hearth | 10,184–9700 | |
| | 8890 \pm 40 | Beta-167397 | c | hearth | 10,187–9798 | |
| | 8900 \pm 40 | Beta-181679 | c | hearth | 10,190–9896 | |
| | 8910 \pm 40 | Beta-167399 | c | hearth | 10,188–9908 | |
| | 8950 \pm 40 | Beta-167395 | c | hearth | 10,221–9917 | |
| | 9030 \pm 70 | AA-51254 | c | hearth | 10,374–9913 | |
| | 9080 \pm 50 | Beta-183108 | c | hearth | 10,386–10,176 | |
| average | 8882 \pm 17 | | | | 10,156–9911 | |
| Little Delta Dune C4 | 8880 \pm 40 | Beta-232392 | c | stratigraphic | 10,179–9789 | Potter et al. 2007b |
| XBD-306 | 8930 \pm 90 | Beta-220216 | c | stratigraphic | 10,238–9710 | Potter et al. 2007a |
| Erodeaway | 8640 \pm 170 | WSU-3683 | c | hearth | 10,184–9305 | Holmes 1988 |
| Gerstle River C4 | 8660 \pm 40 | Beta-167396 | c | hearth | 9697–9539 | Potter 2005 |
| Carlo Creek C1 | 8400 \pm 200 | WSU-1700 | c | hearth | 9910–8774 | Bowers 1980 |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2-σ range) | Reference |
|------------------------|----------------------------|-------------|-----------------------|----------------------|--------------------------|-------------------------|
| | 8690 ± 330 | Gx-5132 | c | hearth | 10,650–8793 | |
| average | 8478 ± 171 | | | | 10,113–9012 | |
| Gerstle River C5 | 7600 ± 140 | WSU-4888 | c | stratigraphic (ULD) | 8716–8047 | Potter 2005 |
| | 8280 ± 60 | Beta-98434 | c | stratigraphic (LLD) | 9442–9037 | |
| average | 8174 ± 55 | | | | 9283–9008 | |
| XBD-307 | 8070 ± 60 | Beta-220217 | c | stratigraphic | 9240–8720 | Potter et al. 2007a |
| XBD-340 | 8000 ± 50 | Beta-232399 | c | stratigraphic | 9011–8652 | Potter et al. 2007b |
| XBD-289 | 7960 ± 70 | Beta-219649 | c | stratigraphic | 9004–8610 | Potter et al. 2007a |
| Houdini Creek | 7880 ± 60 | Beta-74737 | c | stratigraphic | 8978–8552 | Bowers et al. 1995 |
| Panguingue Creek C2 | 8600 ± 200 | AA-1689 | c | stratigraphic | 10,189–9137 | Powers and Maxwell 1986 |
| | 7595 ± 405 | Gx-13012 | c | stratigraphic | 9405–7676 | |
| | 7430 ± 270 | AA-1688 | c | stratigraphic | 8971–7691 | |
| | 7130 ± 180 | Beta-15094 | c | stratigraphic | 8326–7623 | |
| | 7850 ± 180 | Beta-15093 | c | stratigraphic | 9242–8330 | |
| average | 7749 ± 97 | | | | 8970–8371 | |
| Lucky Strike site | 7760 ± 50 | Beta-196499 | c | stratigraphic | 8627–8425 | Reuther et al. 2003 |
| XBD-326 | 7740 ± 60 | Beta-219663 | c | stratigraphic | 8627–8410 | Potter et al. 2007a |
| Broken Mammoth CZ2 | 7200 ± 205 | UGa-6281D | c | hearth –split | 8401–7659 | Holmes 1996 |
| | 7600 ± 160 | WSU-4264 | c | hearth –split | 8850–8023 | |
| | 7700 ± 80 | WSU-4508 | c | hearth | 8633–8372 | |
| average | 7628 ± 78 | | | | 8592–8222 | |
| Owl Ridge C2 | 7230 ± 100 | Beta-11437 | c | stratigraphic | 8305–7850 | Phippen 1988 |
| | 7660 ± 100 | Beta-11436 | c | stratigraphic | 8643–8206 | |
| | 8130 ± 140 | Beta-5418 | c | stratigraphic | 9426–8645 | |
| average | 7584 ± 63 | | | | 8539–8213 | |
