

COMPREHENSIVE RADIOCARBON ANALYSIS OF BENZENE POLYCARBOXYLIC ACIDS (BPCAs) DERIVED FROM PYROGENIC CARBON IN ENVIRONMENTAL SAMPLES

U M Hanke^{1*} • L Wacker² • N Haghipour³ • M W I Schmidt¹ • T I Eglinton³ • C P McIntyre^{2,3,4}

¹Department of Geography, University of Zurich, Switzerland.

²Laboratory of Ion Beam Physics, ETH Zurich, Switzerland.

³Geological Institute, ETH Zurich, Switzerland.

⁴Present address: AMS Laboratory, SUERC, East Kilbride, United Kingdom.

ABSTRACT. Compound-specific radiocarbon analysis (CSRA) of benzene polycarboxylic acids (BPCAs) yields molecular-level, source-specific information necessary to constrain isotopic signatures of pyrogenic carbon. However, the purification of individual BPCAs requires a multistep procedure that typically results in only microgram quantities of the target analyte(s). Such small samples are highly susceptible to contamination by extraneous carbon, which needs to be minimized and carefully accounted for in order to yield accurate results. Here, we undertook comprehensive characterization and quantification of contamination associated with molecular radiocarbon (¹⁴C) BPCA analyses through systematic processing of multiple authentic standards with both fossil and modern ¹⁴C signatures at various concentrations. Using this approach, we precisely apportion the contribution of extraneous carbon with respect to the four implemented subprocedures. Assuming a constant source and quantity of extraneous carbon we correct and statistically evaluate uncertainties in resulting ¹⁴C data. Subsequently, we examine the results of triplicate analyses of reference materials representing four different environmental matrices (sediment, soil, aerosol, riverine natural organic matter) and apportion their BPCA sources in terms of carbon residues derived from biomass or fossil fuel combustion. This comprehensive approach to CSRA facilitates retrieval of robust ¹⁴C data, with application in environmental studies of the continuum of pyrogenic carbon.

KEYWORDS: BPCA, constant contamination, extraneous carbon, pyrogenic carbon, radiocarbon.

INTRODUCTION

The incomplete combustion of biomass and fossil fuels generates solid pyrogenic carbon (PyC) residues, which occur ubiquitously in the environment (Bird et al. 2015). Investigations of these products of incomplete combustion are particularly challenging in complex matrices due to a large variety of possible structural configurations of condensed aromatic compounds (Bird et al. 2015; Blumer 1976; Hedges et al. 2000). In addition to conventional methods that trace PyC on a bulk level (Gustafsson et al. 2001; Masiello et al. 2002; Meredith et al. 2012), specific polycyclic aromatic hydrocarbons (PAHs) and benzene polycarboxylic acids (BPCAs) can serve as molecular-level tracers of combustion residues (Blumer and Youngblood 1975; Glaser et al. 1998; Lima et al. 2005). Recently, we have shown that combined analysis of the abundances and distributions of both PAHs and BPCAs yields a more comprehensive view of the continuum of combustion residues occurring in environmental matrices (Hanke et al. 2016). It follows that the addition of corresponding ¹⁴C signatures could greatly enhance our understanding of the origin of PyC accumulating in the environment. Radiocarbon (¹⁴C) analyses of combustion products in environmental samples have the unique ability of apportioning sources of PyC to either biomass or fossil fuel combustion (Currie and Murphy 1977). However, the preparation of microgram quantities of sample material for ¹⁴C analysis at the molecular level requires an extended chemical scheme and careful attention in order to minimize the addition of extraneous carbon prior to isotopic measurement.

BPCA molecular markers are produced by the oxidative chemical degradation of combustion residues (charred biomass) and condensates (soot-like structures) (Glaser et al. 1998; Roth et al. 2012).

*Corresponding author. Email: umhanke@googlemail.com.

The BPCAs are composed of benzene rings with a varying number of carboxylic acid groups (Glaser et al. 1998). The method was originally developed for quantitative gas chromatographic (GC) measurement of BPCAs as their corresponding methyl ester derivatives (Glaser et al. 1998), this GC-based has subsequently been extended with a focus on compound-specific ^{14}C analysis (CSRA) of BPCAs (Ziolkowski and Druffel 2009; Coppola et al. 2013). In these latter studies, BPCAs were isolated via preparative capillary gas chromatography (PCGC) and subsequently graphitized for accelerator mass spectrometry (AMS) measurements. More recently, the BPCA analysis method has been adapted for liquid chromatography (LC) (Dittmar 2008; Schneider et al. 2011). This LC-based method allows for preparative isolation of individual BPCAs for isotopic analysis without derivatization prior to chromatographic separation (Gierga et al. 2014; Wiedemeier et al. 2016).

Advances in AMS technology now enable ^{14}C measurements of carbon dioxide gas at natural abundance levels on an ultra-microscale (2–100 $\mu\text{g C}$) (Fahrni et al. 2013; Wacker et al. 2013). Direct measurements of CO_2 supersede and negate the additional steps required for graphitization, albeit at the expense of precision. Recently, Lang and collaborators have developed a method based on wet chemical oxidation (WCO) for ^{14}C analysis of dissolved organic carbon (Lang et al. 2013, 2016). Compared with conventional approaches, such as those employing ultraviolet radiation (Armstrong et al. 1966; Beaupré et al. 2007), closed-tube combustion with cupric oxide (Boutton et al. 1983; Mann et al. 2015), or oxidation in the presence of permanganate (Leonard et al. 2013), the WCO method serves as a simple and effective approach that increases sample throughput and reduces costs. It is especially well suited to CSRA of water-soluble carbon compounds with low volatility. It has recently been demonstrated that BPCAs can be purified using preparative LC (LC_{prep}), then oxidized via WCO and measured for ^{14}C using a gas ion source-equipped AMS (Gierga et al. 2014). The integration of multiple process standards with minimal additional effort introduces a new level of blank assessment that can greatly improve the quality of small-scale ^{14}C data.

Preparative chromatographic purification of individual compounds commonly yields less than 50 $\mu\text{g C}$, thus molecular-level ^{14}C techniques are prone to extraneous carbon (McNichol and Aluwihare 2007) introduced during chemical pretreatment, purification and combustion steps. The smaller the sample sizes are, the higher the impact of extraneous carbon, particularly at the microgram scale level (5–25 μg) (Kirner et al. 1995; Shah and Pearson 2007; Santos et al. 2010). To obtain accurate results, a careful assessment and correction of CSRA data is required, this is most readily achieved using the model of constant contamination, i.e. assuming a constant source and quantity of extraneous carbon. This represents a simple yet comprehensive approach to derive robust estimates of the true ^{14}C value of unknowns via mass balance and propagation of errors (Shah and Pearson 2007; Hwang and Druffel 2008; Ziolkowski and Druffel 2009; Santos et al. 2010; Wacker and Christl 2012). Contamination in sample materials comes from multiple sources and can carry fossil (e.g. petrochemical-derived solvents) as well as modern (e.g. ambient air) ^{14}C signatures. Thus, an assessment of extraneous carbon requires the use of ^{14}C standards with complementary $^{14}\text{C}_{\text{depleted}}$ and $^{14}\text{C}_{\text{modern}}$ signatures in order to enable the determination of actual mass and $F^{14}\text{C}$ of the total contamination (Mollenhauer and Rethemeyer 2009).

