

NEW DATA ON CHRONOLOGY OF LANDSCAPE-PALEOClimATIC STAGES IN NORTHWESTERN RUSSIA DURING THE LATE GLACIAL AND HOLOCENE

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ABSTRACT. Two lake and bog sediments have been thoroughly studied using palynological and radiocarbon dating methods. These are the Lembolovskoye Lake section located in the southern part of the Karelian Isthmus and the Mshinskoye bog section located in the southwestern part of the Leningrad province. The data obtained allow us to reconstruct the main features of the vegetation cover evolution, the chronology for the appearance and a real distribution of the main arboreal species from the south of the Leningrad province to the north, and to construct curves of the paleoclimate parameter changes for the area under study. Thirty-six ¹⁴C dates were obtained for the Lembolovskoye Lake section (7 m thick). According to those dates, the organic gyttja formation in the lake began 9870 ± 170 BP. Spore-pollen spectra with high percentage of herbs, arborescent, and shrub-birch dated back to the Younger Dryas were found out in clay layers at a depth of 6.5 m. An appearance of spruce and alder pollen is dated at 6860 ± 120 and 7510 ± 150 BP, respectively. The maximal percentage of broad-leaved species falls on the first half of the Atlantic (AT-1). Thirty-two dates were obtained for the Mshinskoye bog section (6 m thick): from 60 ± 70 to 9520 ± 170 BP (the last date fixed the beginning of peat formation); 12 palinotones from the Preboreal to the Subatlantic were recognized there. The spruce and alder pollen began to appear 7520 ± 110 and 7670 ± 130 BP, respectively. The maximal amount of broad-leaved species is observed at 4690 ± 80 BP. The detailed reconstruction of changes in vegetation communities during the Late Glacial and Holocene was correlated with paleoclimatic characteristics, which have been reconstructed for the sections under study by using the information from a statistical method of spore-pollen data processing.

INTRODUCTION

During 1993–1999 we carried out systematically paleobotanical and geochronological investigations of continuously formed bog and lake sediments in northwestern Russia. The study was undertaken to reconstruct the stages of changing vegetation and paleoclimates during the Late Glacial and Holocene. The results of the study for five reference sections in the Leningrad and Novgorod provinces and two sections in Karelia, for which 320 radiocarbon dates have been obtained, were published earlier (Arslanov et al. 1999; Elina et al. 1996). Later, in 1998–1999, using the unified procedure we studied two more sections in the Leningrad province: the Lembolovskoye Lake and the Mshinskoye bog. The object was to obtain additional paleobotanical and geochronological data for more precise definition and detailing of the stages mentioned above.

METHODS

We analyzed palynologically every 10-cm layer of continuously formed bog and lake sediments along the whole section and the great majority of the layers were radiocarbon dated at the Geochronological Laboratory of the Geographical Research Institute, St. Petersburg University. The pre-treatment of samples (peat and gyttja), the benzene synthesis from the pretreated samples and the ¹⁴C activity measurements in the synthesized samples were carried out by the procedures published earlier (Arslanov et al. 1993). For palynological analysis the peat and gyttja samples were boiled in 10% NaOH for 5 min, then rinsed by distilled water and the residue was separated by centrifugation and then analyzed. The pretreatment of mineralized samples or samples with a small amount of organic matter was carried out by the procedure described in Arslanov et al. (1999), which allows us to separate organic and mineral fractions by using a heavy liquid (PD-6 or KK-2.6). The percentage

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of arboreal pollen and shrub groups (AP) and herbs (NAP), individual arboreal and herb species and also spores was calculated taking the total pollen AP and NAP sum as 100%. Using the TILIA-GRAPH computer program the chrono-palynological diagrams were constructed.

RESULTS

Lembolovskoye Lake Section (60°21'45"N, 30°18'80"E)

Lembolovskoye Lake is located on the eastern slope of the Karelian Highland. It is shallow (average depth 1.5–2.0 m), stretched out from the north to the south (9 km long, 2 km wide at the maximum), divided into two parts by a narrow isthmus. The water level in the lake stays at 42.9 m above sea level (asl).

In the spring of 1998 a borehole was drilled through ice, which cropped out 760-cm-thick sediments of the lake. Dense clays occur at the basement of the lake sediments' core (110 cm thick). They give place to greenish-brown organogenic sapropel (650 cm thick).

Seventy-six samples were studied by palynological method in every 10-cm layer. All the samples contained a great quantity of well-preserved pollen and spores except for the clay layer. Spores of pre-Quaternary (probably redeposited) vegetation were found out in clayey sediments; they represent 5–7% of the pollen and spores sum in situ.

Eleven local pollen assemblage zones (LPAZ), which correlated well with chronozones of the Blytt-Sernander scheme, were marked out on the chrono-palynological diagram of the lake sediments (Figure 1, Table 1).

LEM-1, 750–680 cm, varved clay, Younger Dryas. Pollen of arboreal birch (*Betula sect. Albae*) dominate here and have a small maximum (about 25%). The portion of shrub (*Betula sect. Fruticosae*) and dwarf (*Betula nana*) birch is much smaller (about 10%). The amount of *Pinus* pollen is about 10–15%. The herb pollen (NAP) amounts to 25–35% with predomination of such herbs as *Artemisia*, *Poaceae*, *Chenopodiaceae*, and *Cyperaceae*. *Selaginella selaginoides* spores are also present.

According to the pollen data, birch and pine forest tundra vegetation with grasses were widespread during the Younger Dryas.