| Swan Point CZ2 | 7400 ± 80 | WSU-4426 | c | stratigraphic | 8372–8039 | Holmes et al. 1996 |
| XBD-325 | 7360 ± 40 | Beta-220682 | c | stratigraphic | 8312–8040 | Potter et al. 2007a |
| XBD-291 | 7350 ± 60 | Beta-219650 | c | stratigraphic | 8318–8024 | Potter et al. 2007a |
| Jay Creek Ridge C1 | 6970 ± 210 | Beta-7304 | c | stratigraphic | 8187–7435 | Dixon et al. 1985 |
| | 7240 ± 110 | Beta-7306 | c | stratigraphic | 8321–7853 | |
| average | 7182 ± 97 | | | | 8189–7794 | |
| Teklanika West C1 | 7130 ± 98 | Gx-18518 | c | stratigraphic | 8170–7754 | West 1996a |
| Owl Ridge C3 | 6900 ± 265 | D-3070 | c | stratigraphic | 8302–7272 | Phippen 1988 |
| | 7035 ± 380 | Gx-13009 | c | stratigraphic | 8642–7029 | |
| average | 6944 ± 217 | | | | 8195–7424 | |
| Campus Area J6 | 6850 ± 70 | Beta-97212 | c | stratigraphic | 7833–7579 | Pearson and Powers 2001 |
| XBD-313 | 6750 ± 60 | Beta-219651 | c | hearth | 7691–7504 | Potter et al. 2007a |
| Long Lake | 6606 ± 115 | UGa-949 | c | stratigraphic | 7672–7293 | Reger and Bacon 1996 |
| XBD-311 | 6490 ± 50 | Beta-220215 | c | stratigraphic | 7490–7289 | Potter et al. 2007a |
| Mead CZ2 | 6070 ± 170 | | c | stratigraphic | 7316–6505 | Dilley 1998 |
| XBD-288 | 6060 ± 60 | Beta-219654 | c | stratigraphic | 7156–6749 | Potter et al. 2007a |
| XBD-282 | 5920 ± 50 | Beta-221332 | c | stratigraphic | 6882–6644 | Potter et al. 2007a |
| Gerstle River C6 | 5050 ± 90 | N-4958 | c | stratigraphic (ULD) | 5984–5603 | Potter 2005 |
| | 6220 ± 80 | WSU-4892 | c | stratigraphic (LLD) | 7308–6912 | |
| average | 5704 ± 60 | | | | 6656–6324 | |
| Moose Creek C3 | 5680 ± 50 | Beta-106041 | c | stratigraphic | 6631–6321 | Pearson 1999 |
| XBD-317 | 5610 ± 50 | Beta-219653 | c | stratigraphic | 6485–6301 | Potter et al. 2007a |
| XBD-335 C1 | 5400 ± 40 | Beta-232391 | c | stratigraphic | 6292–6020 | Potter et al. 2007b |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2- σ range) | Reference |
|---------------------------|----------------------------|-------------|-----------------------|-----------------------------|----------------------------------|-------------------------|
| Flat Knoll C1 | 5230 \pm 140 | Beta-7695 | c | stratigraphic | 6290–5663 | Dixon et al. 1985 |
| XBD-328 | 5170 \pm 50 | Beta-219664 | c | stratigraphic | 6171–5750 | Potter et al. 2007a |
| Whitmore Ridge C2 | 5080 \pm 130 | I-4231 | c? | stratigraphic | 6179–5588 | West et al. 1996b |
| | 5480 \pm 300 | UGa-530 | so | stratigraphic | 6942–5603 | |
| average | 5143 \pm 119 | | | | 6189–5616 | |
| Butte Lake C2 | 5030 \pm 200 | Beta-10751 | c | Feature 11 | 6270–5323 | Betts 1987 |
| XBD-283 | 5000 \pm 50 | Beta-219660 | c | stratigraphic | 5893–5615 | Potter et al. 2007a |
| XBD-342 | 4670 \pm 40 | Beta-232395 | c | stratigraphic | 5575–5312 | Potter et al. 2007b |
| Swan Point CZ1B | 4620 \pm 40 | unreported | u | unreported | 5469–5087 | Holmes 2004 |
| Broken Mammoth CZ1B | 4525 \pm 90 | WSU-4458 | c | hearth –split | 5454–4873 | Holmes 1996 |
