# Off-centre gravity induces large-scale flow patterns in spherical Rayleigh-Bénard convection 

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#### Abstract

We perform direct numerical simulations to study the effect of the gravity centre offset in spherical Rayleigh-Bénard convection. When the gravity centre is shifted towards the south, we find that the shift of the gravity centre has a pronounced influence on the flow structures. At low Rayleigh number $R a$, a steady-state large-scale meridional circulation induced by the baroclinic imbalance, created by the misalignment of the gravity potentials and isotherms, is formed. At high $R a$, an energetic jet is created on the northern side of the inner sphere that is directed towards the outer sphere. The large-scale circulation induces a strong co-latitudinal dependence in the local heat flux. Nevertheless, the global heat flux is not affected by the changes in the large-scale flow organization induced by the gravity centre offset.


Key words: Bénard convection, turbulent convection, plumes/thermals

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## 1. Introduction

The effect of an off-centre gravity location on the flow structures in turbulent thermal convection has not been studied previously. Here, we use spherical Rayleigh-Bénard (RB) convection as an idealized model system to study the effect of an off-centre gravity location on thermal convection in a spherical system. The system consists of a fluid layer enclosed between two spherical shells, heated from the inner sphere and cooled from the outer one. Convection in spherical shells differs from the classical RB convection (Ahlers, Grossmann \& Lohse 2009; Lohse \& Xia 2010; Chillà \& Schumacher 2012) owing to the curvature of the boundaries, non-uniform gravity, and the geometrical asymmetry between the boundary layer at the inner and outer spheres (Busse 1970; Spiegel 1971; Jarvis, Glatzmaierand \& Vangelov 1995; Tilgner 1996; Shahnas et al. 2008; Deschamps, Tackley \& Nakagawa 2010; O’Farrell, Lowman \& Bunge 2013; Gastine, Wicht \& Aurnou 2015).

Most studies on spherical RB convection focus on the case in which the gravity centre coincides with the geometric one. Busse (1975) used perturbation analysis to show the qualitative difference between convective flow patterns of odd and even spherical harmonic order just above the onset of convection. Subsequently, in an infinite Prandtl number ( Pr ) fluid with application to mantle convection, Bercovici, Schubert \& Glatzmaier (1989) showed that these convective patterns persist for Rayleigh numbers ( $R a$ ) up to 100 times larger than the critical value. For $R a>10^{5}$, the axisymmetric convective patterns break down when the flow starts to show time-dependent behaviour (Iwase \& Honda 1997; Deschamps et al. 2010; Futterer et al. 2013; Yanagisawa, Kameyama \& Ogawa 2016). In spherical RB convection, the influence of curvature is important. For constant $R a$, the length scale of the preferentially convective patterns decreases with increasing radius ratio (Deschamps et al. 2010). Yanagisawa \& Yamagishi (2005) find that the convective rolls become smaller with increasing $R a$ when the radius ratio is kept constant. Gastine et al. (2015) show that for $\operatorname{Pr}=1$, sheet-like thermal plumes are formed near the inner and outer spheres, and these plumes undergo morphological changes into mushroom-like plumes when they eject from the boundary layers to the bulk. These plume dynamics are similar to the situation for classic RB convection (Puthenveettil \& Arakeri 2005; Shishkina \& Wagner 2008; Zhou \& Xia 2010; Chillà \& Schumacher 2012). However, due to the asymmetries between the hot and cold surfaces in spherical shells, the mushroom-like plumes emitted from the outer sphere are thicker than those emitted from the inner sphere (Gastine et al. 2015). Nevertheless, studies have shown that there are similarities between spherical and classical planar RB convection. For example, Gastine et al. (2015) showed that the effective heat transport (characterized by the Nusselt number $N u$ ) scaling exponent $\alpha$ in $N u \sim R a^{\alpha}$ in spherical RB convection has a similar $R a$ number dependence as in planar RB convection, where for $R a \lesssim 10^{11}$, typically $0.28 \lesssim \alpha \lesssim 0.31$ (Ahlers et al. 2009).

However, the misalignment between the gravity centre and the geometric one in the spherical RB convection set-up has been overlooked to date. In this study, we use three-dimensional direct numerical simulations to investigate the effect of the gravity centre offset on the flow structures and heat transfer in spherical RB convection. The paper is organized as follows. In $\S 2$, we explain the numerical method and discuss the considered parameter regime. In §§3.1-3.3, we discuss the large-scale flow pattern induced by the off-centre gravity. The effect of the large-scale structure on heat transfer is shown in § 3.4. In §3.5, we study the effect of the gravity profile and $R a$ on flow structures and heat transfer. The paper ends with a summary of our findings and conclusions in § 4.


Figure 1. Schematic of spherical RB convection in between two concentric spheres with radius ratio $\eta=r_{i} / r_{o}$ between the inner and outer sphere, and gap size $d=r_{o}-r_{i}$. A no-slip boundary condition with constant temperature is used on the inner (hot) and outer (cold) sphere. In spherical coordinates the longitudinal, co-latitudinal and radial directions are represented by $\hat{\theta}, \hat{\varphi}$ and $\hat{r}$, respectively. The gravity centre $G$ is offset from the geometric centre $O$, and $P$ is an arbitrary fluid point in the spherical shell.

## 2. Numerical method and parameters

### 2.1. Set-up of convection in spherical shells

The spherical RB geometry is illustrated schematically in figure 1. Fluid fills a spherical shell between the inner sphere of radius $r_{i}$ and the outer sphere of radius $r_{o}$. The radius ratio between the inner and outer spheres is given by

$$
\begin{equation*}
\eta=r_{i} / r_{o} \tag{2.1}
\end{equation*}
$$

In this study, we consider a fixed aspect ratio $\eta=0.3$. The surface temperatures of the inner and outer spheres are kept constant at $T_{i}$ and $T_{o}$, respectively, with $T_{i}>T_{o}$. No-slip boundary conditions are imposed on both boundaries. Point $G$ indicates the gravity centre, which is offset from the geometric centre $O$. The gravity shift is defined as

$$
\begin{equation*}
\varepsilon=|O G| / r_{i} \tag{2.2}
\end{equation*}
$$

where $|O G|$ indicates the displacement of the gravity compared to the geometrical centre; see figure 1. In the absence of rotation (and magnetic field), for a sphere, there are no north and south directions, therefore any diametrically opposite symmetry is equivalent to another. Nevertheless, since the spherical coordinates are naturally clustered around the polar axis, it is beneficial to consider the jet aligned along this axis, referred to as north-south, from a computation point of view. For convenience, we always shift the gravity centre $G$ to the south.