Assessing the extraneous carbon in multistep CSRA procedures and correcting the measured data for the total contamination is challenging yet crucial in order to yield refined molecular isotopic data that allows for a robust interpretation of the results. In the present study, we analyze a variety of process standards that encompass the full range of subprocedures. Additionally, the potential of the ^{14}C -BPCAs methodology for real-world samples was tested using triplicate analysis of four different standard reference materials. The objectives of this study

were to (1) quantify extraneous carbon in subprocedures of ^{14}C -BPCA analyses, (2) allocate sources of extraneous carbon and propagate the errors, and (3) derive robust $F^{14}\text{C}$ values for BPCAs from four reference materials.

MATERIALS AND METHODS

Overview

The methodology used for BPCAs compound-specific isotopic analyses has been described and visually presented in detail elsewhere (Wiedemeier et al. 2016). Here, the focus is the characterization of the subprocedures; an overview of the steps and standard materials is given below and in Table 1. Each subprocedure uses a pair of ^{14}C -modern and ^{14}C -depleted materials (Table S1). For the entire procedure, we used chromatography-grade solvents and glassware that had recently (<48 hr) been precombusted (5 hr, 500°C).

Chemical Pretreatment: PyC Oxidation and BPCA Extraction

The archeological (MA) and laboratory (W450) model charcoals are described in a previous study (Gierga et al. 2014). The archeological charcoal was found in pyroclastic deposits at Maninjau, Sumatra (Alloway et al. 2004) and had an $F^{14}\text{C} = 0.003 \pm 0.001$ (ETH-50456) whereas the laboratory charcoal was obtained from the pyrolysis of chestnut woodchips under controlled conditions in the laboratory (Hammes et al. 2006) and had an $F^{14}\text{C} = 1.149 \pm 0.004$ (ETH-50457). For both samples, bulk ^{14}C values were determined after acid-base-acid pretreatment ahead of the analysis to remove extraneous organic matter (Gierga et al. 2014).

An initial mass of model charcoals was back-calculated from previous data, targeting concentrations of $100 \pm 30 \mu\text{g C mL}^{-1}$ (per vial) for the sum of (1,2,3,4,5)- and (1,2,3,4,5,6)-benzene polycarboxylic acids. Weighed sample masses were heated with 2 mL HNO_3 (65%) in Teflon pressure chambers at 170°C for 8 hr. After a vacuum filtration using glass syringes (6 mL) and a glass fiber filter (Chromabond® 0.7 μm mesh size) an aliquot (40%) was pipetted onto a cation exchange resin (Dowex® 50WX8) and then further cleaned by solid phase extraction (Discovery® DSC-18, 500 mg tubes) before being transferred into analytical vials with 1 mL ultrapure water (Milli-Q® Advantage A10, ≤ 3 ppb TOC).

Table 1 BPCA ^{14}C subprocedures including the type and range of different process standards and always consisting of both a ^{14}C depleted and modern material.

Sub-procedures	Work steps	Model Charcoals Geological samples	Model BPCAs Pure compounds	Sucrose and Phthalic acid Pure compounds	HOxII* and ancient CO_2 Diluted in He
Chemical pretreatment	Oxidative degradation of condensed aromatic structures	✓	✗	✗	✗
Purification	Isolation of compounds using preparative LC	✓	✓	✗	✗
Wet chemical oxidation	Closed tube combustion with sodium persulfate	✓	✓	✓	✗
Measurement	Gas ion-source equipped AMS	✓	✓	✓	✓

*NIST SRM4990C.

Purification: Preparative Liquid Chromatography (LC_{prep})

Commercially available (1,2,3)-, (1,2,4)-, (1,2,4,5)-, (1,2,3,4,5)- and (1,2,3,4,5,6)-benzene polycarboxylic acid ($\geq 98\%$, Fluka and Aldrich) were used for the quantification of BPCAs in external standard solutions. These five pure BPCA compounds each had bulk ^{14}C values of $F^{14}\text{C} \leq 0.0008$ (ETH-57051 to 57055). For the modern counterpart, we used (1,2,4)-benzene polycarboxylic acid ($\geq 98\%$) prepared by custom synthesis (Chempur, Germany) with a bulk $F^{14}\text{C} = 1.0422 \pm 0.0035$ (ETH-62938). Solutions of $200 \mu\text{g mL}^{-1}$ BPCAs in ultrapure water were equivalent to an ideal concentration of BPCAs in environmental samples aiming at yields of $30 \mu\text{g C}$ in purified fractions.

The LC_{prep} method for BPCA ^{14}C had a total runtime of 30 min, using a reversed phase analytical C-18 column (Agilent, $2.7 \mu\text{m}$) with two mobile phases: ultrapure water ($1.7\% \text{H}_3\text{PO}_4$) and acetonitrile ($\geq 99.98\%$, Scharlau, $F^{14}\text{C} < 0.004$, ETH-74445). The elution times for BPCAs varied from 5.98 to 19.05 min depending on their polarity and structural features. Target compounds included either individual BPCA or combined BPCAs consisting of (1,2,3,4,5,6)-BPCA and (1,2,3,4,5)-BPCA for all environmental samples. After quantifying the respective compounds in solution ($1 \mu\text{L inj. vol.}^{-1}$) using a diode array detector (60 mm path length) the injection volume (1 to $20 \mu\text{L inj.}^{-1}$) was adjusted to equal $1 \mu\text{g C BPCAs}^{-1}$ at this level baseline separation was maintained. All BPCAs were collected in separate vials ($\leq 6 \text{mL}$) in repetitive runs and the fractions were placed on a heating plate ($\sim 70^\circ\text{C}$) under a gentle ultrapure nitrogen stream to evaporate the solvent for 3 hr. Subsequently, samples were stored in screw-cap vials with Teflon seals in the refrigerator until the next analytical step. We also collected blank eluents with a time window immediately preceding the target BPCA. The time window was 0.3 min for (1,2,3,4,5,6)-BPCA and 0.24 min for all the other BPCAs. Blanks were spiked with either sucrose or phthalic acid in concentrations of 15, 25, $40 \mu\text{g C}$ prior to the wet chemical oxidation (WCO) step.

Wet Chemical Oxidation

The WCO procedure was performed according to Lang et al. (2013). Standards of phthalic acid ($\geq 99.5\%$, Sigma-Aldrich) and sucrose ($\geq 99.5\%$, Sigma) were dissolved in ultrapure water at a concentration of $2 \mu\text{M}$ and $1.5 \mu\text{M}$ respectively. The bulk ^{14}C values were $F^{14}\text{C} \leq 0.0008$ (ETH-57049) for phthalic acid and $F^{14}\text{C} = 1.149 \pm 0.0048$ (ETH-57050) for sucrose. The standard solutions were prepared with the identical ultrapure water we had used for the preparation of the oxidant on the same day.

The WCO procedure converts BPCAs to carbon dioxide (CO_2), which then can directly be measured for ^{14}C using a gas ion source-equipped AMS (Lang et al. 2013). About $30 \mu\text{g C sample}^{-1}$ and 5 mL of purified sodium persulfate solution were transferred into 12 mL gas-tight borosilicate Exetainer vials with butyl rubber septum (Labco, UK) using acidified ultrapure water. The screw-capped vials were then purged with ultra high purity helium at 100mL min^{-1} for 8 min. Oxidation of the samples is achieved by placing the vials in a heating block at 95°C for 1 hr (Lang et al. 2016).