| | 4540 \pm 90 | WSU-4456 | c | hearth –split | 5466–4881 | |
| | 4545 \pm 90 | WSU-4457 | c | hearth | 5467–4833 | |
| | 4690 \pm 110 | WSU-4350 | c | hearth | 5642–5048 | |
| average | 4565 \pm 47 | | | | 5446–5047 | |
| XBD-287 | 4490 \pm 50 | Beta-219648 | c | stratigraphic | 5307–4973 | Potter et al. 2007a |
| Panguingue Creek C3 | 4510 \pm 95 | Gx-13011 | c | stratigraphic | 5448–4867 | Powers and Maxwell 1986 |
| Mount Hayes 35 | 4450 \pm 140 | unreported | u | associated with “dwelling”? | 5573–4654 | Mobley 1982 |
| Landmark Gap Trail | 4330 \pm 135 | Beta-1726 | c | stratigraphic | 5309–4533 | Mobley 1982 |
| XBD-301 | 4360 \pm 50 | Beta-219657 | c | stratigraphic | 5255–4836 | Potter et al. 2007a |
| North Gerstle Point C2 | 4290 \pm 285 | unreported | c | hearth | 5589–4092 | VanderHoek et al. 1997 |
| Jay Creek Mineral Lick C1 | 4100 \pm 60 | Beta-5464 | c | stratigraphic | 4825–4440 | Dixon et al. 1985 |
| | 4440 \pm 120 | Beta-7698 | c | stratigraphic | 5462–4729 | |
| | 4250 \pm 110 | Beta-7697 | c | stratigraphic | 5270–4443 | |
| average | 4184 \pm 48 | | | | 4844–4574 | |
| Borrow C site C1 | 4020 \pm 65 | DIC-2283 | c | stratigraphic | 4810–4294 | Dixon et al. 1985 |
| | 4570 \pm 100 | Beta-7844 | c | stratigraphic | 5531–4890 | |
| average | 4183 \pm 54 | | | | 4847–4539 | |
| XBD-343 | 4160 \pm 40 | Beta-232396 | c | stratigraphic | 4831–4571 | Potter et al. 2007b |
| Tok Terrace C1 | 4160 \pm 100 | Beta-40724 | c | stratigraphic | 4952–4419 | Sheppard et al. 1991 |
| | 4020 \pm 90 | Beta-40717 | c | stratigraphic | 4821–4248 | |
| average | 4083 \pm 67 | | | | 4821–4427 | |
| XMH-166 | 4100 \pm 270 | I-4592 | | unreported | 5434–3866 | West 1972 |
| XBD-316 | 4050 \pm 50 | Beta-219652 | c | stratigraphic | 4807–4418 | Potter et al. 2007a |
| TLM-207 C1 | 4030 \pm 220 | Beta-9897 | c | stratigraphic | 5261–3871 | Dixon et al. 1985 |
| Delta River Overlook C3 | 3980 \pm 150 | Gx-6752 | c | stratigraphic | 4840–3997 | Bacon and Holmes 1980 |
| Flat Knoll C2 | 3920 \pm 100 | Beta-7842 | c | stratigraphic | 4797–4000 | Dixon et al. 1985 |
| Rock Creek East | 3866 \pm 47 | Gx-17392 | c | unreported | 4416–4154 | McKay 1981 |
| Gerstle River C7 | 3800 \pm 70 | N-4959 | c | stratigraphic | 4413–3986 | Potter 2005 |
| Dry Creek C4 | 3430 \pm 75 | SI-2332 | c | stratigraphic | 3871–3480 | Powers et al. 1983 |
| | 3655 \pm 60 | SI-1934 | c | stratigraphic | 4149–3835 | |
| | 4670 \pm 95 | SI-1937 | c | stratigraphic | 5595–5053 | |
| average | 3783 \pm 42 | | | | 4346–3989 | |
| Little Delta River #4 | 3700 \pm 70 | Beta-123332 | c | stratigraphic | 4239–3848 | Higgs et al. 1999 |
| XBD-297 | 3620 \pm 50 | Beta-219661 | c | stratigraphic | 4088–3777 | Potter et al. 2007a |
| TLM-169 C1 | 3410 \pm 80 | Beta-10794 | c | stratigraphic | 3860–3467 | Dixon et al. 1985 |
| North Arrow site | 3220 \pm 90 | Beta-7299 | c | stratigraphic | 3684–3245 | Dixon et al. 1985 |
| | 3675 \pm 160 | Gx-5630 | c | hearth? | 4495–3586 | |