The dynamics of spherical RB convection is controlled by the Rayleigh and Prandtl numbers

$$
\begin{equation*}
R a=\frac{\beta g_{o} \Delta T d^{3}}{\kappa v}, \quad \operatorname{Pr}=\frac{v}{\kappa} \tag{2.3a,b}
\end{equation*}
$$

where $\beta$ is the thermal expansion coefficient, $g_{o}$ is the surface-averaged gravity at the outer sphere, $v$ is the kinematic viscosity, and $\kappa$ is the thermal diffusivity. We normalize the fields and distances by the length scale $d=r_{o}-r_{i}$, the temperature difference $\Delta T$ between the inner and outer spheres, and the free-fall velocity $U=\sqrt{\beta g_{o} \Delta T d}$.

### 2.2. Underlying dynamical equation and numerical method

We solve the Navier-Stokes equations under the Boussinesq approximation in spherical coordinates, which in dimensionless form read

$$
\begin{gather*}
\frac{\partial u}{\partial t}+\boldsymbol{u} \cdot \nabla \boldsymbol{u}=-\nabla p+\sqrt{\frac{P r}{R a}} \nabla^{2} \boldsymbol{u}+C_{g} g T, \quad \nabla \cdot \boldsymbol{u}=0,  \tag{2.4}\\
\frac{\partial T}{\partial t}+\boldsymbol{u} \cdot \nabla T=\frac{1}{\sqrt{R a P r}} \nabla^{2} T, \tag{2.5}
\end{gather*}
$$

where $\boldsymbol{u}, p, T$ and $\boldsymbol{C}_{g} g$ denote the fluid velocity, pressure, temperature and gravitational acceleration acting in the three directions. The coefficients $C_{g, i}$ (where $i=1,2,3$ correspond to the $\hat{\theta}, \hat{r}$ and $\hat{\varphi}$ directions, respectively) denote the decomposition of $g$, which is detailed in Appendix A. Since we always shift the gravity centre $G$ to the south, the three components of the buoyancy coefficient can be written as

$$
\begin{equation*}
C_{g} g=\left(0, r+r_{i} \varepsilon \cos \varphi,-\varepsilon r_{i} \sin \varphi\right) r_{o}^{-n_{g}} L_{G P}^{n_{g}-1} \tag{2.6}
\end{equation*}
$$

in which $L_{G P}=|\boldsymbol{G P}|$ indicates the distance between a fluid point and the gravity centre; see figure 1 . We consider the gravity profile

$$
\begin{equation*}
g=\left(L_{G P} / r_{o}\right)^{n_{g}} \tag{2.7}
\end{equation*}
$$

and we use $n_{g}=-2,-1,0$ and 1 . The value of $n_{g}$ can be inferred by physical considerations on the mass ratio between the nucleus ( $r<r_{i}$ ) and the spherical shell ( $r_{i}<$ $\left.r<r_{o}\right)$. The gravity is constant $\left(n_{g}=0\right)$ when the mass of the spherical shell is negligible. This assumption is typically used to model the Earth's mantle (Bercovici et al. 1989). If the mass is condensed centrally, then one obtains $n_{g}=-2$. Gastine et al. (2015) showed that only for this gravity distribution function is there an analytical relation between the viscous dissipation rate $\left(\epsilon_{U}\right)$ and $N u$, i.e. $\epsilon_{U}=\left(3 /\left(1+\eta+\eta^{2}\right)\right)(1 / \operatorname{Pr})(N u-1)$. When the density is constant within the fluid layer, and no mass is contained in the inner sphere, gravity is directly proportional to the radial coordinate $r\left(n_{g}=1\right)$ (Tilgner 1996).

The governing equations (2.4) and (2.5) are discretized by a staggered central second-order finite-difference scheme in spherical coordinates. The numerical scheme is based on the method by Verzicco \& Orlandi (1996), which has been extended recently to spherical coordinates by Santelli, Orlandi \& Verzicco (2021). The advantage of this scheme is that it allows arbitrary non-uniform grids in the radial and co-latitudinal directions. Here, it is worthwhile to mention that singularities at the poles are prevented by introducing a new set of quantities $\left(u_{\theta}, u_{r} r^{2}, u_{\varphi} \sin \varphi\right)$, which results in trivial boundary conditions $u_{\varphi} \sin \varphi=0$ at the north pole $(\varphi=0)$ and south pole $(\varphi=\pi)$ in the co-latitudinal direction. The code has been validated carefully by Santelli et al. (2021), and we refer to that work for details on the method. For additional validation results relevant to the flows considered here, we refer the reader to Appendix B.

### 2.3. Choice of parameters

To study the effect of the gravity centre location on the flow dynamics in spherical RB convection, we considered different gravity distributions ( $n_{g} \in\{-2,-1,0,1\}$, see (2.7)) and gravity centre offset ( $\varepsilon$ from 0 to 0.8 , see (2.2)) for $R a=10^{7}$ and $10^{8}$. All simulations in this study are for $\operatorname{Pr}=1$. To ensure that the flow is resolved fully, we place a sufficient number of computational grid points in the bulk and the boundary
layers (Stevens, Verzicco \& Lohse 2010). As we will use spectral analysis to study the flow structures, we use a uniform grid in the longitudinal and co-latitudinal directions. In the radial direction, the grid cells are clustered towards the inner and outer spheres to ensure that the boundary layers are resolved adequately (Shishkina et al. 2010).

We calculate the Nusselt number $N u$ from the normalized averaged temperature gradients at the inner and outer spheres as follows:

$$
\begin{equation*}
N u_{i}=-\left.\eta \frac{\mathrm{d} \overline{\langle T\rangle_{\theta, \varphi}}}{\mathrm{d} r}\right|_{r_{i}}, \quad N u_{o}=-\left.\frac{1}{\eta} \frac{\mathrm{~d} \overline{\langle T\rangle_{\theta, \varphi}}}{\mathrm{d} r}\right|_{r_{o}} \tag{2.8a,b}
\end{equation*}
$$