Measurement: Gas Ion Source AMS

Gaseous samples were measured with a MIni CARbon DAting System (MICADAS, Ionplus, Switzerland) AMS fitted via a carbonate handling system modified with a needle for sparging solutions (Fahrni et al. 2013; Wacker et al. 2013). Briefly, sample CO_2 (10 – $100 \mu\text{g C}$) was sparged from the solution in septa-sealed vials onto a zeolite molecular sieve trap. The CO_2 gas was subsequently thermally desorbed and transferred to a gas-tight syringe with helium, then slowly injected into the gas-accepting ion source of the AMS at a constant rate (Lang et al. 2016). The measurement time was set to a maximum of 8 min in order to attain a precision of

better than 1% on modern standard. The measured background was better than 44,000 ¹⁴C yr. AMS data were normalized and background corrected with in-house standard gases prepared from Oxalic acid II (NIST SRM 4990C, F¹⁴C = 1.3407; ETH-54368) and ancient carbon dioxide (F¹⁴C ≤ 0.0041; ETH-41218). Gas samples (10–60 μg C) were measured in a single sequence of in total 64 AMS measurements.

Statistical Evaluation of Radiocarbon Data

Data were processed using BATS software (Wacker et al. 2010) and the measured ¹⁴C/¹²C ratios are reported as F¹⁴C according to Reimer et al. (2004). To extract the F¹⁴C of samples (F¹⁴C_S) from the measured AMS data (F¹⁴C_M), we applied the model of constant contamination represented by Equation (1)

$$F^{14}C_S = \frac{F^{14}C_M * m_M - F^{14}C_C * m_C}{m_M - m_C} \quad (1)$$

where F¹⁴C_C is the contaminant F¹⁴C, the total mass of the sample measured by the AMS (m_M) is the sum of the mass of the sample (m_S) and the mass of the contaminant (m_C) (Wacker and Christl 2012). The associated uncertainty (σ_{F¹⁴C_S}) of F¹⁴C_S was calculated by propagation of respective errors according to Equation (2)

$$\begin{aligned} \sigma_{F^{14}C_S}^2 = & \left[\sigma_{m_C} \left(\frac{F^{14}C_M * m_M - F^{14}C_C * m_C}{(m_M - m_C)^2} \right) - \left(\frac{F^{14}C_C}{m_M - m_C} \right) \right]^2 \\ & + \left[\sigma_{m_M} \left(\frac{F^{14}C_M}{m_M - m_C} - \frac{F^{14}C_M * m_M - F^{14}C_C * m_C}{(m_M - m_C)^2} \right) \right]^2 \\ & + \left[\sigma_{F^{14}C_M} \frac{m_M}{m_M - m_C} \right]^2 + \left[\sigma_{F^{14}C_C} \frac{-m_C}{m_M - m_C} \right]^2 \end{aligned} \quad (2)$$

In order to obtain F¹⁴C_S, it is necessary to first determine the contaminant F¹⁴C_C and m_C. For this, we used multiple process standards of known F¹⁴C in concentrations that spanned the range of measured samples (10–60 μg C). Pairs of complementary ¹⁴C standard materials, one ¹⁴C depleted (F¹⁴C = 0) and the other modern (F¹⁴C ≥ 1), were used for evaluation of the contamination using an end-member approach. Different standards were used to assess the contamination of the different subprocedures used in sample preparation (Table 1).

The depleted and modern process standards were used to determine the modern and the dead contamination, respectively, of the sample.

F¹⁴C_C is then

$$F^{14}C_C = \frac{F^{14}C_{C_{modern}} * m_{C_{modern}}}{m_{C_{depleted}} + m_{C_{modern}}} \quad (3)$$

and m_C

$$m_C = m_{C_{depleted}} + m_{C_{modern}} \quad (4)$$

for the respective subprocedures. The mass of ¹⁴C-depleted and modern contamination was determined by a fit (minimizing the chi-squares) to the measured standards. The amount

of m_C , separated into ^{14}C -depleted and modern components, was varied to evaluate the best fit of the modeled curves to the measured data. Reduced chi-squared statistics were then used to assess the statistical quality of the evaluation (Pearson's chi-square test). Subsequently the modern and fossil contamination was used to calculate $F^{14}\text{C}_C$ (Equation 3) and the m_C (Equation 4) for the different subprocedures. The uncertainties in the modeled $\sigma F^{14}\text{C}_C \pm \sigma_{mc}$ cannot be determined independently; we thus assigned a 10% uncertainty for each parameter as it accounts well for the observed scatter of the measured process standards. The proportional contribution of contamination for the individual steps of sample preparation was then calculated by subtraction. An example spreadsheet of calculations for the charcoal samples is given in the supplemental information.

Environmental Samples: Reference Materials

Four reference materials representing diverse yet common types of environmental samples were selected and analyzed for bulk ^{14}C and BPCA ^{14}C . Marine sediment Standard Reference Material (SRM) 1941b and urban dust SRM 1649a were obtained from the U.S. National Institute of Standards and Technology (NIST) and full details are given in the certificate of analysis (NIST 2007; 2015). The chernozem soil black carbon reference material was obtained from University of Zurich. Further details are given in Hammes et al. (2008) and online at <http://www.geo.uzh.ch/en/units/2b/Services/BC-material.html>. Natural organic matter (NOM) from the Suwannee River (2R101N) was obtained from the International Humic Substances Society (IHSS). Samples were analyzed for bulk ^{14}C using conventional graphite preparation at the laboratory of ion beam physics, ETH Zürich. CSRA of BPCAs was performed in triplicate, with resulting data used to apportion fossil ($F^{14}\text{C} = 0$) and modern ($F^{14}\text{C} \sim 1$) sources of combustion products according to Currie et al. (1989). To account for different possible modern sources we used two end members: $F^{14}\text{C}_{\text{contemporary}}$, which is the year of sampling (annual production) and $F^{14}\text{C}_{\text{biomass}}$ integrating 30 years of biomass growth (perennial production) similar to Xu et al. (2012).

RESULTS AND DISCUSSION

Comprehensive characterization and quantification of extraneous carbon in small-scale ^{14}C analyses allow sample ^{14}C values to be accurately determined and facilitates robust data interpretation. Figure 1 illustrates the results of measured process standards (^{14}C modern and depleted) for the three subprocedures—*chemical pretreatment*, *purification*, and *WCO*. The measured results on reference materials can be found in Table 2, including bulk $F^{14}\text{C}$, ^{14}C -BPCA uncorrected (after normalization of MICADAS instrumental measurement uncertainty) and corrected values including the propagation of errors from chemical pretreatment, purification, WCO and AMS. Worthy of note is that after the preparative purification of target compounds, all samples were combusted on a single day ($n = 55$) and then measured on the AMS in a single sequence the following day. The processed AMS data (BATS software, sequence: #C160520UHG1) provided by ETH Zurich is included in the supporting information (Table S2).

Constant Contamination

In small-scale ^{14}C analyses, the total mass measured comprises the mass of the sample and the mass of the contaminant, the latter of which can be assessed via external process standards of known $F^{14}\text{C}$, as well as via spiked process blanks (Lang et al. 2016; Pearson et al. 1998; Ziolkowski and Druffel 2009). Encompassing the entire BPCA ^{14}C analysis, the process standard charcoals yield a constant contamination of $3.50 \pm 0.35 \mu\text{g C}$ and $0.443 \pm 0.044 F^{14}\text{C}$.