LEM-2, 680–650 cm, organic gyttja, Preboreal. *Betula sect. Albae* predominates the pollen spectrum (35–55%); the amount of the *Betula sect. Fruticosae* and *Betula nana* pollen reduced to 2–5%. The percentage of *Pinus* pollen did not exceed 2% and of NAP –20%.

The data obtained shows that the area occupied by tundra vegetation with shrub and dwarf birch, willow, *Artemisia* and *Chenopodiaceae* was diminished. Arboreal birch became the main forest-forming species.

LEM-3a, 650–595 cm, organic gyttja, Early Boreal. The amount of arboreal species (AP) increased sharply (to 97%) and that of NAP fall to 3%. Pollen of *Pinus* predominated (55–70%), and *Betula sect. Albae* (15–30%) came next.

At the first half of the Boreal, the role of xerophytes in the landscape is reduced sharply. Pine and birch-pine forest predominated.

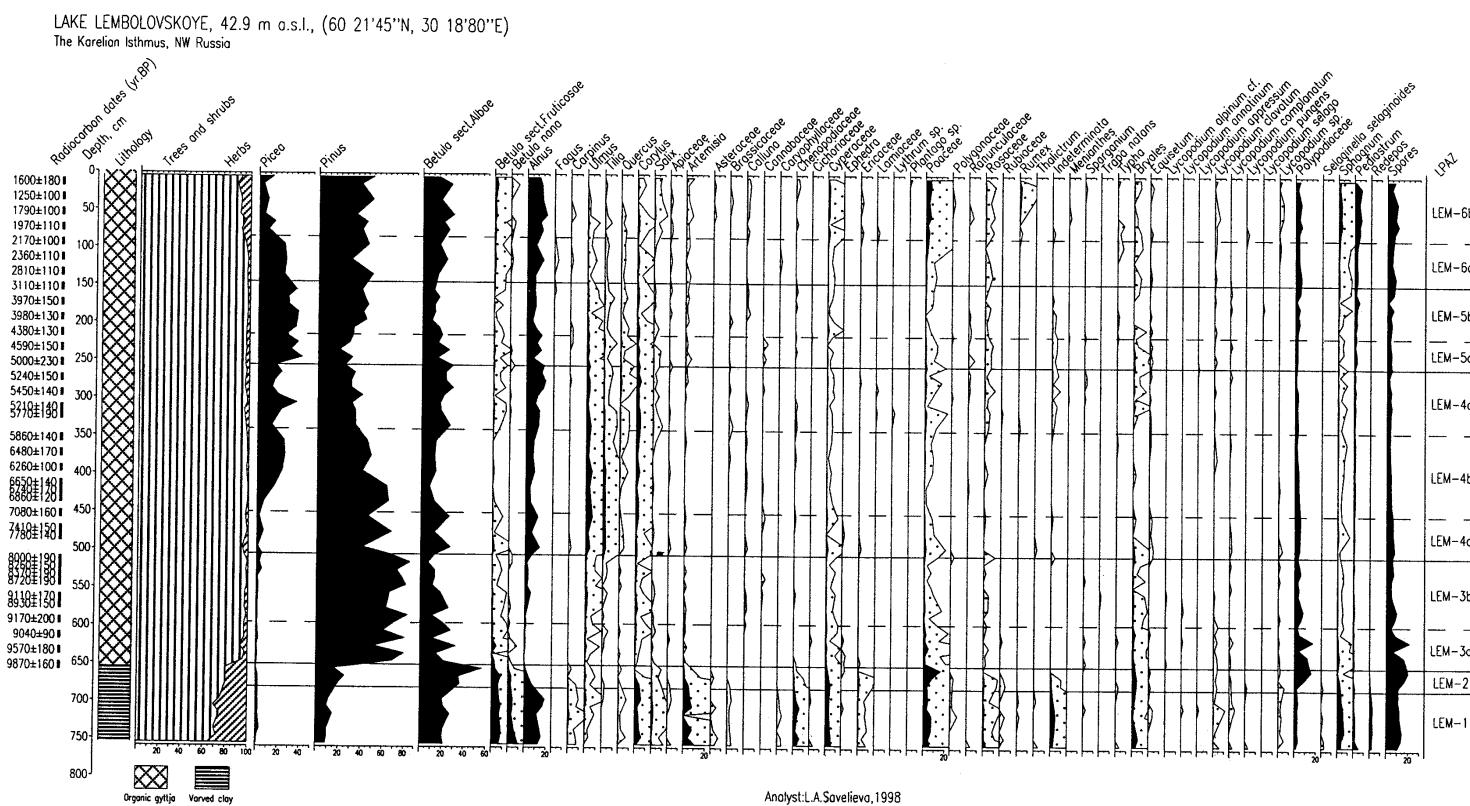


Figure 1 Chrono-palynological diagram of the Lembolovskoye Lake sediments

Table 1 ^{14}C dates of the Lembolovskoye Lake sediments

Depth (cm)	Lab code	^{14}C age (BP)	Calibrated age ^a AD/BC (1 σ)
LU-			
30–40	4112	1250 ± 100	685–885 AD
50–60	4111	1790 ± 100	125–380 AD
70–80	4110	1970 ± 110	115–210 AD
90–100	4109	2170 ± 100	360–95 BC
110–120	4108	2360 ± 110	760–235 BC
130–140	4107	2810 ± 110	1125–835 BC
150–160	4106	3110 ± 110	1515–1215 BC
170–180	4105	3970 ± 150	2840–2205 BC
190–200	4104	3980 ± 130	2835–2235 BC
210–220	4103	4380 ± 130	3330–2880 BC
230–240	4102	4590 ± 150	3520–3095 BC
250–260	4101	5000 ± 230	4215–3520 BC
270–280	4100	5240 ± 150	4315–3820 BC
290–300	4099	5450 ± 140	4450–4050 BC
310–320	4098	5210 ± 140	4225–3805 BC
320–330	4131	5770 ± 190	4835–4365 BC
350–360	4096	5860 ± 140	4905–4545 BC
380–390	4128	6270 ± 160	5465–5005 BC
390–400	4094	6260 ± 100	5320–5060 BC
400–410	4127	6650 ± 190	5730–5375 BC
410–420	4093	6650 ± 140	5710–5475 BC
420–430	4126	6740 ± 170	5775–5485 BC
430–440	4092	6860 ± 120	5865–5635 BC
450–460	4091	7080 ± 160	6155–5785 BC
460–470	4124	7310 ± 160	6375–6015 BC
470–480	4090	7410 ± 150	6405–6095 BC
490–500	4089	7510 ± 150	6495–6115 BC
510–520	4088	8000 ± 190	7300–6650 BC
520–530	4121	8260 ± 150	7515–7080 BC
530–540	4087	8370 ± 190	7585–7085 BC
540–550	4120	8720 ± 190	8195–7585 BC
560–570	4119	9110 ± 170	8600–7975 BC
570–580	4083	8930 ± 150	8270–7840 BC
590–600	4081	9170 ± 200	8720–8025 BC
610–620	4079	9040 ± 90	8411–7973 BC
620–630	4115	9830 ± 150	9605–8940 BC
630–640	4116	9570 ± 180	9215–8655 BC
650–660	4078	9870 ± 160	9685–9185 BC

^aRadiocarbon ages are corrected to calibrated ages according to the Groningen Calibration Program version CAL25.

LEM-3b, 595–505 cm, organic gyttja, Late Boreal. *Pinus* pollen still dominated (60–80%), *Betula sect. Albae* was a little less than 30%. *Ulmus-* and *Corylus-* pollen were present constantly (2–3%) and *Alnus* appeared (2–5%).

The data suggests wide spreading pine forests with an admixture of birch and alder. The thermophilic flora occurred in the forests, elm and hazel in the first turn but their amount is negligible.