Here, $\langle\cdots\rangle_{\theta, \varphi}$ represents the average over a spherical surface with fixed $r$, and $\cdots$ indicates a time average. Also, $N u_{i}$ and $N u_{o}$ as functions of the co-latitudinal direction are defined as

$$
\begin{equation*}
N u_{i}(\varphi)=-\left.\eta \frac{\mathrm{d} \overline{\langle T\rangle_{\theta}}}{\mathrm{d} r}\right|_{r_{i}}, \quad N u_{o}(\varphi)=-\left.\frac{1}{\eta} \frac{\mathrm{~d} \overline{\langle T\rangle_{\theta}}}{\mathrm{d} r}\right|_{r_{o}}, \tag{2.9a,b}
\end{equation*}
$$

where $\langle\cdots\rangle_{\theta}$ represents the average over the azimuthal direction. The difference in $N u$ values obtained at the inner and outer spheres is always less than $0.2 \%$. Also, we verify that $N u$ calculated at the spheres is the same as the value obtained from the volume-averaged Nusselt number $N u_{v}$, which is calculated from the surface-averaged Nusselt number $N u_{h}$ averaged over concentric spherical surfaces with fixed $r$, with

$$
\begin{equation*}
N u_{h}(r)=\frac{\sqrt{\operatorname{RaPr}}\left\langle u_{r} T\right\rangle_{\theta, \varphi}-\partial_{r}\langle T\rangle_{\theta, \varphi}}{-\partial_{r} T_{c}}, \tag{2.10}
\end{equation*}
$$

where $T_{c}(r)=\eta /\left[(1-\eta)^{2} r\right]-\eta /(1-\eta)$ is the purely conductive temperature profile for constant temperature boundary conditions. Subsequently, we perform radial integration and time averaging to obtain $N u_{v}$ as

$$
\begin{equation*}
N u_{v}=\overline{\frac{1}{d} \int_{r_{i}}^{r_{o}} N u_{h}(r) \mathrm{d} r} \tag{2.11}
\end{equation*}
$$

The volume-averaged root-mean-square (r.m.s.) Reynolds number is given by

$$
\begin{equation*}
R e_{r m s}=\sqrt{R a / P r} \sqrt{\left\langle u_{i} u_{i}\right\rangle}, \quad i=1,2,3 . \tag{2.12}
\end{equation*}
$$

Here, $\langle\cdots\rangle$ denotes the average over the whole volume of the spherical shell. The details of the simulations considered in this study are summarized in table 1.

## 3. Results

Before we discuss the turbulent flows in §§3.2-3.5, we first focus on the large-scale circulation formation induced by off-centre gravity in a steady axisymmetric flow obtained at $R a=10^{3}$. As indicated in (2.6), the radial and co-latitudinal buoyancy coefficients depend on the gravity centre offset $\varepsilon$ and the gravity distribution $n_{g}$. To disentangle these effects, for turbulent flows we first examine the effect of the gravity centre offset for constant gravity $\left(n_{g}=0\right)$ at $R a=10^{8}$ in $\S \S 3.2-3.4$. The effect of $n_{g}$ and $R a$ is studied in § 3.5.

| $R a$ | $n_{g}$ | $N_{\theta} \times N_{r} \times N_{\varphi}$ | $\varepsilon$ | $N u$ | $N u_{v}$ | $R e_{r m s}$ |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| $10^{6}$ | 0 | $241 \times 71 \times 121$ | 0.01 | 4.89 | 4.88 | 151.4 |
|  |  |  | 0.05 | 4.86 | 4.86 | 152.6 |
| $10^{7}$ | -2 | $501 \times 101 \times 251$ | 0 | 16.09 | 16.05 | 919.5 |
|  |  |  | 0.2 | 16.18 | 16.17 | 965.7 |
|  |  |  | 0.4 | 16.23 | 16.22 | 976.9 |
| $10^{7}$ | -1 | $401 \times 91 \times 201$ | 0.8 | 16.20 | 16.08 | 936.8 |
|  |  |  | 0.2 | 11.89 | 11.90 | 657.5 |
|  |  |  | 0.4 | 11.96 | 11.95 | 692.5 |
|  |  |  | 0.8 | 11.41 | 11.14 | 526.6 |
| $10^{7}$ | 0 | $301 \times 81 \times 151$ | 0 | 8.73 | 8.77 | 485.6 |
|  |  |  | 0.01 | 8.92 | 8.91 | 486.1 |
|  |  |  | 0.05 | 8.94 | 8.93 | 493.6 |
|  |  |  | 0.4 | 8.91 | 8.91 | 497.6 |
|  |  |  | 0.8 | 8.81 | 8.96 | 494.4 |
| $10^{7}$ | 1 | $321 \times 85 \times 161$ | 0 | 7.04 | 7.05 | 358.3 |
|  |  |  | 0.2 | 6.93 | 6.89 | 356.3 |
|  |  |  | 0.4 | 7.06 | 7.03 | 359.5 |
| $10^{8}$ | 0 | $501 \times 101 \times 251$ | 0 | 17.36 | 17.38 | 1487.3 |
|  |  |  | 0.01 | 17.36 | 17.35 | 1487.9 |
|  |  |  | 0.05 | 17.37 | 17.42 | 1501.1 |
|  |  |  | 0.1 | 17.30 | 17.29 | 1505.2 |
|  |  |  | 0.2 | 17.15 | 17.15 | 1529.0 |
|  |  |  | 0.4 | 17.31 | 17.24 | 1513.7 |
|  |  |  | 16.62 | 16.53 | 1302.8 |  |

Table 1. Details of the simulations. The columns from left to right indicate: $R a$, the gravity profile exponent $n_{g}$ (see (2.7)); the number of grid points in the longitudinal, radial and co-latitudinal directions, $N_{\theta} \times N_{r} \times N_{\phi}$; the shift of the gravity centre with respect to the geometrical centre $\varepsilon$ (see (2.1)); the average heat transfer $N u$ across the inner $N u_{i}$ and outer $N u_{o}$ spheres (see (2.8a,b)); $N u_{v}$ (see (2.11)); and $R e_{r m s}$ (see (2.12)).

### 3.1. Dynamics of steady flow

When the gravity centre offsets with the geometric one in spherical RB convection, the crossing between temperature gradient $(\nabla T)$ and gravitational acceleration ( $g$ ) results in the baroclinic torque, which plays an important role in the generation of vorticity by density variations. Figure $2(a)$ shows gravitational potential isolines to intersect the isotherms in an axisymmetric flow. The curl of the buoyancy term $(\nabla \times \boldsymbol{g} T)$ denotes the Boussinesq baroclinic torque under the Boussinesq approximation (Schladow, Patterson \& Street 1989; Heifetz, Maor \& Guha 2020). As shown in figure 2(b), the baroclinic torque is strongest in the vicinity of the inner sphere where the angle between $\nabla T$ and $g$ is the largest. The baroclinic torque drives a northward-directed flow along the inner sphere. This results in a large-scale circulation pattern as shown in figure 2(c). In steady flows, large-scale circulations induced by baroclinic torque are also reported for side-heated cavity (Jiang, Sun \& Calzavarini 2019), stratified flow in a semicylindrical cavity (Kumar \& Pothérat 2020), and natural convection in spherical annuli (Scurtu, Futterer \& Egbers 2008).