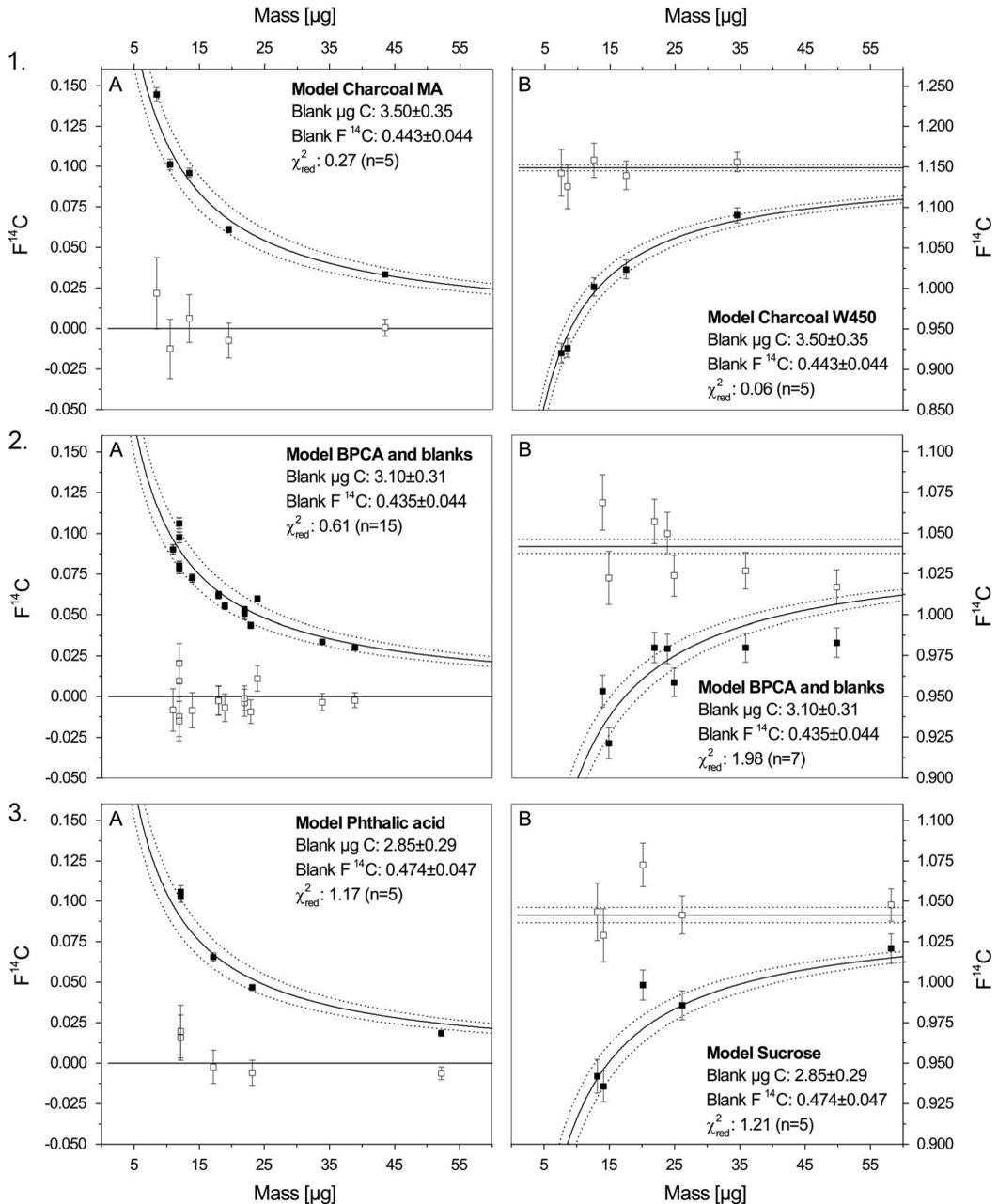


Figure 1 Measured AMS data are plotted as ^{14}C content ($F^{14}\text{C}$) versus sample size ($\mu\text{g C}$) for measured process standards (solid black squares) for the different subprocedures with $A = F^{14}\text{C}$ depleted standards and $B = F^{14}\text{C}$ modern standards: 1. Chemical pretreatment (process standards that span the entire procedure), 2. Purification (preparative chromatography standards) and 3. Wet Chemical Oxidation (WCO standards). The solid curved lines with the 1-sigma error ranges (dotted) are the chi-square fitted curves. The open squares show the corrected $F^{14}\text{C}$ with propagated uncertainties that scatter around the known bulk $F^{14}\text{C}$ value of the standard materials (solid straight lines).

Table 2 Results from bulk $F^{14}C$ and BPCAs ^{14}C analysis for four different complex environmental matrices without ($F^{14}C_{\text{uncorrected}}$) and with correction for the external contamination ($F^{14}C_{\text{corrected}}$) using the model of constant contamination.

ETH #	Name	Reference	TC [%]	Bulk $F^{14}C$	Benzene polycarboxylic acids (BPCA)		
					[g kg ⁻¹]	$F^{14}C_{\text{uncorrected}}$	$F^{14}C_{\text{corrected}}$
68463-68465	Organics in marine sediment	NIST 1941b	3.0	$0.568 \pm 0.003^*$	1.39 ± 0.02	0.241 ± 0.002	0.217 ± 0.009
68466-68468	Soil chernozem	University of Zurich	2.0	$0.874 \pm 0.004^*$	2.09 ± 0.02	0.557 ± 0.004	0.574 ± 0.008
68460-68462	Urban dust	NIST 1649a	17.7	$0.509 \pm 0.002^*$	9.76 ± 0.10	0.126 ± 0.002	0.088 ± 0.009
68469-68471	Riverine natural organic matter	IHSS	50.7	$1.057 \pm 0.005^*$	20.85 ± 0.36	0.950 ± 0.005	1.020 ± 0.007

*Bulk $F^{14}C$ values refer to ETH#: 67206-67209.

The model BPCAs and blanks (Purification and WCO) are slightly lower with $3.10 \pm 0.31 \mu\text{g C}$ and $0.435 \pm 0.044 F^{14}C$ while the contamination for phthalic acid and sucrose (WCO) is $2.85 \pm 0.29 \mu\text{g C}$ and $0.474 \pm 0.047 F^{14}C$.

A few studies have previously reported extraneous carbon in CSRA using either a different methodological setup or only certain subprocedures of multistep analyses. For example, Lang and collaborators reported a total contamination for the same *WCO* procedure of 0.68 ± 0.26 to $1.05 \pm 0.23 \mu\text{g C}$ and 0.027 ± 0.015 to $0.107 \pm 0.051 F^{14}C$ (Lang et al. 2016). The feasibility study by Gierga and collaborators that provides the foundation for the present approach assessed the uncertainties from AMS as well as WCO, found a contamination of $1.6 \pm 0.2 \mu\text{g C}$ and $0.90 \pm 0.14 F^{14}C$ (Gierga et al. 2014). Another study focused on purification of a different suite of target compounds, tetra-ether lipids, via LC_{prep} (with closed-tube combustion), found a contamination of $1.0 \pm 0.2 \mu\text{g C}$ (Shah and Pearson 2007). Contamination in CSRA has also been reported for PCGC (with vacuum line and graphitization) with $<1.0 \pm 0.5 \mu\text{g C}$ for fatty acids and $<0.17 \mu\text{g C}$ for phospholipid fatty acids and alkanes (Santos et al. 2010). In the conventional ^{14}C -BPCA method using PCGC (with cupric oxide combustion and graphitization), the extraneous carbon has been estimated to be $2.0 \pm 1.0 \mu\text{g C}$ and $0.158 \pm 0.026 F^{14}C$ in 2011 and $1.6 \pm 0.9 \mu\text{g C}$ and $0.067 \pm 0.025 F^{14}C$ in 2012 (Coppola et al. 2013) demonstrating a good reproducibility of the method.