| average | 3329 \pm 78 | | | | 3819–3384 | |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2-σ range) | Reference |
|--|----------------------------|-------------------------|-----------------------|----------------------|--------------------------|--|
| Fog Creek C1 | 3160 ± 70 | Beta-7687 | c | stratigraphic | 3557–3215 | Dixon et al. 1985 |
| | 3180 ± 170 | Beta-7685 | c | stratigraphic | 3829–2957 | |
| | 3270 ± 90 | Beta-7300 | c | stratigraphic | 3716–3271 | |
| | 3290 ± 60 | Beta-7302 | c | stratigraphic | 3679–3387 | |
| | average 3239 ± 40 | | | | 3559–3382 | |
| Tuff Creek North C2 | 3210 ± 80 | DIC-2286 | c | stratigraphic | 3632–3262 | Dixon et al. 1985 |
| Usibelli | 3195 ± 295 | Gx-13013 | c | stratigraphic | 4228–2739 | Hoffecker 1985 |
| Healy Lake Village Levels 4-5 | 4010 ± 110 | Gx-2163 | c | stratigraphic | 4827–4159 | Cook 1996 |
| | 3020 ± 50 | Beta-76063 | c | stratigraphic | 3358–3074 | |
| | average 3190 ± 32 | | | | 3465–3362 | |
| XBD-336 | 3040 ± 40 | Beta-232398 | c | stratigraphic | 3362–3082 | Potter et al. 2007b |
| Fish Creek concB1, C2 | 3005 ± 135 | Gx-4110 | c | stratigraphic | 3476–2808 | Cook et al. 1977 |
| | 3065 ± 115 | Gx-4109 | c | stratigraphic | 3555–2947 | |
| | average 3040 ± 88 | | | | 3442–2976 | |
| Campus (Mobley) | 2725 ± 125 | Beta-7075 | c | stratigraphic | 3240–2488 | Mobley 1991 |
| | 2860 ± 180 | Beta-4260 | c | stratigraphic | 3450–2500 | |
| | 3500 ± 140 | Beta-6829 | c | stratigraphic | 4151–3444 | |
| average | 3025 ± 83 | | | | 3392–2968 | |
| Owl Knoll | 3010 ± 110 | Beta-123340 | c | hearth | 3443–2886 | Potter et al. 2007c; Higgs et al. 1999 |
| Healy Lake Village Levels 2-3 | 2875 ± 140 | Gx-2169 | c | stratigraphic | 3357–2753 | Cook 1996 |
| | 3580 ± 140 | Gx-2165 | c | stratigraphic | 4286–3484 | |
| | 3655 ± 426 | AU-4 | c | stratigraphic | 5270–2899 | |
| | 2660 ± 100 | Gx-2176 | c | stratigraphic | 3003–2367 | |
| | average 2965 ± 69 | | | | 3342–2952 | |
| TLM-096 | 2750 ± 215 | DIC-2285 | c | stratigraphic | 3373–2350 | Dixon et al. 1985 |
| XBD-281 | 2760 ± 40 | Beta-221333 | c | stratigraphic | 2951–2774 | Potter et al. 2007a |
| Rainbow Lake Loc. 1 | 2090 ± 130 | Gx-6009 | c | stratigraphic | 2349–1740 | Bacon and Holmes 1980 |
| | 4145 ± 240 | UGa-3172 | c | stratigraphic | 5316–3985 | |
| | average 2556 ± 114 | | | | 2855–2349 | |
| Red's Ravine | 2485 ± 75 | Beta-33300/ ETH-5901 | c | stratigraphic | 2733–2361 | |
| McCurdy archaeological site | 2410 ± 100 | Beta-14508 | c | hearth | 2744–2183 | US Bureau of Indian Affairs 1986 |
| Dixthada C1 | 2420 ± 60 | P-1834 | c | stratigraphic | 2706–2346 | Shinkwin 1979 |
| Windy Knoll site | 2340 ± 145 | DIC-1903 | c | hearth | 2745–2011 | Dixon et al. 1985 |
| Yardang Flint Station C1 | 2300 ± 180 | I-647 | c | stratigraphic | 2751–1901 | Reger et al. 1964 |
| Lake Minchumina C1 (Levels 4/5, Blueberry phase) | 1950 ± 320 | Gx-7116 | c | stratigraphic | 2716–1293 | Holmes 1986 |