For different $n_{g}$, figure $3(a)$ shows the radial $\left(C_{g, r} g \hat{r}\right)$ and co-latitudinal ( $C_{g, \varphi} g \hat{\varphi}$ ) buoyancy coefficients for $\varepsilon=0.8$ on the mid-plane of the spherical shell. It shows that $g$ increases monotonically from the north to the south pole when $n_{g}<0$. For $n_{g}>0$,

(b)

(c)


Figure 2. Baroclinic imbalance induces a regular large-scale circulation due to the offset gravity centre towards the south $\left(\varepsilon=0.2\right.$, see (2.2)) for a steady flow with $R a=10^{3}$. (a) The misalignment of the gravitational potentials (blue dashed circles) and isotherms (red circles) in a meridional cut. The direction is indicated by the blue arrows. (b) The corresponding Boussinesq baroclinic torque $\nabla \times g T$. Two dashed arrows correspond to the direction of the baroclinic torque. (c) The vector field (coloured by the velocity magnitude) and streamlines (grey curves) of the velocity in the meridian plane.
the opposite trend is observed, while for $n_{g}=0$, the gravity is constant as function of the co-latitude location. We keep $R a=10^{3}$ and vary $n_{g}$ from -2 to 1 . The baroclinic torque $(\nabla \times g T)$ for $\varepsilon=0.8$ is shown in figure $3(b): \nabla \times g T$ concentrates in the south polar region when $n_{g}=-2$, whereas the strongest $\nabla \times g T$ gradually moves to the equator with increasing $n_{g}$. The corresponding large-scale circulation is shown in figure $3(c)$; the velocity field strengthens close to the inner sphere and its highest value moves from the south to the equator, which is the same trend as for the baroclinic torque. In addition, the magnitude of the velocity field decreases with increasing $n_{g}$ due to the decreasing mean buoyancy power ( $\overline{\left\langle g u_{r} T\right\rangle}$, see Gastine et al. 2015) averaged over the spherical shell.

### 3.2. Generation of convective jet and large-scale circulation

Figure 4 shows the temperature field on the meridional cut in figures $4(a-d)$, the equatorial cut in figures $4(e-h$, and the surfaces corresponding to the locations of the outer and inner thermal boundary layer heights in figures $4(i-p)$. Figures $4(b-d)$ show that for $\varepsilon \geq 0.05$, the jet ejects hot plumes from the inner sphere towards the outer sphere. When $\varepsilon$ is larger, e.g. $\varepsilon \geq 0.4$, thermal plumes preferably ascend northwards along the inner sphere, as shown in figures $4(o, p)$. This happens because the plumes that form along the southern part of the inner sphere move upwards, following the large-scale flow pattern induced by the baroclinic torque; see figure $2(c)$. Consequently, the plumes combine at the northern part of the inner sphere to form the northbound plume towards the outer sphere (figures $4 c, d$ ). This creates a convective jet impinging on the north pole of the outer sphere, after which thermal plumes descend southwards along the outer sphere, as shown in figures $4(k, l)$. As a result, a limited number of plumes are ejected radially from the hot sphere to the cold sphere in the equatorial region, as shown in figures $4(g, h)$. In addition, we visualize the advection of temperature fluctuation by baroclinic flow $U \cdot \nabla T^{\prime}$ as shown in figures $4(q-t)$. We find that $U \cdot \nabla T^{\prime}$ is intense where $T$ is high, which indicates that baroclinic flow works on the temperature perturbation to promote the generation of thermal plumes on the northern side of the sphere. This confirms that the plumes combine at the northern part of the inner sphere to form the northbound plume towards the outer sphere.


Figure 3. (a) The radial ( $C_{g, r} g \hat{r}$, red arrow) and co-latitudinal ( $C_{g, \varphi} g \hat{\varphi}$, blue arrow) decomposition of buoyancy $g \boldsymbol{G P}$ (black arrow) in (2.6) for $\varepsilon=0.8$. Different panels indicate the buoyancy components on the mid-plane of the spherical shell for $n_{g}=-2,-1,0,1$; see (2.7). The lengths of the arrows are scaled by the co-latitudinal averaged gravity in the mid-plane. $(b, c)$ Baroclinic imbalance induces regular large-scale circulation due to the offset gravity centre $\varepsilon=0.8$ for a steady flow with $R a=10^{3}$ and different gravity profiles. (b) The Boussinesq baroclinic torque $\nabla \times \boldsymbol{g} T$ in a meridional cut. (c) The vector field (coloured by the velocity magnitude) and streamlines (grey curves) of the velocity in the meridian planes.

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Figure 4. $(a-p)$ Snapshots of the temperature fields for $R a=10^{8}, n_{g}=0$ and $\varepsilon=0,0.05,0.4,0.8$. (a-d) Meridional cut and temperature isosurface for $T=0.2$. ( $e-h$ ) Temperature field in the equatorial plane, and $(i-l, m-p)$ temperature field at the outer (inner) thermal boundary layer heights along the inner and outer spheres, displayed in hammer projections. ( $q-t$ ) Advection of temperature fluctuation by the baroclinic flow $U \cdot \nabla T^{\prime}$ at the outer thermal boundary layer height.

We compare the unsteady flow fields for $\varepsilon=0.01$ and 0.05 , and different $R a$, in figures $5(a)$ and $5(b)$, respectively. The $R a=10^{3}$ cases are shown as reference in the leftmost panel. For the very small $\varepsilon=0.01$ in figure $5(a)$, the upward convective jet is the strongest for $R a=10^{7}$. However, the convective jet and meridional circulation are weakened significantly for large $R a=10^{8}$. For $\varepsilon=0.05$, figure $5(b)$ shows that the convective jet is the strongest for $R a=10^{6}$. Again, the convective jet weakens with increasing $R a$. In figure $5(c)$, we compare different $\varepsilon$ for $R a=10^{8}$. As expected, the convective jet is enhanced significantly with increasing $\varepsilon$. However, the shape of the large-scale circulation is almost unchanged for $\varepsilon \leq 0.4$. When $\varepsilon$ increases to 0.8 , the large-scale circulation becomes a crescent shaped circulation, which results in its centre moving northwards.


Figure 5. The vector field (coloured by the velocity magnitude) and streamlines (grey curves) of the longitudinal and time-averaged velocity in the meridian plane for: $(a) \varepsilon=0.01$, and $(b) \varepsilon=0.05$, with increasing $R a=10^{3}$ (steady flow), $R a=10^{6}, R a=10^{7}$ and $R a=10^{8}$; (c) $R a=10^{8}$ and $n_{g}=0$ with increasing $\varepsilon=0.1,0.2,0.4,0.8$.