LEM-4a, 505–450 cm, organic gyttja, Early Atlantic. The amount of *Picea* pollen decreased sharply (to 45%) and that of *Alnus-*, *Ulmus-*, *Tilia-* and *Quercus*-pollen increased.

During this period pine and birch remained, as before, the main forest-forming species. The broad-leaved species (elm and lime), hazel and alder have increased in importance.

LEM-4b, 450–340 cm, organic gyttja, Middle Atlantic. A rise in the *Picea* pollen curve (to 25%) is observed. The amount of *Pinus*-pollen was 40–63% and that of *Alnus*-pollen was about 15%. The percentage of *Alnus* pollen diminished and that of the broad-leaved species varied from 4 to 8%.

Pollen data shows in the middle of the Atlantic an intensive expansion of spruce forests; spruce pine and spruce-birch-pine, with an admixture of broad-leaved species, hazel and alder began to be formed.

LEM-4c, 340–255 cm, organic gyttja, Late Atlantic. *Pinus*- and *Betula sect. Albae* pollen curves are similar. The amount of *Picea*-pollen was changed from 14 to 36%, that of *Alnus* from 8 to 19% and the share of broad-leaved species varied from 3 to 7%.

Pollen data demonstrate that towards the end of the Atlantic, spruce lost its important role in the composition of the forest.

LEM-5a, 255–215 cm organic gyttja, Early Subboreal. Pollen of *Pinus* (32–45%) and *Picea* (29–37%) dominated in the spectra. The percentage of *Betula sect. Albae* and *Alnus* decreased noticeably: to 9–10% and 5–8%, respectively. Pollen of *Tilia* and *Quercus* occurred sporadically.

The beginning of the Subboreal was characterized by new spruce expansion: spruce formed great forest tracts (spruce and birch-pine-spruce forests). Thermophilic flora was present only as a negligible admixture.

LEM-5b, 215–145 cm, organic gyttja, Late Subboreal. The maximum of *Picea*-pollen is fixed (43%). Other components of the pollen spectra (such as *Pinus*, *Betula sect. Albae*, *Alnus*) remained in the same relationship as in LPAZ of LEM-5a. *Fagus*-pollen appeared among the broad-leaved species but its total amount did not exceed 2%.

Thus, spruce remained the main forest-forming species but the role of pine and birch increased gradually; the thermophilic flora elements vanished.

LEM-6a, 145–85 cm, organic gyttja, Early Subatlantic. The percentage of *Picea*-pollen decreased to 25% and that of *Betula sect. Albae* and *Alnus* increased moderately. The amount of the broad-leaved species pollen (*Carpinus*, *Fagus*, *Ulmus*, *Tilia*, *Quercus*) occurred sporadically.

The pollen data indicate that at the beginning of the Subatlantic spruce forests began to diminish gradually. Pine and birch forests with an admixture of spruce were formed.

LEM-6b, 85–00 cm, organic gyttja, Late Subatlantic. Picea-pollen were dominant (43–60%); the amount of *Betula sect. Albae* pollen was 15–30% and that of *Alnus* was 12–15%. The percentage of *Picea* pollen did not exceed 15%. The amount of NAP pollen increased to 13%.

During the Late Subatlantic pine forests with birch were spread; alder grew on the over-moistened areas. At the same time herbs (mostly, *Gramineae*, sedge, sorrel) and shrub were increasing in importance.