| | 2365 ± 140 | UGa-634 | c | hearth | 2752–2062 | |
| | average 2298 ± 128 | | | | 2713–2005 | |
| Delta River Overlook C5 | 2285 ± 145 | Gx-6750 | c | stratigraphic | 2724–1952 | Bacon and Holmes 1980 |
| Broken Mammoth CZ1A | 2280 ± 40 | unreported | c | hearth | 2352–2157 | Holmes 2001 |
| XBD-337 | 2180 ± 40 | Beta-232400 | c | stratigraphic | 2327–2063 | Potter et al. 2007b |
| Fish Creek concA11 | 2115 ± 140 | Gx-4108 | c | Feature 1 | 2434–1719 | Cook et al. 1977 |
| Tok Terrace C2 | 1650 ± 60 | Beta-40603 | c | stratigraphic | 1696–1409 | Sheppard et al. 1991 |
| | 1980 ± 70 | Beta-40712 | c | stratigraphic | 2121–1740 | |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2- σ range) | Reference |
|--|----------------------------|--------------|-----------------------|------------------------------------|----------------------------------|--|
| | 2110 \pm 170 | Beta-40721 | c | stratigraphic | 2684–1631 | |
| | 2630 \pm 90 | Beta-40716 | c | stratigraphic | 2951–2369 | |
| | 2690 \pm 90 | Beta-42600 | c | stratigraphic | 3062–2497 | |
| | 2820 \pm 180 | Beta-40720 | c | stratigraphic | 3398–2487 | |
| average | 2114 \pm 35 | | | | 2296–1993 | |
| TLM-217 C2 | 2070 \pm 60 | Beta-9899 | c | stratigraphic | 2298–1886 | Dixon et al. 1985 |
| XBD-324 | 2070 \pm 50 | Beta-219662 | c | stratigraphic | 2287–1899 | Potter et al. 2007a |
| Portage site C1 | 2009 \pm 225 | not reported | u | unreported | 2675–1414 | West 1972 |
| XBD-296 | 2010 \pm 40 | Beta-221334 | c | stratigraphic | 2103–1876 | Potter et al. 2007a |
| XBD-286 | 1860 \pm 50 | Beta-220213 | c | stratigraphic | 1921–1634 | Potter et al. 2007a |
| Little Panguingue Creek C2 | 1825 \pm 68 | AA-1699 | c | hearth | 1896–1569 | Hoffecker and Powers 1996 |
| FAI-045 | 1820 \pm 70 | DIC-1552 | c | associated with caribou bones | 1893–1567 | Dixon et al. 1980 |
| Donnelly Ridge | 1790 \pm 300 | Beta-650 | c | stratigraphic | 2452–1062 | West 1967, 1981 |
| | 1830 \pm 200 | Beta-649 | c | stratigraphic | 2302–1338 | |
| average | 1818 \pm 166 | | | | 2135–1377 | |
| Tuff Creek North C3 | 1800 \pm 55 | DIC-2284 | so | stratigraphic | 1867–1571 | Dixon et al. 1985 |
| Hurricane Bluff C2 | 1750 \pm 40 | Beta-123338 | c | stratigraphic | 1810–1553 | Potter et al. 2007c; Higgs et al. 1999 |
| TLM-216 | 1880 \pm 50 | Beta-9892 | w | stratigraphic | 1930–1705 | Dixon et al. 1985 |
| | 1670 \pm 50 | Beta-9898 | w | stratigraphic | 1702–1417 | |
| | 1530 \pm 80 | Beta-10125 | w | stratigraphic | 1596–1293 | |
| average | 1735 \pm 32 | | | | 1714–1557 | |
| Lake Minchumina C2 (Levels 2 and 3, Cranberry phase) | 1610 \pm 150 | Gx-4233 | c | stratigraphic | 1870–1273 | Holmes 1986 |
| Watana Depression site | 1580 \pm 110 | Beta-7846 | | depression | 1715–1291 | Dixon et al. 1985 |
| Swan Point CZ1A | 1220 \pm 70 | WSU-4523 | c | stratigraphic | 1285–982 | Holmes et al. 1996 |
| | 1570 \pm 70 | WSU-4524 | c | stratigraphic | 1607–1316 | |
| | 1670 \pm 60 | WSU-4522 | c | stratigraphic | 1706–1415 | |