### 3.3. Modal analysis to identify large-scale structures

When the gravity centre is moved towards the south pole (increasing $\varepsilon$ ), the total turbulent kinetic energy (TKE), which is defined as

$$
\begin{equation*}
\mathrm{TKE}=\frac{1}{2} \frac{R a}{P r} \overline{\left\langle u_{i} u_{i}\right\rangle}, \quad i=1,2,3, \tag{3.1}
\end{equation*}
$$



Figure 6. (a) Volume-averaged $\operatorname{TKE}(\varepsilon) / \operatorname{TKE}(\varepsilon=0)$ as a function of $\varepsilon$ for $R a=10^{8}$ and $n_{g}=0$. The error bars indicate the corresponding standard deviations. (b) Plot of $\left\langle C_{m=0}\right\rangle$ of TKE for $\ell>0$, normalized by $\left\langle C_{m=0}\right\rangle$ for $\varepsilon=0$. The inset shows $\left\langle C_{\ell, m=0}\right\rangle$, normalized by $\left\langle C_{\ell, m=0}\right\rangle$ for $\varepsilon=0$, as a function of the spherical harmonic order $\ell$, for different $\varepsilon$.
first increases slightly $(\varepsilon \leq 0.4)$ and then decreases considerably $(\varepsilon=0.8)$ (see figure $6 a)$. To study the effect of the off-centre gravity on the distribution of TKE, we perform a modal analysis of the flow. In spherical coordinates, a surface spherical harmonic function is the expansion of plane waves, where the order $m$ and degree $\ell$ represent the wavenumbers along the longitudinal and co-latitudinal directions, respectively. Here, we define the power spectrum on a spherical surface as

$$
\begin{equation*}
C_{\ell, m}=a_{\ell, m}^{2}+b_{\ell, m}^{2} \tag{3.2}
\end{equation*}
$$

where $a_{\ell, m}$ and $b_{\ell, m}$ are the spherical harmonics expansion coefficients, which are detailed in Appendix C. We define the power spectrum in the longitudinal direction as

$$
\begin{equation*}
C_{m}=\frac{1}{n-m+1} \sum_{\ell=m}^{n} C_{\ell, m}, \quad n=\min \left(N_{\varphi}-1,\left(N_{\theta}+1\right) / 2\right) \tag{3.3}
\end{equation*}
$$

where $N_{\varphi}$ and $N_{\theta}$ represent the numbers of grid points in the co-latitudinal and longitudinal directions, respectively. Radial integration and temporal averaging are used to obtain $\overline{\langle\cdots\rangle}_{r}$. The large-scale circulation is represented by ( $m=0, \ell>0$ ) modes as demonstrated in Appendix C.

Figure $6(a)$ shows that the average TKE is almost unchanged for $\varepsilon \leq 0.4$, but the distribution among the different modes is changed. As we have seen in figure 5, a large-scale circulation is formed when $\varepsilon \geq 0.05$ for $R a=10^{8}$. Figure $6(b)$ shows that the axisymmetric $(m=0)$ mode, which indicates the large-scale mode, increases strongly with $\varepsilon$. For $\varepsilon=0.8$, both the $m=0$ mode and the total TKE decrease sharply, while it is still enhanced compared with $\varepsilon=0$. Interestingly, only a small gravity centre offset $(\varepsilon \geq 0.05)$ is required to form the convective jet and the corresponding meridional circulation for $R a=10^{8}$. The inset of figure $6(b)$ shows the distribution of the $m=0$ mode over the different co-latitudinal structures $\left(\overline{\left\langle C_{\ell, m=0}\right\rangle_{r}}\right)$, which shows that the large-scale circulation is dominated by the co-latitudinal modes with $\ell \leq 3$. With increasing $\varepsilon, \overline{\left\langle C_{\ell, m=0}\right\rangle_{r}}$ is enhanced for low ( $\ell \leq 3$ ) and high ( $\ell \geq 70$ ) wavenumbers, compared with $\varepsilon=0$.


Figure 7. (a) Mean temperature $\overline{\langle T\rangle_{\theta}}(r, \varphi)$ averaged over time and longitudinal direction shown in the meridian plane for $\varepsilon=0.05,0.4$ and 0.8. (b) Mean temperature profiles $\left(\overline{\langle T\rangle_{\theta}}(r)\right)$ at three co-latitudes for $\varphi=0, \pi / 2$ and $\pi$, respectively. The inset for $\varphi=\pi / 2$ shows the temperature gradient $\mathrm{d} \overline{\langle T\rangle_{\theta}}(r) / \mathrm{d} r$.

### 3.4. Effect of large-scale flow on heat transfer distribution

As we have shown above, the gravity centre offset influences strongly the flow structures. This raises the question of how the local temperature profiles and heat fluxes are affected. Figure $7(a)$ shows the mean temperature in the meridian plane. For $\varepsilon=0$, the temperature is uniform in the co-latitude direction. However, with increasing $\varepsilon$, the effect of the north-pole convective jet, which brings hot fluid from the inner to the outer sphere, becomes more pronounced. The effect of the jet is more visible in figure $7(b)$, which shows the temperature profiles at selected angular positions in the co-latitudinal direction. Most of the temperature drop occurs in the thermal boundary layer regardless of $\varepsilon$ and $\varphi$. However, for $\varepsilon=0.4$ and 0.8 , the positive temperature gradient ( $\mathrm{d} \overline{\langle T\rangle_{\theta}} / \mathrm{d} r$ ), observed at $\varphi=\pi / 2$ and shown in the inset, indicates some thermal stratification in the bulk region.

Next, we examine how the local heat fluxes depend on $\varepsilon$. Figure 8 shows the Nusselt number along the inner $N u_{i}(\varphi)$ and outer $N u_{o}(\varphi)$ spheres. It is observed that the local heat fluxes on the outer sphere decrease from $\varphi=0$ (north pole) to $\varphi=\pi$ (south pole), whereas the opposite trend is observed on the inner sphere. This happens because the hot convective


Figure 8. Nusselt number $N u$ on $(a)$ the inner sphere and $(b)$ the outer sphere, as functions of the co-latitudinal direction for $R a=10^{8}$ and $n_{g}=0$. The inset of $(b)$ shows $N u(\varepsilon) / N u(\varepsilon=0)$ as a function of $\varepsilon$. The error bars represent standard deviations. We note that the error bars for standard errors are smaller than the size of the symbols, which are representative for the uncertainty in the presented global Nusselt values.
jet impinges on the outer boundary layer, resulting in a thinner thermal boundary layer and larger heat flux than without the jet. On the inner sphere, the formation of the convective jet reduces temperature gradients, and therefore the heat transport, at the north pole of the inner sphere, as shown in figure $7(a)$. Surprisingly, the formation of the large-scale circulation and corresponding non-uniform heat flux distribution has a limited influence on the overall heat flux, which globally is not sensitive to the large-scale dynamics in the flow (Ahlers et al. 2009) (see inset in figure $8 b$ ), although some reduction in the heat transport is observed for $\varepsilon=0.8$ when the flow becomes stratified more thermally.

### 3.5. Effect of Ra and gravity profile

So far, we have focused on the effect of the gravity centre offset on the flow structures. In this subsection, instead, we show that convective jet and large-scale flow structures are formed for a wide range of $n_{g}$ and $R a$, and that the control parameters modify significantly the intensity of the large-wavelength structures. However, as we have seen above, the global heat flux depends only weakly on $\varepsilon$ in the considered parameter regime.