Thirty-six ^{14}C dates have been obtained for this section at the Geochronological Laboratory of the Geographical Research Institute, St Petersburg University (Table 1).

Mshinskoye Bog Section (59°52'00"N, 29°55'00"E)

Mshinskoye bog is located at the south of the Leningrad province. It stretches 30 km from south to north and 10–12 km from west to east. The lake occupies its central part; its southern stretch is called Vyalie Lake, the northern one—Strechno Lake. The lake length from the south to the north is 16 km; its surface is about 36 km². In October 1997, a 600-cm-deep borehole was drilled. It cropped out a thickness of peat (590 cm), underlying sand (5 cm), and blue varved clay (10 cm) layers.

Sixty samples, collected every 10 cm, have been studied by palynological methods. The chrono-palynological diagram allows marking out 11 local pollen assemblage zones (LPAZ) (Figure 2, Table 2). It has been found that the peat formation began in the Late Preboreal.

MSH-1a, 605–575 cm, varved clay, sand, peat, Early Preboreal (PB-1). The zone is characterized by predomination of pollen of *Betula sect. Albae* (to 65%). *Betula sect. Fruticosae* is 5–20% and *Betula nana* –2–5%. *Salix* pollen is present (5–7%). A relatively high pollen percentage of herbs (NAP 15–38%) is of a great importance for the zone, *Cyperacea* and *Poacea* are dominant among them.

According to the spore-pollen data, sparse birch forests with an admixture of shrub and dwarf birch, and also willow were dominant in the Early Preboreal. Herb groups from *Cyperacea* and *Poacea* grew on the open area.

MSH-1b, 575–520 cm, peat, Late Preboreal (PB-2). A reduction of herbs (to 20%) and an increase of *Pinus*-pollen (15–20%) is characteristic for the zone. *Ulmus*-pollen occurred. Other components of the spore-pollen spectra remained in the same relationship. In the Late Preboreal pine began to be of importance in pine-birch forests; herbs were spread as earlier.

MSH-2, 520–450 cm, peat, Boreal (BO). A sharp increasing of *Betula sect. Albae*-pollen (60–70%) is observed in the zone. The amount of *Pinus*-pollen changes from 7 to 25%. The percentage of NAP did not exceed 10%. An abrupt rise in the *Alnus*-curve is observed to the end of the zone. Pollen of *Tilia*, *Quercus*, *Corylus* and also *Picea* appeared.

The data shows that birch forests with an admixture of pine were widespread on the territory under study during the Boreal. The open area occupied by herb groupings from *Cyperacea*, *Poacea*, *Artemisia*, *Chenopodiacea* diminished abruptly. Spruce and thermophilic flora occurred in the forest composition.

MSH-3a, 450–370 cm, peat, Early Atlantic (AT-1). A reduction of the amount of *Betula sect. Albae* to 30% and *Pinus* pollen to 10% and also an increase in that of *Alnus* (to 25%), *Picea* (to 12%), *Ulmus* (to 8%), and *Corylus* (to 10%) are the key features for this zone. The amount of NAP is very low.