Allocation of Constant Contamination to Subprocedures

The multistep design of our approach allows the comprehensive assessment of extraneous carbon in CSRA and the quantitative attribution of contamination to individual subprocedures. As all process standards incrementally include one to three subprocedures, the partial contamination can be calculated backwards from the *WCO* step. Here, the contamination by WCO accounts for 81% ($2.85 \pm 0.29 \mu\text{g C}$) of total contamination carrying a ^{14}C signature of $0.474 \pm 0.047 F^{14}C$. Additionally, the LC_{prep} adds $0.25 \pm 0.12 \mu\text{g C}$ ($F^{14}C = 0.0 \pm 0.5$) indicating that the source of the contaminant is ^{14}C depleted and probably stems from the incomplete removal of LC mobile phase acetonitrile. The subprocedure *chemical pretreatment* adds $0.40 \pm 0.14 \mu\text{g C}$ ($F^{14}C = 0.50 \pm 0.25$), suggesting that extensive sample transfers and exposure to ambient air may have introduced extraneous carbon from variable sources. Together, the *chemical pretreatment* and the *purification* procedures result in an accumulated contamination of $0.65 \pm 0.14 \mu\text{g C}$ ($F^{14}C = 0.31 \pm 0.06$). The propagation

of error calculations returned uncertainties of about 20 to 50%, depending on the strength of the statistical fit of modeled and measured data. The measured data for the subprocedures *WCO* and *chemical pretreatment* fall within the 1-sigma error range, whereas the *purification* procedure exhibits a slightly larger deviation in the modern standards (Figure 1). Nonetheless, it is evident that the vast majority of total contamination stems from the *WCO* procedure.

In considering the origin of the *WCO* contaminant, two possible sources may be taken into consideration—the purified oxidant sodium persulfate and the ultrapure water (5 mL sample⁻¹). A test measurement of the purified oxidant several days ahead of the AMS analysis sequence did not indicate any significant impurity. Thus we infer that the contamination likely originated from the water. About 2 L of MilliQ® water (TOC \leq 3 ppb) was subsampled on the day of the analyses. The water treatment system was maintained in good condition, including use of fresh (10-week-old) MilliQ® cartridges Q-Guard®T2 and Quantum®TEX. However, the distilled source water feeding the ultra-purification device can vary in purity, even though the displayed TOC may suggest an excellent water quality (\sim 3 ppb TOC). Further precautionary steps, such as UV oxidation of water immediately prior to use, are therefore recommended to minimize this potential source of contamination. Lang et al. (2016) showed that the contamination from *WCO* can be as low as 0.68 $\mu\text{g C}$ which is at least consistent with blank values reported for the other combustion techniques and would result in a total contamination of 1.33 $\mu\text{g C}$ if attained in this study. Additionally, this study used the same experimental setup as Lang et al. (2013) of 5 mL solvent (sample and oxidant in water) in 12 mL Exetainer vials yet, this could possibly be reduced to 3 mL solvent. This would ideally have lowered the contamination from *WCO* for 40%. However, such changes to the protocol require test measurements on the gas ion source-equipped AMS for revalidation.

Application and Source Apportionment to Reference Materials

Results for bulk and CSRA ^{14}C signatures and quantities of BPCAs for the four reference materials, representing different environmental settings and matrices, are given in Table 2. Triplicate BPCAs ^{14}C measurements on these standard materials were performed in order to test the applicability and reproducibility of the method. The average ^{14}C abundances in total carbon (TC) vary from $F^{14}\text{C} = 0.509 \pm 0.002$ to 0.874 ± 0.004 , implying that some proportion of TC in each sample is either of fossil origin or was pre-aged, thereby diluting the contemporary ^{14}C signatures at the time of deposition. In contrast, the riverine natural organic matter exhibits contemporary ^{14}C concentrations of 1.057 $F^{14}\text{C}$ (Suwannee River NOM, sampled via reverse osmosis in 2012).

BPCA quantification returned values of between 1.39 and 20.85 g kg^{-1} of sample, accounting for between 4.8 and 10.4% of TC. For CSRA, measured concentrations of (1,2,3,4,5,6)-BPCA and (1,2,3,4,5)-BPCA were between 78.1 ± 6.4 and $128.4 \pm 5.3 \mu\text{g C mL}^{-1}$. The masses of BPCAs ^{14}C recovered from reference materials varied between 24 to 36 $\mu\text{g C}$. The measured ^{14}C abundances of different BPCA ^{14}C encompassed almost the entire spectrum of natural abundance ^{14}C values, ranging from $0.088 \pm 0.009 F^{14}\text{C}_{\text{corrected}}$ for atmospheric particulates to $1.020 \pm 0.007 F^{14}\text{C}_{\text{corrected}}$ for the riverine NOM. The latter is indistinguishable from that of Coppola et al. (2015) who reported $1.02 \pm 0.02 F^{14}\text{C}$ for the same material using a different molecular ^{14}C approach.

Combustion-derived carbon can be apportioned to fossil and contemporary sources by $F^{14}\text{C}$ measurements (Currie and Murphy 1977; Currie et al. 1989; Currie 2000; Szidat et al. 2006).

In reference to the growth phase of biomass (May to August), we calculated the mean atmospheric $F^{14}C$ using the extended IntCal13 $^{14}CO_2$ concentration of the Northern Hemisphere (Levin and Kromer 2004; Levin et al. 2013). As trees integrate the atmospheric $F^{14}C$ over their period of growth (Lewis et al. 2004), we took a 30-year average prior to the sampling date in order to account for wood-derived combustion residues ($F^{14}C_{\text{biomass}}$). PyC from annual biomass (e.g. grasses) constitute an additional BPCAs source, thus we also considered the atmospheric ^{14}C abundance at the year of sampling ($F^{14}C_{\text{contemporary}}$) (Xu et al. 2012). Uncertainties include individual ^{14}C data as well as standard deviations from average atmospheric $F^{14}C$ values. Results of the NIST 1941b (marine sediment) yield a f_{fossil} of $82.9 \pm 2.8\%$ for $F^{14}C_{\text{biomass}}$ and $80.4 \pm 3.2\%$ for $F^{14}C_{\text{contemporary}}$. For comparison, Reddy et al. (2002) applied ^{14}C analysis on residues derived from the thermochemical oxidation (CTO-375) of environmental samples (Gustafsson and Gschwend 1997; Gustafsson et al. 2001) and found about 98.5–98.8% f_{fossil} for the predecessor NIST 1941a. BPCA ^{14}C analysis of NIST 1649a (urban dust) yields f_{fossil} $93.4 \pm 2.6\%$ for $F^{14}C_{\text{biomass}}$ and $93.5 \pm 2.6\%$ for $F^{14}C_{\text{contemporary}}$. Similar values were reported by Reddy et al. (2002) for CTO-375 with f_{fossil} of 93.5–95.2% and by Currie et al. (2002) for chemical oxidation residues (dichromate method) with f_{fossil} 88.7% as well as for thermal optical kinetic (TOK) procedure with f_{fossil} 97.4% always referring to $F^{14}C_{\text{contemporary}}$. Source apportionment for riverine NOM (2R101N) results in f_{fossil} of $8.9 \pm 3.0\%$ for $F^{14}C_{\text{biomass}}$ and $1.9 \pm 3.3\%$ for $F^{14}C_{\text{contemporary}}$.