Figures $9(a-d)$ show visualizations of the convective jet for $R a=10^{7}, \varepsilon=0.4$, and different gravity profiles. The figures show that the convective jet forms with $u_{r}>0$ in the north pole region when the gravity centre shifts to the south. As expected, the jet becomes more turbulent when $n_{g}$ is decreased, as this increases the mean buoyancy power over the spherical shell. As shown previously in figure 3(a), the co-latitudinal buoyancy component $C_{g, \varphi} g \hat{\varphi}$ points to the north pole in all circumstances. However, the radial buoyancy coefficient $C_{g, r} g \hat{r}$ is strongly dependent on the co-latitude. It is well accepted that the stronger radial velocity component tends to induce more intense thermal plumes. However, the convective jet is still observed in the north pole region rather than the south pole region when $n_{g}=-2$, as shown in figure $9(d)$. This confirms that the large-scale circulation is a result of the baroclinic imbalance due to the misalignment of the gravity isopotential and isothermal lines, which is indicated by steady flows shown in figures $3(b, c)$.

The corresponding $\operatorname{Re} e_{r m s}(r)$, presented in figure $10(a)$, shows that the turbulent intensity is enhanced greatly with decreasing $n_{g}$. This reveals that the flow becomes more


Figure 9. ( $a-d$ ) Instantaneous isosurfaces of $T=0.2$ for different gravity distributions from $n_{g}=1$ to $n_{g}=-2$ (see (2.7)) for $R a=10^{7}$ and $\varepsilon=0.4$. Panel (e) shows $R a=10^{8}, n_{g}=0$ and $\varepsilon=0.4$, for comparison. The colour indicates $u_{r}$ at the isosurface. (Two movies for $R a=10^{7}, n_{g}=0$ and $\varepsilon=0.8$ are provided; see supplementary movies 1 and 2 .)
turbulent when $n_{g}$ decreases, due to the increased buoyancy power over the volume. Correspondingly, figure $10(b)$ shows that the TKE spectrum of $\overline{\left\langle C_{m}\right\rangle_{r}}$ increases for all wavenumbers. We note that $\left\langle C_{m=0}\right\rangle_{r}$, which dominates TKE and indicates the strength of the large-scale circulation, increases with decreasing $n_{g}$. With increasing $R a$ from $10^{7}$ to $10^{8}$, the normalized $R e_{r m s}(r)$ is enhanced as shown in figure $10(c)$. As shown in figure $10(d)$, the enhancement happens for all wavenumbers, especially for the high wavenumber structures. Thus for a constant $\varepsilon=0.4$, the large-scale circulation is more energetic for higher $R a$ and smaller $n_{g}$, due to the increased mean buoyancy power over the volume and the relatively stronger baroclinic torque as indicated in figures $3(b, c)$.

The generation of large-scale structures can induce highly non-uniform co-latitudinal heat flux distribution. We showed previously that the global heat transfer is relatively insensitive to $\varepsilon$ for $n_{g}=0$ and $R a=10^{8}$. Figure 11 shows that a similar trend is observed for different $n_{g}$ and $R a$. In particular, for $\varepsilon \leq 0.8$, the difference between $N u(\varepsilon>0)$ and $N u(\varepsilon=0)$ is less than $8 \%$, which indicates that the effect of the large-scale flow structures on the overall heat transport is relatively limited.

## 4. Summary and conclusions

We demonstrated using direct numerical simulations that even a small shift of the gravity centre location can introduce pronounced changes in the flow organization and local heat fluxes in spherical RB convection. In the Rayleigh number ( $R a$ ) space explored in the present study, we observe the presence of the baroclinic flow, which induces a large-scale


Figure 10. Plots of $(a, c) R e_{r m s}(r)$ as function of the gap width, and $(b, d)$ the TKE spectrum of $\overline{\left\langle C_{m}\right\rangle_{r}}$ for $\varepsilon=0.4$. (a,b) Here, $R a=10^{7}$ with different $n_{g} ;(c, d)$ here, $n_{g}=0$ with different $R a$. Figures are normalized by $R a=10^{7}$ with $n_{g}=0$. The values of buoyancy power for $R a=10^{7}$ and $n_{g}=1,0,-1,-2$, and $R a=10^{8}$ and $n_{g}=0$, are 3.3, 5.55, 12.3, 42.8 and 9.46 , respectively.


Figure 11. Plots of $N u(\varepsilon) / N u(\varepsilon=0)$ as a function of $\varepsilon$ for different $R a$ and gravity profiles $n_{g}$; see (2.7).
circulation and jet for $\varepsilon \geq 0.05$. However, for very small $\varepsilon=0.01$ and large $R a=10^{8}$, the baroclinic flow is weakened significantly due to the dominance of the strong convective flow. In turbulent flows with $\varepsilon \geq 0.05$, a strong convective jet on the northern side of the inner sphere is formed, which leads to hot/cold fluid going towards the north/south along the inner/outer sphere. The formation of the large-scale circulations has been confirmed by modal analysis. Due to the large-scale flow organization, the heat transfer at the inner and outer sphere is highly non-uniform. In particular, around the north pole, the heat flux is enhanced on the outer sphere due to the impingement of the convective jet, while the formation of the convective jet at the north pole of the inner sphere reduces temperature gradients, and therefore the local heat transport. However, surprisingly, the global heat flux is relatively insensitive to the gravity centre shift. These findings do not seem to depend much on $R a$ (over our explored range) or the particular gravity profile.

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## Appendix A. Buoyancy components of off-centre gravity

The two-step transformation of a vector GP in Cartesian coordinates xyz to spherical coordinates $\varphi \theta r\left(x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}\right)$ is shown in figure 12. A random fixed point $G\left(x_{G}, y_{G}, z_{G}\right)$ is representing the gravity centre location. The other random point, $P(r \sin \varphi \cos \theta, r \sin \varphi \sin \theta, r \cos \varphi)$, is a moving point representing a fluid parcel. Initially, both $P$ and $G$ are in Cartesian coordinates $x y z$. Using two steps, we will transfer $\boldsymbol{G P}$ to spherical coordinates $\varphi \theta r\left(x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}\right)$.
(i) Rotating $x y z$ to $x^{\prime} y^{\prime} z^{\prime}$ about $z$ by an angle $\theta$ using the right-hand side rule. The corresponding rotation matrix is

$$
R_{z}(\theta)=\left[\begin{array}{ccc}
\cos \theta & -\sin \theta & 0  \tag{A1}\\
\sin \theta & \cos \theta & 0 \\
0 & 0 & 1
\end{array}\right]
$$