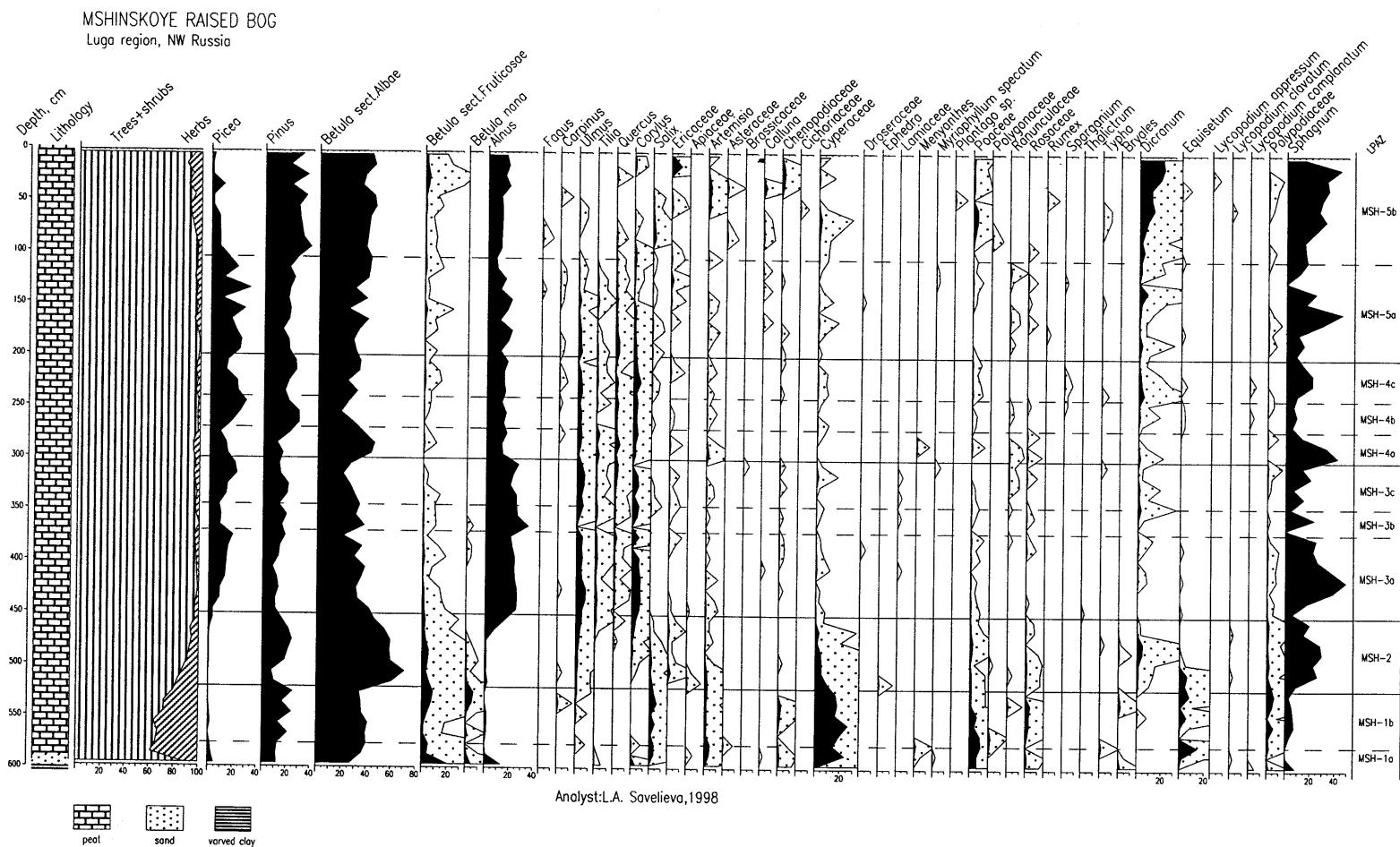


Figure 2 Chrono-palynological diagram of the Mshinskoye raised bog sediments

Spruce-pine-birch forests with an admixture of broad-leaved trees (elm, lime, oak), hazel and alder predominated on the territory during the Early Atlantic.

MSH-3b, 370–345 cm, peat, Middle Atlantic (AT-2). There is a maximum of 35% on the *Alnus* pollen curve. Also the *Picea* pollen curve has a small maximum of about 20%. The percentage of *Ulmus*, *Tilia*, *Quercus* and *Corylus* pollen decreased sharply. The amount of *Betula sect. Albae*- and *Pinus*-pollen varied from 25 to 40% and from 15 to 20%, respectively.

In the Middle Atlantic the share of thermophilic flora reduced; arboreal birch and pine remained the main forest-forming species.

MSH-3c, 345–305 cm, peat, Late Atlantic (AT-3). A maximum pollen amount of broad-leaved species in total is characterized for this zone: *Ulmus* (9%), *Tilia* (1–2%), *Quercus* (2–3%). (It is the maximum amount for the whole section). A small maximum of *Picea*-pollen (20%) is observed to the end of the zone.

Against the background of a domination of spruce-pine-birch forests, the favorable climatic conditions in the Late Atlantic promoted the wide expansion of thermophilic flora.

MSH-4a, 305–270 cm, peat, Early Subboreal (SB-1). The amount of *Betula sect. Albae*-pollen increased sharply (to 45%), that of *Pinus* is 8–12%, *Picea* 7–12%, *Alnus* 5–17%. A share of *Ulmus* reduced 2–3% and that of *Tilia* and *Quercus* increased 1–2%.

During the Early Subboreal, birch was the main forest-forming species. Birch forests with pine and spruce were developed on the area under study. The broad-leaved trees presented also a small admixture.

MSH-4b, 270–240 cm, peat, Middle Subboreal (SB-2). A maximum of *Picea*-pollen (30%) for the whole section is fixed in this zone. The percentage of *Pinus* increased to 20–25% and *Betula sect. Albae* decreased to 15–20%. Among the broad-leaved species pollen of *Carpinus* occurred.

In the Middle Subboreal, spruce achieved the maximum expansion and became the forest-forming species.

MSH-4c, 240–200 cm, peat, Late Subboreal (SB-3). The amount of pollen from the broad-leaved species reduced gradually. The percentage of *Betula sect. Albae* was 20–35%, *Pinus* 20%, and *Picea* 10–15%.

Birch, pine and spruce remained the forest-forming arboreal species. Thermophilic flora elements vanished from the forest composition.

MSH-5a, 200–195 cm, peat, Early Subatlantic. The pollen curves for the principal components of spore-pollen spectra in this zone are similar to that of LPAZ LEM-6a, what is represented in the diagram of the Lembolovskoye Lake section (Figure 1).

MSH-5b, 105–00 cm, peat, Late Subatlantic. In this zone the pollen curves for the principal components of spore-pollen spectra in this zone are similar to that of LPAZ LEM-6b (Figure 1). We obtained 32 ¹⁴C dates the Mshinskoye bog section (Table 2).

DISCUSSION

Correlating chrono-palynological diagrams of the sections in Northwestern Russia: Lambolovskoye Lake and Mshinskoye Bog with the sections studied earlier (sections: Nikolsko-Lutinskoye Bog,