| | 1750 \pm 80 | WSU-4521 | c | stratigraphic | 1872–1423 | |
| average | 1552 \pm 34 | | | | 1526–1369 | |
| Brown Scraper Kame site | 1420 \pm 70 | Beta-5653 | c | associated with calcined bone | 1515–1181 | Dixon et al. 1985 |
| Mead CZ1 | 1420 \pm 60 | WSU-4348 | c | stratigraphic | 1507–1184 | Dilley 1998 |
| Red Scraper site C2 | 1380 \pm 155 | DIC-2246 | c | stratigraphic | 1603–961 | Dixon et al. 1985 |
| Borrow C site C3 | 1260 \pm 80 | Beta-7845 | c | Feature 1 | 1305–983 | Dixon et al. 1985 |
| | 1400 \pm 55 | DIC-2245 | c | stratigraphic | 1405–1184 | |
| average | 1355 \pm 45 | | | | 1345–1179 | |
| Healy Lake Garden | 1260 \pm 90 | GaK-1885 | c | hearth? | 1312–975 | Cook 1969 |
| | 1270 \pm 80 | GaK-1884 | c | hearth? | 1311–985 | |
| average | 1266 \pm 60 | | | | 1294–1063 | |
| East Cove site | 1360 \pm 120 | Gx-5129 | c | house floor? | 1522–1002 | Holmes 1986 |
| | 1140 \pm 135 | Gx-5997 | c | hearth | 1300–788 | |
| average | 1262 \pm 90 | | | | 1314–975 | |
| Mount Hayes 130 | 1220 \pm 190 | I-4232 | c | possible association with housepit | 1515–742 | West 1972 |
| XBD-290 | 1170 \pm 40 | Beta-219655 | c | stratigraphic | 1223–975 | Potter et al. 2007a |
| Kosina Creek B | 1160 \pm 100 | DIC-1878 | c | hearth | 1290–918 | Dixon et al. 1985 |
| Lake Minchumina C3 (Level 1, Rasp-berry phase) | 1140 \pm 120 | Gx-2828 | b | from hearth | 1292–797 | Holmes 1986 |
| Flood's Cabins Cache Pits | 1050 \pm 60 | WSU-2584 | c | from storage pit | 1118–795 | Holmes 1986 |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (Continued)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2-σ range) | Reference |
|--|----------------------------|-------------|-----------------------|-------------------------------------|--------------------------|----------------------------|
| Chugwater C3 | 870 ± 50 | Beta-7565 | c | hearth? | 910–693 | Maitland 1986; Lively 1996 |
| | 950 ± 105 | Beta-9248 | c | hearth? | 1059–680 | |
| | 1120 ± 90 | Beta-7566 | c | hearth | 1269–803 | |
| average | 932 ± 40 | | | | 928–744 | |
| Flat Knoll C3 | 840 ± 60 | Beta-7692 | c | stratigraphic | 908–674 | Dixon et al. 1985 |
| | 1060 ± 70 | Beta-7693 | c | stratigraphic | 1170–796 | |
| average | 933 ± 46 | | | | 929–743 | |
| Healy Lake Village Level 1 | 900 ± 90 | GaK-1886 | c | hearth | 962–675 | Cook 1996 |
| Nenana River Dune site | 800 ± 50 | Beta-196497 | c | from cache pit | 896–663 | Potter 2004b |
| Little Bones Ridge | 740 ± 70 | DIC-2253 | c | structural timber | 793–552 | Dixon et al. 1985 |
| Ringling site | 460 ± 100 | Gx-4391 | c | from storage pit 770302 | 652–305 | Workman 1976 |
| | 695 ± 115 | Gx-4300 | c | from storage pit 29 | 904–509 | |
| | 760 ± 125 | Gx-4299 | c | from storage pit 50 | 924–538 | |
| | 765 ± 125 | Gx-4298 | c | from storage pit 29 | 925–540 | |
| | 720 ± 60 | WSU-4922 | c | timbers from housepit 95-36 | 762–553 | Hanson 1999 |
| average | 780 ± 70 | WSU-4923 | c | hearth | 905–563 | Hanson 1999 |