Radiocarbon analyses on soils have a great potential for the study of carbon turnover, yet time lags between assimilation and incorporation in soils as well as lateral and vertical spatial heterogeneity can challenge interpretations (Trumbore 2009; Van der Voort et al. 2016). We therefore chose not to attempt source apportionment on the chernozem soil due to the absence of sufficient age constraints on charred soil organic matter (Carcaillet 2001).

The corrected ^{14}C -BPCA data compare well with the limited available literature data and lend support to the applicability of PyC source apportionment in different complex matrices using this particular method. Analyzed samples range from $F^{14}C < 0.1$ to $F^{14}C \geq 1$. Environmental matrices from locations with a pronounced human (urban, industrial) influence contain combustion products of mainly fossil origin. In contrast, the Suwannee river catchment drains swamp areas with a high fire return frequency and thus yields modern ^{14}C signatures. The two different proportions calculated for atmospheric ^{14}C concentrations ($F^{14}C_{\text{biomass}}$ and $F^{14}C_{\text{contemporary}}$) influence the fractional contribution of fossil carbon only to a less extent during this period of waning thermo-nuclear (“bomb”) ^{14}C concentrations.

Implications for Dating and Source Apportionment

Molecular ^{14}C data on pyrogenic carbon provides an important window on the origin and dynamics of this important component of the global carbon cycle. By combining (1) comprehensive characterization and quantification of contamination associated with sample processing for small-scale ^{14}C analysis, (2) assessments of reproducibility through triplicate analyses of samples, a new level of confidence in CSRA has been achieved. Adopting the constant contamination approach to account for contamination associated with analysis of reference materials led to lower BPCAs $F^{14}C$ values, with calculated fossil carbon contributions of 10.1, 30.2, 6.9% for marine sediment, urban dust and riverine NOM, respectively. As such, these carefully corrected molecular ^{14}C data provide a robust benchmark against which to place quantitative constraints on PyC sources and cycling in the environment. The large differences between the ^{14}C -BPCA_{uncorrected} and ^{14}C -BPCA_{corrected} (Table 2) clearly highlight the uncertainty that contamination can introduce with respect to PyC source apportionment.

Even though the extent to which contamination affects the result is mass dependent, an accurate F^{14}C signature of the contaminant is nevertheless required to make the correct adjustment. An interpretation of molecular ^{14}C data of combustion products without a correction for extraneous carbon could lead to either under- or overestimating different pools of PyC in samples. The need for accurate and precise data is particularly high for the validation of numerical models that mimic PyC pools and fluxes in the global environment. Currently, the extrapolation of our results to the global scale is limited by the lack of available quantitative and qualitative data, particularly in sediments (Bird et al. 2015) as well as by the high uncertainty (often larger than two orders of magnitude) in modeling the global pyrogenic carbon cycle (Landry and Matthews 2016). Here, the thoroughly corrected empirical ^{14}C data can serve as a reference to better constrain and validate models in future attempts.

In archaeology, visually distinct charcoal particles are commonly picked by hand (e.g., with the aid of an optical microscope) and used for ^{14}C dating. Here, measurement and comprehensive correction of BPCA ^{14}C data opens up the possibility of investigating matrices where charred residues are present as submicroscopic finely disseminated particles that would not be amenable to conventional approaches.

With respect to applications of ^{14}C -BPCA for the purpose of apportioning sources of combustion residues, calculations using the end member model of $\text{F}^{14}\text{C}_{\text{depleted}}$ and $\text{F}^{14}\text{C}_{\text{modern}}$ reveal that the two different scenarios for modern F^{14}C (contemporary vs. biomass) have a lesser impact (15.2% in sediment, 1.4% in urban dust, 7.6% in riverine NOM) compared to the correction of measured data for extraneous carbon (6.9 to 30.2%, as discussed above). This important observation strengthens the concept of source apportionment based on thoroughly corrected ^{14}C data. In this study, we only used reference materials (surface deposits) collected post-1950s and thus the modern F^{14}C end member should reflect the high atmospheric ^{14}C concentrations during the assimilation of carbon into biomass. This period of time might serve as an extreme case to test this concept of source apportionment and suggests an even lower susceptibility for samples predating this time period. Constraining the actual contemporary F^{14}C of feedstock materials in natural environments may prove challenging but fortunately it appears to be not imperative for the robust apportionment of PyC sources. Instead, differences in source areas (proximal versus distal), mode of transport (e.g. runoff versus atmospheric deposition), predepositional histories (e.g. temporary storage in soils) are likely to result in greater sources of uncertainty in apportioning PyC derived from fossil fuels and biomass.

CONCLUSIONS

Comprehensive quantification and isotopic characterization of extraneous carbon was achieved via use of multiple process standards that encompass each of the four subprocedures applied in BPCA ^{14}C analysis: (1) oxidative degradation of condensed aromatics, (2) BPCA isolation by LC_{prep} , (3) wet chemical oxidation (WCO) of purified BPCAs, and (4) the subsequent measurement of resulting CO_2 using a gas ion source-equipped AMS. Each subprocedure introduced some extraneous carbon, with the highest proportion determined to stem from the WCO step. Despite the significant extraneous carbon associated with this procedure, the ease and rapidity of WCO carries considerable advantages, enabling high-throughput, lower cost processing of standards and samples that ultimately yield more comprehensive BPCAs ^{14}C datasets.

The comprehensive assessment of (and correction for) contamination coupled with triplicate analyses of BPCAs isolated from four standard reference samples encompassing a range of

complex environmental matrices yielded robust ^{14}C values that provide new isotopic benchmarks for these reference materials. To fully exploit the potential of CSRA on pyrogenic carbon, particularly with respect to ^{14}C dating applications, the methodological protocol should include process standards, each encompassing end-members in ^{14}C natural abundance and spanning an equivalent concentration range to that of the target analytes. Following such comprehensive assessment of extraneous carbon, BPCAs ^{14}C can serve as a versatile tool for improving our understanding of fate and residence time of combustion products in the environment.

ACKNOWLEDGMENTS

We thank Simon Fahrni (Ionplus AG, Switzerland) for his kind on-call help with the AMS outside regular work hours. Michael Hilf and Guido Wiesenberg (University of Zurich, UZH, Switzerland) kept the preparative LC in good repair. We acknowledge Philippa Ascough and Michael Bird for providing us the archeological ^{14}C -depleted charcoal (MA), we appreciate the many helpful discussions on preparative chromatography with Daniel Montluçon (ETH Zürich, Switzerland) and Daniel Wiedemeier (UZH).