Table 2 ^{14}C dates of Mshinskoye raised bog sediments

Depth (cm)	Lab code LU-	$\delta^{14}\text{C}$ (‰) or ^{14}C age (BP)	Calibrated age (AD/BC)
30–40	4053	1087 ± 8.0	
40–50	4054	60 ± 70	
50–60	4055	70 ± 80	
100–110	4063	1380 ± 100	560–775 AD
110–120	4041	1810 ± 240	45 BC–530 AD
130–140	4040	1870 ± 90	29–313 AD
150–160	4039	2120 ± 80	351–3 BC
160–170	4084	2500 ± 90	787–519 BC
170–180	4038	2770 ± 100	1020–810 BC
190–200	4037	2740 ± 100	1000–805 BC
210–220	4036	3030 ± 100	1405–1130 BC
230–240	4035	3400 ± 90	1871–1531 BC
240–250	4076	3800 ± 100	2455–2045 BC
250–260	4034	3990 ± 80	2621–2349 BC
290–300	4032	3790 ± 80	2397–2047 BC
310–320	4031	4290 ± 100	3085–2700 BC
330–340	4030	4620 ± 100	3625–3105 BC
350–360	4029	4690 ± 80	3627–3371 BC
370–380	4028	5390 ± 100	4335–4050 BC
390–400	4027	5570 ± 100	4520–4260 BC
410–420	4026	6130 ± 150	5260–4855 BC
420–430	4077	6020 ± 90	5035–4799 BC
440–450	4080	7150 ± 120	6200–5845 BC
450–460	4024	7520 ± 110	6445–6245 BC
460–470	4082	7670 ± 130	6650–6390 BC
470–480	4023	8160 ± 150	7470–6835 BC
490–500	4022	8890 ± 180	8260–7765 BC
500–510	4018	8920 ± 330	8410–7610 BC
510–520	4017	9740 ± 190	9385–8790 BC
530–540	4016	9700 ± 140	9275–8815 BC
540–550	4015	9700 ± 160	9280–8805 BC
550–560	4014	9450 ± 70	9093–8611 BC
560–570	4013	9450 ± 60	9089–8629 BC
570–580	4012	9950 ± 160	9735–9250 BC
580–590	4011	9520 ± 170	9165–8635 BC

Lammin-Suo Bog, Vishnevskoye Lake, Sakkala Bog, Suo Bog, see Figures 1–5 and Tables 1–5 in Arslanov et al. 1999; Chernova et al. 1997), we revealed characteristic peculiarities of the spore-pollen spectra for the Younger Dryas and Holocene. A presence of herbs (from 20% to 40% in the total composition) is characteristic for the spore-pollen spectra during the Younger Dryas; the pollen of *Artemisia*, *Chenopodiaceae*, *Poaceae* and *Cyperaceae* are dominant. As for arboreal species, *Betula* predominates there: *Betula* sect. *Albae* 15–25%, *Betula* sect. *Fruticosae* 10–40%, and *Betula nana* 5–20%. The percentage of *Pinus* is 10–20%. The pollen of *Alnus* is present in all spore-pollen diagrams relating to the Younger Dryas and its percentage reaches sometimes 15–20% (see Lembo-

lovskoye Lake). The broad-leaved species (mainly *Carpinus*, *Ulmus* and *Corylus*) occur only sporadically.

At the boundary of the Younger Dryas and Preboreal, the share of *Betula sect. Albae* pollen increased noticeably (to 50–60%) and that of *Betula sect. Fruticosae* and *Betula nana* decreased to 10–15% and 5–10%, respectively. The percentage of herbs pollen diminished to 20%.

In the Preboreal, *Betula sect. Albae* became the dominant species and reached even 70% (see Mshinskoye bog). The percentage of pollen varied between 15 and 45%. *Alnus* and *Picea* pollen were fixed in all diagrams of the section under study. *Poaceae* and *Cyperaceae* are in the herb pollen group.

An increase of the amount of *Pinus* pollen and a decrease of that of herbs served the main criterion for establishing the boundary between the Preboreal and Boreal.

In the Boreal, the percentage of *Pinus* pollen in lake sediments achieved 65% and in the bog it did not exceed 30%; that of *Betula sect. Albae* varied from 15 to 60%. Pollen of *Alnus*, broad-leaved species (*Ulmus*, *Tilia*, *Quercus*) and *Corylus* are constantly present in the Holocene spore-pollen spectra since the end of the Boreal.

The Boreal/Atlantic boundary is defined by the sharp rise of *Alnus*- (to 15–29%), broad-leaved species- and *Corylus*- curves and also by the occurrence of *Picea*-pollen (approximately 2–3%).

In the Atlantic spore-pollen spectra *Betula sect. Albae*, *Alnus* and *Pinus* predominate but their relationship varies strongly. The maximum percentage (in total) of broad-leaved species- and *Corylus*-pollen is characteristic for the whole Holocene. At the middle of Atlantic the amount of *Picea*-pollen was about 10% but to the end of that achieved 20–25% (Lembolovskoye Lake, Mshinskoye bog). The percentage of herbs did not exceed 5%.

The Atlantic/Subboreal boundary is drowned by increasing the amount of *Picea*-pollen and decreasing that of *Alnus* and broad-leaved species (in total).

In the Subboreal, *Pinus* (30–60%) and *Betula sect. Albae* (10–40%) predominate in spore-pollen spectra; along with this, an appreciable amount of *Picea*-pollen (varied between 15 and 40%) is observed. At the same time a sporadic presence of *Fagus* and *Carpinus* and the maximum percentage of *Quercus* are often fixed.

A rise of *Pinus*- and herb pollen curves marks the Subboreal/Subatlantic boundary. At the first half of the Subatlantic an appreciable amount of *Picea*-pollen (to 25–30%) is characteristic for the spore-pollen spectra. An increase in the pollen percentage of the *Pinus*- or *Betula sect. Albae* is observed. The share of herbs increases often 2–3 times in comparison with the Subboreal and achieves 10–15% with predomination of *Poacea*, *Cyperacea*, *Ericacea* and motley grass.