(ii) Rotating $x^{\prime} y^{\prime} z^{\prime}$ to $x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}$ about $y^{\prime}$ by an angle $\varphi$ using the right-hand side rule. The corresponding rotation matrix is

$$
R_{y}(\varphi)=\left[\begin{array}{ccc}
\cos \varphi & 0 & \sin \varphi  \tag{A2}\\
0 & 1 & 0 \\
-\sin \varphi & 0 & \cos \varphi
\end{array}\right]
$$

After the above two steps we obtain that $x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}$ overlaps with $\varphi \theta r$. The transformation matrix $\mathcal{M}_{\theta \varphi}$ from $\boldsymbol{G} \boldsymbol{P}_{x y z}$ to $\boldsymbol{G} \boldsymbol{P}_{x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}}$ is defined by

$$
\begin{align*}
\mathcal{M}_{\theta \varphi}=R_{z}(\theta) \times R_{y}(\varphi) & =\left[\begin{array}{ccc}
\cos \theta & -\sin \theta & 0 \\
\sin \theta & \cos \theta & 0 \\
0 & 0 & 1
\end{array}\right]\left[\begin{array}{ccc}
\cos \varphi & 0 & \sin \varphi \\
0 & 1 & 0 \\
-\sin \varphi & 0 & \cos \varphi
\end{array}\right] \\
& =\left[\begin{array}{ccc}
\cos \varphi \cos \theta & -\sin \theta & \sin \varphi \cos \theta \\
\cos \varphi \sin \theta & \cos \theta & \sin \varphi \sin \theta \\
-\sin \varphi & 0 & \cos \varphi
\end{array}\right] . \tag{A3}
\end{align*}
$$



Figure 12. Schematic figure of the two-step transformation of a vector $\boldsymbol{G P}$ in Cartesian coordinates $x y z$ to a spherical coordinates $\varphi \theta r\left(x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}\right)$ frame. Circular arrows about the $z$ - and $y^{\prime}$-axes show rotation axes.

In the Cartesian coordinate frame, $\boldsymbol{G P}_{x y z}=\left(r \sin \varphi \cos \theta-x_{G}, r \sin \varphi \sin \theta-y_{G}, r \cos \varphi-\right.$ $z_{G}$ ). Using the above transformations, we find that $\boldsymbol{G P}$ in spherical coordinates $x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}$ is

$$
\begin{gather*}
\boldsymbol{G} \boldsymbol{P}_{x^{\prime \prime} y^{\prime \prime} z^{\prime \prime}}=\boldsymbol{G} \boldsymbol{P}_{x y z} \times \mathcal{M}_{\theta \varphi}=\left(-x_{G} \cos \varphi \cos \theta-y_{G} \cos \varphi \sin \theta+z_{G} \sin \varphi,\right. \\
\left.x_{G} \sin \theta-y_{G} \cos \theta, r-x_{G} \sin \varphi \cos \theta-y_{G} \sin \varphi \sin \theta-z_{G} \cos \varphi\right) . \tag{A4}
\end{gather*}
$$

In the dimensionless equations for the velocity $\boldsymbol{u}$, the buoyancy term is given by $g T G P / L_{G P}$, where

$$
\begin{equation*}
L_{G P}=|\boldsymbol{G P}|=\sqrt{\left(r \sin \varphi \cos \theta-x_{G}\right)^{2}+\left(r \sin \varphi \sin \theta-y_{G}\right)^{2}+\left(r \cos \varphi-z_{G}\right)^{2}} \tag{A5}
\end{equation*}
$$

The three components in the $\hat{\varphi}, \hat{\theta}$ and $\hat{r}$ directions can be decomposed as

$$
\left.\begin{array}{rl}
C_{g, \theta} & =\left(x_{G} \sin \theta-y_{G} \cos \theta\right) / L_{G P}  \tag{A6}\\
C_{g, r} & =\left(r-x_{G} \sin \varphi \cos \theta-y_{G} \sin \varphi \sin \theta-z_{G} \cos \varphi\right) / L_{G P}, \\
C_{g, \varphi} & =\left(-x_{G} \cos \varphi \cos \theta-y_{G} \cos \varphi \sin \theta+z_{G} \sin \varphi\right) / L_{G P}
\end{array}\right\}
$$

## Appendix B. Code validations

We validated extensively our code against the results presented by Gastine et al. (2015), who performed a systematic parameter study for spherical RB convection with $\operatorname{Pr}=1$ covering radius ratios $0.2 \leq \eta \leq 0.95$ for different gravity profiles in table 2 . The table shows that $N u$ and the volume-averaged Reynolds number $R e_{r m s}$ agree within $2 \%$ with the results from Gastine et al. (2015) for all cases. To ensure that the off-centre gravity is implemented correctly, we verified that the flow dynamics for the same $L_{O G}$ (see figure 1) are identical for five different gravity centre locations. The last five rows of table 2 confirm that both $N u$ and $R e_{r m s}$ obtained from these simulations are consistent.

## Appendix C. Spectral analysis in spherical coordinates

This appendix describes the spectral analysis procedure used in this study. For further reading on this topic, we refer to pp. 122-145 of MacRobert (1947). In planar configurations, Fourier transformations can be applied when performing spectral analysis.

| $R a$ | $\eta$ | $n_{g}$ | $G(\theta, r, \varphi)$ | $N_{\theta}, N_{r}, N_{\varphi}$ | $N u$ | $N u(\mathrm{ref})$ | $R_{r m s}$ | $R e_{r m s}(\mathrm{ref})$ |
| :--- | :---: | :---: | :---: | ---: | :---: | ---: | ---: | ---: |
| $7 \times 10^{6}$ | 0.3 | 1 | $O$ | $181,73,181$ | 6.42 | 6.40 | 282.2 | 287.2 |
|  | 0.3 | 0 | $O$ | $181,73,181$ | 8.19 | 8.15 | 377.5 | 377.8 |
| $3 \times 10^{7}$ | 0.3 | 1 | $O$ | $251,93,163$ | 9.39 | 9.38 | 592.0 | 595.5 |
| $5 \times 10^{6}$ | 0.35 | 1 | $O$ | $257,55,257$ | 6.80 | 6.74 | 274.8 | 274.1 |
| $3 \times 10^{8}$ | 0.35 | 1 | $O$ | $577,119,577$ | 21.47 | 21.23 | 1815.6 | 1824.8 |
| $5 \times 10^{6}$ | 0.6 | 0 | $O$ | $257,71,257$ | 11.85 | 11.68 | 437.3 | 442.7 |
| $10^{6}$ | 0.6 | -2 | $O$ | $325,71,325$ | 9.02 | 8.90 | 255.4 | 259.2 |
| $3 \times 10^{4}$ | 0.6 | -2 | $O$ | $129,49,97$ | 3.42 | 3.40 | 44.1 | 44.0 |
|  |  | -2 | $\left(0, r_{i} / 2, \pi / 2\right)$ | $129,49,97$ | 3.197 | - | 40.8 | - |
|  |  | -2 | $\left(\pi / 2, r_{i} / 2, \pi / 2\right)$ | $129,49,97$ | 3.187 | - | 40.8 | - |
|  |  | -2 | $\left(0, r_{i} / 2,0\right)$ | $129,49,97$ | 3.191 | - | 41.1 | - |
|  | -2 | $\left(\pi / 4, r_{i} / 2, \pi / 4\right)$ | $129,49,97$ | 3.182 | - | 40.9 | - |  |
|  | -2 | $\left(\pi / 7, r_{i} / 2,3 \pi / 5\right)$ | $129,49,97$ | 3.186 | - | 41.0 | - |  |