| | 707 ± 36 | | | | 721–562 | |
| Birches site | 640 ± 95 | I-2617 | c | from house 5 | 737–504 | West 1978 |
| Tok Terrace Cluster G, upper component | 640 ± 70 | Beta-34233 | c | from burned rock layer in steambath | 451–0 | Gerlach et al. 1989 |
| GUL-076 | 550 ± 135 | Gx-3855 | c | hearth | 738–300 | Clark 1974 |
| | 690 ± 135 | Gx-3859 | c | planking from housepit | 921–498 | |
| average | 620 ± 95 | | | | 730–502 | |
| Tok Terrace C3 | 920 ± 90 | Beta-40722 | c | stratigraphic | 1042–677 | Sheppard et al. 1991 |
| | 570 ± 80 | Beta-40713 | c | stratigraphic | 677–495 | |
| | 450 ± 90 | Beta-40718 | c | stratigraphic | 643–305 | |
| average | 640 ± 50 | | | | 671–545 | |
| Dixthada C2 | 770 ± 40 | P-1832 | c | base of midden | 764–661 | Shinkwin 1979 |
| | 390 ± 50 | P-1833 | c | base of midden | 515–315 | |
| average | 622 ± 31 | | | | 659–551 | |
| Batzulnetas Village | 410 ± 80 | Beta-56552 | c | unreported | 619–296 | AHRS (n.d.) |
| | 570 ± 100 | Beta-56551 | c | unreported | 722–322 | |
| average | 472 ± 62 | | | | 639–319 | |
| VAL-206 | 460 ± 70 | Beta-6692 | c | hearth | 635–316 | Reger 1985 |
| TLM-253 | 430 ± 130 | Beta-10796 | c | associated with FCR, calcined bone | 668–0 | Dixon et al. 1985 |
| TLM-250 | 370 ± 80 | Beta-10798 | c | hearth | 535–156 | Dixon et al. 1985 |
| Nenana River Gorge site C1 | 460 ± 115 | I-9882 | c | hearth | 666–288 | Plaskett 1977 |
| | 260 ± 75 | I-9883 | c | hearth | 498–0 | |
| average | 320 ± 63 | | | | 505–154 | |
| Tsusena Creek C1 | 300 ± 70 | DIC-2252 | c | hearth | 506–0 | Dixon et al. 1985 |
| O'Brian Creek | 280 ± 60 | Beta-56665 | c | unreported | 496–0 | Reger et al. 1975 |

Table A1 Radiocarbon-dated component date list (see Methods section for details). (*Continued*)

| Component ^a | ¹⁴ C assay (BP) | Lab # | Material ^b | Context ^c | Calib. yr BP (2- σ range) | Reference |
|-----------------------------|-------------------------------|-----------|-----------------------|----------------------|-------------------------------------|-------------------|
| Permafrost Creek site C2 | 280 ± 110 | DIC-1905 | c | hearth | 511–0 | Dixon et al. 1985 |
| | 280 ± 245 | DIC-1904 | c | hearth | 650–0 | |
| average | 280 ± 100 | | | | 510–0 | |
| Bendildenden | 260 ± 70 | Beta-6828 | c | hearth | 497–0 | Reger 1985 |
| Butte Lake C4 | 110 ± 60 | DIC-3068 | c | Feature 2 | 281–0 | Betts 1987 |
| | 180 ± 60 | DIC-3069 | c | Feature 6 | 307–0 | |
| average | 145 ± 42 | | | | 283–0 | |

^aSite prefixes relate to USGS 250,000 scale quadrangles (ANC – Anchorage, FAI – Fairbanks, HEA – Healy, GUL – Gulkana, NAB – Nabesna, TLM – Talkeetna Mountains, VAL – Valdez, XBD – Big Delta, XMH – Mount Hayes). Components are labeled as C# – component, CH# – cultural horizon, CZ# – cultural zone.

^bMaterial: c – charcoal; iv – ivory; o – organic residue; so – soil organics; u – unreported.

^cAbbreviations: ULD – upper limiting date; LLD – lower limiting date; FCR – fire-cracked rock.