Comparing curves of the total amount of arboreal species (AP), herbs (NAP), spores and also of individual arboreal species for the sections: Lembolovskoye Lake and Mshinskoye bog and others studied earlier, we revealed general regularities of vegetation development during the Late Pleistocene and Holocene. In pollen zone SA-3, the Preboreal and the Younger Dryas stage the reduction in pollen percentage of arboreal species and increase of that of herbs and spores are observed for all sections under study with the maximal amount of herb pollen in the Younger Dryas stage. *Picea*-curves of all the sections have a maximum in the Subboreal, although the percentage of *Picea*-pollen is high enough at the second half of the Atlantic and at the first half of the Subatlantic. Only towards the end of the Subatlantic its amount decreases sharply. An appearance of *Picea*-pollen in the sedi-

Table 3 Dynamics of the indicator arboreal species and alder during the Holocene in the Leningrad province

Section names	Picea		Alnus	
	EB ^a >2%	Abrupt reduction	EB>2%	
Lembolovskoye lake	Between 6860 ± 120 7080 ± 160	1970 ± 110	Between 7510 ± 150 8000 ± 190	
Mshinskoye bog	7520 ± 110	<1380 ± 110	7670 ± 130	

Section names	Broad-leaved species							
	Ulmus		Tilia		Querens		Corylus	
	EB	Abrupt reduction	EB	Abrupt reduction	EB	Abrupt reduction	EB	Abrupt reduction
Lembolovskoye lake	Between 7410 ± 150 7510 ± 150	3970 ± 150	7510 ± 150	5860 ± 140 5860 ± 140	5450 ± 140	4380 ± 130	7510 ± 150	About 3980 ± 130
Mshinskoye bog	8160 ± 150	About 1870 ± 90	7520 ± 110	About 3990 ± 90	6020 ± 90	Between 1870 ± 90 2120 ± 80	7670 ± 130	Between 1870 ± 90 2120 ± 80

^aThe empirical boundary (EB) is a level after which pollen of particular species occur constantly (Neustadt 1965)

ments of the Lembolovskoye Lake section is dated at 6860 ± 120 BP and in those of the Mshinskoye bog one at 7520 ± 110 BP (Table 3).

Curves for pollen content of *Pinus* and *Betula* are continuous for all sections throughout the Younger Dryas and Holocene. The high content of *Pinus*-pollen in the Lembolovskoye Lake sediments (to 80% in the Preboreal) attracts the special attention. The amount of *Pinus* pollen in the bog sediments of all sections under study is much less than in the lake ones; only in the Subatlantic its content increases to 30–40%. *Betula*-curve has maximum for the Lembolovskoye and Vishnevskoye Lake sediments in the Preboreal and for the bog sediments in the Younger Dryas and Preboreal (Nikolsko-Lutinskoye Bog) and in the Preboreal and Boreal (Mshinskoye Bog, Shirinsky mokh, Lammin-Suo; Figures 1–3 in Arslanov et al. 1999). *Alnus*-curve becomes continuous after 7670 ± 130 BP for the Mshinskoye Bog and 7510 ± 150 BP for the Lembolovskoye Lake sections.

Pollen of thermophilic plants occur much earlier in the sediments of the Mshinskoye Bog section, which is located 180 km to the south of the Lembolovskoye Lake section. The data are in agreement with the regularities stated earlier for appearance of *Pinus*-, *Alnus*- and broad-leaved species-pollen in the sections located from the south towards the north (see Figures 1–5 and Table 6 in Arslanov et al. 1999). A share of broad-leaved species during the Atlantic is noticeable more (to 12%) in the Mshinskoye Bog sediments than in the Lembolovskoye Lake one located more to the north.

We used statistical information method (Klimanov et al. 1976) to reconstruct the quantitative characteristics of the Late Glacial and Holocene climates. The method is based on the statistical relationship between recent spore-pollen spectra and recent climatic conditions. The reconstruction was made on the basis of the palynological and geochronological data obtained for the Lembolovskoye Lake and the Mshinskoye bog sections. The average statistical error in determining the mean July temperature and the annual one is ± 0.6 °C, and that for the average annual precipitation is ± 25 mm.

A great amount of radiocarbon dates enabled to attach the paleoclimatic data to the time scale but the paleoclimatic curves is not very detailed because of scarce density of palynological sampling. As a result, we construct a generalized paleoclimatic curves on the base of two sections. The data obtained are illustrated in Figure 3. and demonstrate variations of the average temperatures of July, January, the annual mean and the average annual precipitation from recent values. The changes of paleotemperatures (Δt^0_{July} , $\Delta t^0_{\text{January}}$, $\Delta t^0_{\text{annual}}$) and of annual precipitation (ΔP) are represented as variations from recent temperature values. Recent climatic parameters (Climatic Atlas, 1990) for the sections under study were, on the average, as follows: mean July temperatures around 17–18 °C, mean January temperatures 8–9 °C, the average annual temperature 4–5 °C and on average precipitation of 600 mm per year.

We will not discuss here the quantitative characteristics of Δt^0_a because they are shown in Figure 3. We shall consider the time of extremes of climatic deterioration and amelioration.