SUPPLEMENTARY MATERIAL

To view supplementary material for this article, please visit <https://doi.org/10.1017/RDC.2017.44>

REFERENCES

- Alloway BV, Pribadi A, Westgate JA, Bird MI, Fifield LK, Hogg A, Smith I. 2004. Correspondence between glass-FT and ^{14}C ages of silicic pyroclastic flow deposits sourced from Maninjau caldera, west-central Sumatra. *Earth and Planetary Science Letters* 227(1–2):121–33. doi: 10.1016/j.epsl.2004.08.014.
- Armstrong FAJ, Williams PM, Strickland JDH. 1966. Photo-oxidation of organic matter in sea water by ultra-violet radiation, analytical and other applications. *Nature* 211(5048):481–3. doi: 10.1038/211481a0.
- Beaupré SR, Druffel ERM, Griffin S. 2007. A low-blank photochemical extraction system for concentration and isotopic analyses of marine dissolved organic matter. *Limnology and Oceanography: Methods* 5(6):174–84. doi: 10.4319/lom.2007.5.174.
- Bird MI, Wynn JG, Saiz G, Wurster CM, McBeath AV. 2015. The pyrogenic carbon cycle. *Annual Review of Earth and Planetary Sciences* 43:273–98. doi: 10.1146/annurev-earth-060614-105038.
- Blumer M. 1976. Polycyclic aromatic compounds in nature. *Scientific American* 234(3):34–45. doi: 10.1038/scientificamerican0376-34.
- Blumer M, Youngblood WW. 1975. Polycyclic aromatic hydrocarbons in soils and recent sediments. *Science* 188(4183):53–5. doi: 10.1126/science.188.4183.53.
- Boutton TW, Wong WW, Hachey DL, Lee LS, Cabrera MP, Klein PD. 1983. Comparison of quartz and pyrex tubes for combustion of organic samples for stable carbon isotope analysis. *Analytical Chemistry* 55(11):1832–3. doi: 10.1021/ac00261a049.
- Carcaillet C. 2001. Soil particles reworking evidences by AMS ^{14}C dating of charcoal. *Comptes Rendus de l'Académie des Sciences – Series IIA – Earth and Planetary Science* 332(1):21–8. doi: 10.1016/S1251-8050(00)01485-3.
- Coppola AI, Ziolkowski LA, Druffel ERM. 2013. Extraneous carbon assessments in radiocarbon measurements of black carbon in environmental matrices. *Radiocarbon* 55(2–3):1631–40. doi: 10.2458/azu_js_rc.55.16303.
- Coppola AI, Walker BD, Druffel ERM. 2015. Solid phase extraction method for the study of black carbon cycling in dissolved organic matter using radiocarbon. *Marine Chemistry* 177(5):697–705. doi: 10.1016/j.marchem.2015.10.010.
- Currie LA. 2000. Evolution and multidisciplinary frontiers of ^{14}C aerosol science. *Radiocarbon* 42(1): 115–26. doi: 10.1017/S003382220005308X.
- Currie LA, et al. 2002. A critical evaluation of interlaboratory data on total, elemental, isotopic carbon in the carbonaceous particle reference material, NIST SRM 1649a. *Journal of Research of the National Institute of Standards and Technology* 107(3):279–98.
- Currie LA, Murphy RB. 1977. Origin and residence times of atmospheric pollutants: application of ^{14}C . *National Bureau of Standards Special Publication* 464:439–47.

- Currie LA, Stafford TW, Sheffield AE, Klouda GA, Wise SA, Fletcher RA, Donahue DJ, Jull AJT, Linick TW. 1989. Microchemical and molecular dating. *Radiocarbon* 31(3):448–63. doi: 10.2458/azu_js_rc.31.1171.
- Dittmar T. 2008. The molecular level determination of black carbon in marine dissolved organic matter. *Organic Geochemistry* 39(4):396–407. doi: 10.1016/j.orggeochem.2008.01.015.
- Fahrni SM, Wacker L, Synal HA, Szidat S. 2013. Improving a gas ion source for ¹⁴C AMS. *Nuclear Instruments and Methods in Physics Research B* 294:320–7. doi: 10.1016/j.nimb.2012.03.037.
- Gierga M, Schneider MPW, Wiedemeier DB, Lang SQ, Smittenberg RH, Hadjas I, Bernasconi SM, Schmidt MWI. 2014. Purification of fire derived markers for µg scale isotope analysis ($\delta^{13}\text{C}$, $\Delta^{14}\text{C}$) using high performance liquid chromatography (HPLC). *Organic Geochemistry* 70(1): 1–9. doi: 10.1016/j.orggeochem.2014.02.008.
- Glaser B, Haumeier L, Guggenberger G, Zech W. 1998. Black carbon in soils: the use of benzene-carboxylic acids as specific markers. *Organic Geochemistry* 29(4):811–9. doi: 10.1016/S0146-6380(98)00194-6.
- Gustafsson Ö, Bucheli TD, Kukulska Z, Andersson M, Largeau C, Rouzaud JN, Reddy CM, Eglinton TI. 2001. Evaluation of a protocol for the quantification of black carbon in sediments. *Global Biogeochemical Cycles* 15(4):881–90. doi: 10.1029/2000gb001380.
- Gustafsson Ö, Gschwend PM. 1997. Soot as a strong partition medium for polycyclic aromatic hydrocarbons in aquatic systems. *ACS Symposium Series* 671:365–81.
- Hammes K, Smernik R, Skjemstad JO, Herzog A, Vogt UF, Schmidt MWI. 2006. Synthesis and characterisation of laboratory-charred grass straw (*Oryza sativa*) and chestnut wood (*Castanea sativa*) as reference materials for black carbon quantification. *Organic Geochemistry* 37(11):1629–33. doi: 10.1016/j.orggeochem.2006.07.003.
- Hammes K, Smernik R, Skjemstad JO, Schmidt MWI. 2008. Characterisation and evaluation of reference materials for black carbon analysis using elemental composition, colour, BET surface area and ¹³C NMR spectroscopy. *Applied Geochemistry* 23(8):2113–22. doi: 10.1016/j.apgeochem.2008.04.023.
- Hanke UM, Eglinton TI, Braun ALL, Reddy CM, Wiedemeier DB, Schmidt MWI. 2016. Decoupled sedimentary records of combustion: causes and implications. *Geophysical Research Letters* 43(10): 5098–108. doi: 10.1002/2016GL069253.
- Hedges JI, et al. 2000. The molecularly-uncharacterized component of nonliving organic matter in natural environments. *Organic Geochemistry* 31(10):945–58. doi: 10.1016/S0146-6380(00)00096-6.
- Hwang J, Druffel ERM. 2008. Blank correction for delta-¹⁴C measurements in organic compound classes of oceanic particulate matter. *Radiocarbon* 47(1):75–87. doi: 10.2458/azu_js_rc.47.2802.
- Kirner DL, Taylor RE, Southon JR. 1995. Reduction in backgrounds of microsamples for AMS ¹⁴C dating. *Radiocarbon* 37(2):697–704. doi: 10.2458/azu_js_rc.37.1722.
- Landry J-S, Matthews HD. 2016. The global pyrogenic carbon cycle and its impact on the level of atmospheric CO₂, over past and future centuries. *Global Change Biology*. doi: 10.1111/gcb.13603.
- Lang SQ, Früh-Green GL, Bernasconi SM, Wacker L. 2013. Isotopic ($\delta^{13}\text{C}$, $\Delta^{14}\text{C}$) analysis of organic acids in marine samples using wet chemical oxidation. *Limnology and Oceanography-Methods* 11:161–75. doi: 10.4319/lom.2013.11.161.
- Lang SQ, McIntyre CP, Bernasconi SM, Früh-Green GL, Voss BM, Eglinton TI, Wacker L. 2016. Rapid ¹⁴C analysis of dissolved organic carbon in non-saline waters. *Radiocarbon* 58(3): 505–15. doi:dx.doi.org/10.1017/RDC.2016.17.
- Leonard A, Castle S, Burr GS, Lange T, Thomas J. 2013. A wet oxidation method for AMS radiocarbon analysis of dissolved organic carbon in water. *Radiocarbon* 55(2–3):545–52. doi: 10.2458/azu_js_rc.55.16277.
- Levin I, Kromer B. 2004. The tropospheric ¹⁴CO₂ level in mid latitudes of the Northern Hemisphere. *Radiocarbon* 46(3):1261–72. doi: 10.1017/S003822200033130.
- Levin I, Kromer B, Hammer S. 2013. Atmospheric D¹⁴CO₂ trend in Western European background air from 2000 to 2012. *Tellus B* 65(20092). doi: 10.3402/tellusb.v65i0.20092.
- Lewis CW, Klouda GA, Ellenson WD. 2004. Radiocarbon measurement of the biogenic contribution to summertime PM-2.5 ambient aerosol in Nashville, TN. *Atmospheric Environment* 38: 6053–61. doi: 10.1016/j.atmosenv.2004.06.011.
- Lima AL, Farrington JW, Reddy CM. 2005. Combustion-derived polycyclic aromatic hydrocarbons in the environment - a review. *Environmental Forensics* 6(2):109–31. doi: 10.1080/15275920590952739.
- Mann PJ, Eglinton TI, McIntyre CP, Zimov N, Davydova A, Vonk JE, Holmes RM, Spencer RGM. 2015. Utilization of ancient permafrost carbon in headwaters of Arctic fluvial networks. *Nature Communications* 6:7856. doi: 10.1038/ncomms8856.
- Masiello CA, Druffel ERM, Currie LA. 2002. Radiocarbon measurements of black carbon in aerosols and ocean sediments. *Geochimica et Cosmochimica Acta* 66(6):1025–36. doi: 10.1016/S0016-7037(01)00831-6.
- McNichol AP, Aluwihare LI. 2007. The power of radiocarbon in biogeochemical studies of the marine carbon cycle: insights from studies of dissolved and particulate organic carbon (DOC and POC). *Chemical Reviews* 107(2):443–66. doi: 10.1021/cr050374g.