Table 2. Parameters of the spherical RB simulations with $\operatorname{Pr}=1$, which are compared to the results ( $\mathrm{Nu}(\mathrm{ref})$ and $\left.R e_{r m s}(\mathrm{ref})\right)$ of Gastine et al. (2015). Here, $G(\theta, r, \varphi)$ is the gravity centre location, and $N_{\theta}, N_{r}, N_{\varphi}$ indicate the numbers of grid points in the longitudinal, radial and co-latitudinal directions, respectively. The last five simulations for $R a=3 \times 10^{4}$ are used to verify that the coefficients $C_{g, i}$ (see (A 6)) are implemented correctly, by putting the gravity centre $G$ in five different locations with the same $L_{O G}$ (see figure 1).

However, for a spherical configuration, the basis wave functions are found by solving the Laplace equation. Analogous to the complex exponential in planar configuration, the spherical harmonics $Y_{\ell}^{m}(\theta, \varphi)$ read

$$
\begin{equation*}
Y_{\ell}^{m}(\theta, \varphi)=P_{\ell}^{m}(\varphi) \mathrm{e}^{\mathrm{i} m \theta} \tag{C1}
\end{equation*}
$$

where $\ell$ and $m(-\ell \leq m \leq \ell)$ represent the wavenumbers along a meridian and the equatorial plane, respectively. The polar angle $\varphi$ ranges from 0 to $\pi$, and $\theta$ is the azimuthal angle, in the range $0 \leq \theta \leq 2 \pi$. Also, $P_{\ell}^{m}(\varphi)$ are the associated Legendre functions. For computational purposes, the normalized associated Legendre functions in (C2) are more attractive since $P_{\ell}^{m}(\varphi)$ can overflow the computer (Swarztrauber 1979):

$$
\begin{equation*}
\bar{P}_{\ell}^{m}(\varphi)=\left[\frac{2 \ell+1}{2} \frac{(\ell-m)!}{(\ell+m)!}\right]^{1 / 2} P_{\ell}^{m}(\varphi) \tag{C2}
\end{equation*}
$$

The use of the spherical harmonics for approximating functions on a sphere is motivated by the fact that they form a complete system of orthogonal functions. Therefore, any function $f(\theta, \varphi)$ that is continuous and has continuous derivatives up to second-order may be expanded in an absolutely and uniformly convergent series

$$
\begin{equation*}
f(\theta, \varphi)=\sum_{\ell=0}^{\infty} \sum_{m=-\ell}^{\ell} \bar{P}_{\ell}^{m}(\varphi)\left(a_{\ell, m} \cos (m \theta)+b_{\ell, m} \sin (m \theta)\right), \tag{C3}
\end{equation*}
$$

where the expansion coefficients are given by

$$
\begin{align*}
& a_{\ell, m}=\frac{1}{\pi} \int_{0}^{2 \pi} \int_{0}^{\pi} f(\theta, \varphi) \bar{P}_{\ell}^{m}(\varphi) \cos (m \theta) \sin \varphi \mathrm{d} \theta \mathrm{~d} \varphi  \tag{C4}\\
& b_{\ell, m}=\frac{1}{\pi} \int_{0}^{2 \pi} \int_{0}^{\pi} f(\theta, \varphi) \bar{P}_{\ell}^{m}(\varphi) \sin (m \theta) \sin \varphi \mathrm{d} \theta \mathrm{~d} \varphi \tag{C5}
\end{align*}
$$



Figure 13. Spectral analysis of a regular large-scale circulation in a steady flow field for $R a=10^{3}$, $n_{g}=-2$ and $\varepsilon=0.4$. (a) The streamlines (curves coloured by the velocity magnitude) and vortical structures (isosurfaces of the $Q$-criterion, $Q=0.006$, coloured black). (b) The equatorial cut, meridional cut and radial surface located at $r=r_{i}+0.05 d$ of the velocity magnitude. (c) Plot of ${\left.\overline{\left\langle C_{\ell, m=0}\right.}\right\rangle_{r} \text { of the velocity field as a }}$ function of the spherical harmonic order $\ell+1$ in logarithmic scale. Also shown is the spherical harmonic basis of $\ell=0,1,2,3,4$ for $m=0$.

The spherical harmonics $Y_{\ell}^{m}(\theta, \varphi)$ are orthonormal, so

$$
\begin{equation*}
\frac{1}{2 \pi} \int_{\varphi=0}^{\pi} \int_{\theta=0}^{2 \pi} Y_{\ell}^{m} Y_{\ell^{\prime}}^{m^{\prime} *} \mathrm{~d} \Omega=\delta_{\ell \ell^{\prime}} \delta_{m m^{\prime}} \tag{C6}
\end{equation*}
$$

where $\delta_{i, j}$ is the Kronecker delta, and $\mathrm{d} \Omega=\sin \varphi \mathrm{d} \theta \mathrm{d} \varphi$ is the differential solid angle in spherical coordinates. Coefficients (C4) and (C5) are obtained by solving (C3) on a uniform spherical grid by using the SPHEREPACK library (Adams \& Swarztrauber 1999).

The power spectrum on a spherical surface $C_{\ell, m}$ and in the longitudinal direction $C_{m}$ is defined in (3.2) and (3.3), respectively. To demonstrate the formation of ( $\mathrm{m}=$ $0, \ell>0$ ) modes corresponding to the meridional large-scale circulation, we perform a simulation for $n_{g}=-2$ and $\varepsilon=0.4$ with $R a=10^{3}$ (much lower than the critical Rayleigh number when $\varepsilon=0$ ). The advantage of using this steady flow is a regular large-scale circulation generated without longitudinal dependent structures $(m \neq 0)$, which can be seen from figures $13(a, b)$. The spectrum of the kinetic energy, i.e. $\overline{\left\langle C_{\ell, m=0}\right\rangle}{ }_{r}$, is shown in figure $13(c)$. Except for the mode of ( $m=0, \ell=0$ ) representing the mean flow, apparently the large-scale circulation is represented by $(m=0, \ell>0)$ modes.

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