The data obtained shows evidence that the Younger Dryas was the coldest period, after which the rhythmic amelioration began. Four progressively warm periods were marked out during the Preboreal. A similar picture is observed for the four deteriorations that separated these warm periods. The mean January and the annual temperatures, especially illustrate it. In the Boreal two amelioration are reconstructed and the last one is dated at 8500 BP as well as in other areas of Northern Eurasia. A noticeable cooling was recorded at the boundary of the Boreal and Atlantic. During the Atlantic several ameliorations were reconstructed and the maximum one—the Holocene maximum—was recorded at about 5500 BP; the next one dated at 6000 was not as warm. At the boundary of the AT and SB we recorded a climatic deterioration dated at around 4500 BP as it was in other regions of

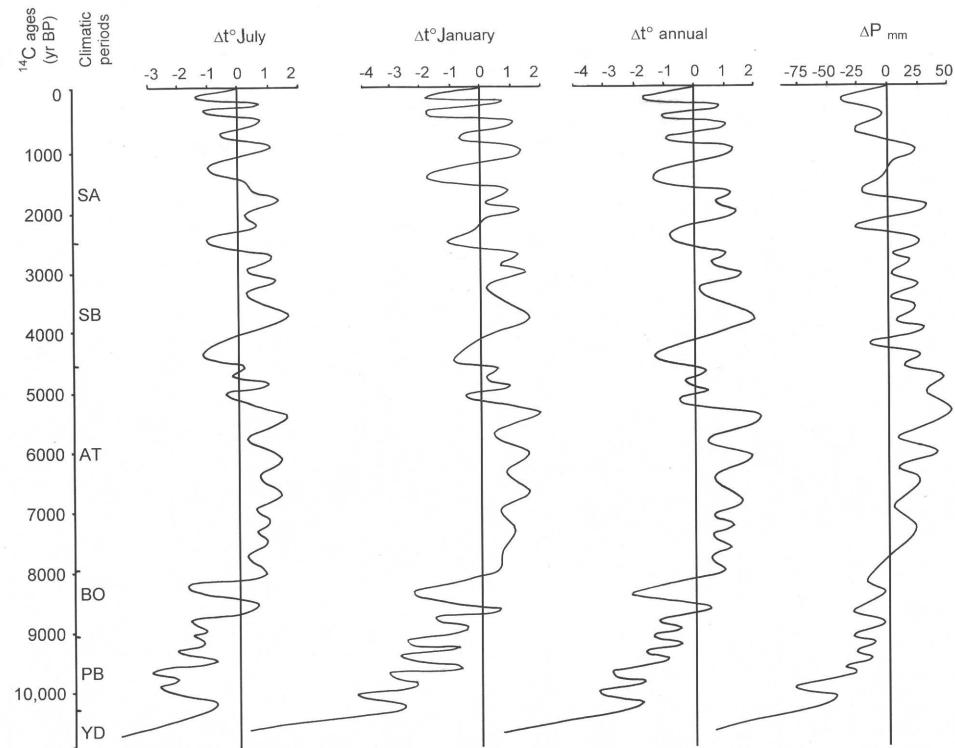


Figure 3 Deviations of the mean July, January, and annual temperatures, as well as annual precipitation from the recent temperature values during the Late Glacial and Holocene in Lenigrad province reconstructed on the bases of chrono-palynological data of the Lembolovskoye Lake and Mshinskoye raised bog sediments (total curve).

Northern Eurasia. In the Subboreal three climatic deteriorations were reconstructed; the first (at around 4000 BP) was the maximum as contrasted to other regions, where the maximum SB warming were observed at around 3500 BP. It could be connected with local characteristics of climate changes or a sparse density of sampling for spore-pollen analysis, i.e. the maximum warming at 3500 BP was missed. At the beginning of the Subatlantic (at around 2500 BP) a climatic deterioration was recorded as in other regions. The trend towards the systematic climatic deterioration occurred during the SA: every next cooling was colder than the previous one. This is well traced by changing the mean July temperatures. The last cooling before present was the coldest: it was the Little Ice Age, with a maximum in the 17th century. As for the relationship between temperatures and precipitation, it can be said that climatic ameliorations were followed by increase of precipitation and vice versa, the deteriorations were followed by a decrease in it.

Thus, a trend towards warming was marked out since the Late Glacial to the Holocene optimum, and towards cooling since the Holocene optimum to the present. The amplitude of changes in the mean January temperatures was greater than that of mean July. In the sections under study, the detailed climatic changes during the Preboreal and Boreal were reconstructed, as distinguished from other regions studied earlier. The ^{14}C data show that wide temperature changes during the Late Glacial and Holocene in the Karelian Isthmus and southwestern Leningrad province were the same as in other regions of Northern Eurasia.

CONCLUSION

A reference chrono-palynological diagram has been constructed on the base of the detailed conjugated palynological and geochronological investigations of the Lembolovskoye Lake and Mshinskoye Bog sections. It enabled us to mark out pollen assemblage zones (PAZ), which were correlated with climatic periods and phases of the Blytt-Sernander scheme since the Younger Dryas stage to the Subatlantic. They were also correlated with the stages of paleoclimatic changes stated on the base of processing the chrono-palynological data by the information-statistic method. The time of appearance and disappearance of the indicator species of trees and alder during the Holocene and also the time of their migration from south towards the north was set up. Comparison of the data obtained now and that obtained earlier for six sections located in the Leningrad and Novgorod provinces shows a similarity of the spore-pollen spectra for arboreal species. The percentage of *Pinus* is found to be noticeably higher in lake sediments than in bog ones. A ^{14}C chronology for the stages of vegetation development and paleoclimate changes during the Younger Dryas and Holocene has been set up. Younger Dryas/Preboreal transition in Northwestern Russia is characterized by transformation of birch and pine forest-tundra vegetation involving shrub and dwarf birch, willow, xerophytic herbs, grasses to sparse birch forests with admixture of shrub and dwarf birch and willow also.

According to paleoclimatic reconstruction the mean annual temperature during the coldest time of Younger Dryas, about 10,500–10,700 BP, was approximately 6 °C lower than annual recent temperature (see Figure 3 in this article and Figure 6 in Arslanov et al. 1999). At the beginning of Holocene mean annual temperature was about 2 °C lower than that at present.

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