- Meredith W, Ascough PL, Bird MI, Large DJ, Snape CE, Sun Y, Tilston EL. 2012. Assessment of hydrolysis as a method for the quantification of black carbon using standard reference materials. *Geochimica et Cosmochimica Acta* 97:131–47. doi: 10.1016/j.gca.2012.08.037.
- Mollenhauer G, Rethemeyer J. 2009. Compound-specific radiocarbon analysis – analytical challenges and applications. *IOP Conference Series: Earth and Environmental Science*. 5. doi: 10.1088/1755-1307/5/1/012006.
- NIST. 2007. *Certificate of Analysis Standard Reference Material 1649a*. National Institute of Standards and Technology.
- NIST. 2015. *Certificate of Analysis Standard Reference Material 1941b*. National Institute of Standards and Technology.
- Pearson A, McNichol AP, Schneider RJ, von Reden KF, Zheng Y. 1998. Microscale AMS ^{14}C measurement at NOSAMS. *Radiocarbon* 40(1): 61–75. doi: 10.2458/azu_js_rc.40.1989.
- Reddy CM, Pearson A, Xu L, McNichol AP, Benner BA Jr, Wise SA, Klouda GA, Currie LA, Eglinton TI. 2002. Radiocarbon as a tool to apportion the sources of polycyclic aromatic hydrocarbons and black carbon in environmental samples. *Environmental Science & Technology* 36(8): 1774–82. doi: 10.1021/es011343f.
- Reimer PJ, Brown TA, Reimer RW. 2004. Discussion: reporting and calibration of post-bomb ^{14}C data. *Radiocarbon* 46(3):1299–304. doi: 10.2458/azu_js_rc.46.4183.
- Roth PJ, Lehndorff E, Brodowski S, Bornemann L, Sánchez-García L, Gustafsson Ö, Amelung W. 2012. Differentiation of charcoal, soot, diagenetic carbon in soil: method comparison and perspectives. *Organic Geochemistry* 46:66–75. doi: 10.1016/j.orggeochem.2012.01.012.
- Santos GM, Southon JR, Drenzek NJ, Ziolkowski LA, Druffel ERM, Xu X, Zhang D, Trumbore SE, Eglinton TI, Hughen KA. 2010. Blank assessment for ultra-small radiocarbon samples: chemical extraction and separation versus AMS. *Radiocarbon* 52(2–3):1322–35. doi: 10.2458/azu_js_rc.52.3631.
- Schneider MPW, Smittenberg RH, Dittmar T, Schmidt MWI. 2011. Comparison of gas with liquid chromatography for the determination of benzenepolycarboxylic acids as molecular tracers of black carbon. *Organic Geochemistry* 42(3): 275–82. doi: 10.1016/j.orggeochem.2011.01.003.
- Shah SR, Pearson A. 2007. Ultra-microscale (5–25 $\mu\text{g C}$) analysis of individual lipids by ^{14}C AMS: Assessment and correction for sample processing blanks. *Radiocarbon* 49(1):69–82. doi: 10.1017/S0033822200041904.
- Szidat S, Jenk TM, Synal HA, Kalberer M, Wacker L, Hajdas I, Kasper-Giebl A, Baltensperger U. 2006. Contributions of fossil fuel, biomass burning, biogenic emissions to carbonaceous aerosols in Zurich as traced by ^{14}C . *Journal of Geophysical Research* 111(D7):1–12. doi: 10.1029/2005JD006590.
- Trumbore SE. 2009. Radiocarbon and soil carbon dynamics. *Annual Review of Earth and Planetary Sciences* 37:47–66. doi: 10.1146/annurev.earth.36.031207.124300.
- Van der Voort TS, Hagedorn F, McIntyre CP, Zell C, Walthert L, Schleppei P, Feng X, Eglinton TI. 2016. Variability in ^{14}C contents of soil organic matter at the plot and regional scale across climatic and geologic gradients. *Biogeosciences* 13:3427–39. doi: 10.5194/bg-13-3427-2016.
- Wacker L, Christl M. 2012. Data reduction for small radiocarbon samples – error propagation using the model of constant contamination. *ETH Zürich Laboratory of Ion Beam Physics: Annual Report 2011*.
- Wacker L, Christl M, Synal HA. 2010. Bats: a new tool for AMS data reduction. *Nuclear Instruments and Methods in Physics Research B* 268(7–8): 976–9. doi: 10.1016/j.nimb.2009.10.078.
- Wacker L, Fahrni SM, Hajdas I, Molnar M, Synal HA, Szidat S, Zhang YL. 2013. A versatile gas interface for routine radiocarbon analysis with a gas ion source. *Nuclear Instruments and Methods in Physics Research B* 294:315–9. doi: 10.1016/j.nimb.2012.02.009.
- Wiedemeier DB, et al. 2016. Characterization, quantification and compound-specific isotopic analysis of pyrogenic carbon using benzene polycarboxylic acids (BPCA). *Journal of Visualized Experiments*. 111. doi: 10.3791/53922.
- Xu L, Zheng M, Ding X, Edgerton ES, Reddy CM. 2012. Modern and fossil contributions to polycyclic aromatic hydrocarbons in PM_{2.5} from North Birmingham, Alabama in the Southeastern U.S. *Environmental Science & Technology* 46(3): 1422–9. doi: 10.1021/es2043189.
- Ziolkowski LA, Druffel ERM. 2009. Quantification of extraneous carbon during compound-specific radiocarbon analysis of black carbon. *Analytical Chemistry* 81(24):10156–61. doi: 10.1021/ac901922s.