





On small-scale and large-scale intermittency of Lagrangian statistics in canopy flow

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The interaction of fluids with surface-mounted obstacles in canopy flows leads to strong turbulence that dominates dispersion and mixing in the neutrally stable atmospheric surface layer. This work focuses on intermittency in the Lagrangian velocity statistics in a canopy flow, which is observed in two distinct forms. The first, small-scale intermittency, is expressed by non-Gaussian and not self-similar statistics of the velocity increments. The analysis shows an agreement in comparison with previous results from homogeneous isotropic turbulence (HIT) using the multifractal model, extended self-similarity and velocity increments' autocorrelations. These observations suggest that the picture of small-scale Lagrangian intermittency in canopy flows is similar to that in HIT and, therefore, they extend the idea of universal Lagrangian intermittency to certain inhomogeneous and anisotropic flows. Second, it is observed that the root mean square of energy increments along Lagrangian trajectories depends on the direction of the trajectories' time-averaged turbulent velocity. Subsequent analysis suggests that the flow is attenuated by the canopy drag while leaving the structure function's scaling unchanged. This observation implies the existence of large-scale intermittency in Lagrangian statistics. Thus, this work presents a first empirical evidence of intermittent Lagrangian velocity statistics in a canopy flow that exists in two distinct senses and occurs due to different mechanisms.

Key words: intermittency, atmospheric flows, mixing and dispersion

1. Introduction

Turbulent flows are often characterized by bursts of activity among long quiescent periods, and thus they are said to be intermittent. Intermittency can occur in turbulence in two different forms. The first is called small-scale intermittency, and it was first reported

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by Batchelor, Townsend & Jeffreys (1949) and reviewed by Frisch (1995) and Tsinober (2009). Small-scale intermittency is evident in statistics of velocity differences, both in the Eulerian and the Lagrangian frames, since their probability distribution functions (p.d.f.s) develop increasingly heavier tails as the scale of separation is reduced (e.g. Kailasnath, Sreenivasan & Stolovitzky 1992; Arnèodo et al. 2008). Despite numerous models that have been suggested, a comprehensive theory for small-scale intermittency is still missing (e.g. She & Leveque 1994; Elsinga, Ishihara & Hunt 2020), and yet it is believed to be a universal feature of high Reynolds number turbulence. The second kind of intermittency is termed large-scale intermittency, and it may occur due to variability of the flow at low frequencies. For example, transitions between the turbulent and non-turbulent states occur in jets or in transitional pipe flows (Corrsin 1943; Wygnanski & Champagne 1973), mesoscales can change local turbulence parameter in the atmospheric boundary layer (Muchinski, Frehlich & Balsley 2004) and strong large-scale velocity and temperature fluctuations can occur in stratified flows (Feraco et al. 2018). This work focuses on flows that are typical of the atmospheric surface layer, so-called canopy flows. In these flows, a fluid flow interacts with large surface-mounted obstacles, leading to high turbulence intensities. Furthermore, turbulence in canopies is said to be non-local since a significant fraction of turbulent kinetic energy is produced at the top of the obstacles and is then transported into the canopy layer itself (Finnigan 2000). The non-local character of the turbulence in canopy flows leads to large-scale intermittency inside the canopy, which is expressed by a velocity skewness, sparse extreme events of momentum and scalar fluxes or time-varying Hölder exponents (e.g. Finnigan 1979; Gao, Shaw & Paw U 1989; Louka, Belcher & Harrison 2000; Keylock et al. 2020). Therefore, canopy flows provide a fruitful ground for observing the two phenomena in conjunction, which is the aim of this work.

Intermittency in turbulence was studied previously mostly in the Eulerian framework, yet the advent of technological advances of the 2000s enabled empirical investigations in the Lagrangian framework as well (as reviewed by Toschi & Bodenschatz (2009)). Previous Lagrangian studies have revealed the existence of anomalous scaling of velocity differences (Chevillard et al. 2003; Arnèodo et al. 2008; Benzi et al. 2010; Huang et al. 2013) and has examined local flow features associated with extreme events (Liberzon et al. 2012; Xu et al. 2014; Watteaux et al. 2019) and proposed modelling strategies (Wilczek et al. 2013; Bentkamp, Lalescu & Wilczek 2019). These works focused on homogeneous isotropic turbulent flows and, inevitably so, focused on small-scale intermittency. Indeed, there is an absence of Lagrangian studies focusing on intermittency in inhomogeneous flows. In addition, Blum et al. (2010) showed that the Lagrangian structure functions depended on the magnitude of the instantaneous large-scale velocity in an oscillating grids experiment. Other than that, there is a lack of studies that focus on large-scale intermittency in the Lagrangian framework. In particular, there are no empirical investigations of intermittency in Lagrangian statistics in canopy flows despite its importance to Lagrangian stochastic models with applications for dispersion and mixing in the environment (Wilson & Sawford 1996; Reynolds 1998; Duman et al. 2016; Keylock et al. 2020; Shnapp et al. 2020; Viggiano et al. 2020).

This work presents an analysis of Lagrangian statistics in a canopy flow using empirical results from a recent wind tunnel experiment (Shnapp *et al.* 2019). The existence of small-scale intermittency is demonstrated in § 3.1, and the results are compared with previous studies from homogeneous isotropic turbulence (HIT) flows. Both qualitative and quantitative agreement is observed, which supports the idea of the universality of small-scale Lagrangian intermittency suggested by Arnèodo *et al.* (2008). Then, in § 3.2, it is demonstrated through conditional statistics that large-scale intermittency existed in

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Figure 1. A schematic diagram of several canopy obstacles in top view in the wind tunnel. The measurement volume was situated upstream of a tall canopy obstacle, and it is highlighted in red in the figure.

the canopy flow as well, and that although it affected the energetics of trajectories, it did not affect the scaling laws of structure functions.

2. Methods

Lagrangian trajectories in a canopy flow were analysed using the results of a wind tunnel, three-dimensional particle tracking velocimetry (3D-PTV) experiment. The full experimental details are given in Shnapp *et al.* (2019), and Lagrangian statistics were analysed in Shnapp *et al.* (2020). For brevity, only the information relevant to this work shall be repeated here.

The experiment was conducted in the environmental wind tunnel laboratory at the Israel Institute for Biological Research, which features a 14 m long open wind tunnel with a $2 \times 2 \text{ m}^2$ cross-sectional area. We used a double-height staggered canopy layout, in which flat plates of height *H* and $\frac{1}{2}H$ were placed in consecutive rows (H = 100 mm). The plates were thin, their width was $\frac{1}{2}H$, and the spacing between the rows was $\frac{3}{4}H$, as shown in the sketch in figure 1. The canopy frontal area density was $\lambda_f = A_f/A_T = \frac{9}{16}$, (where A_f is the element's frontal area and A_T is the lot area of the canopy layer), which categorizes our canopy as moderately dense. The wind velocity was $U_{\infty} = 2.5 \text{ m s}^{-1}$, corresponding to a Reynolds number of $Re_{\infty} = U_{\infty}H/\nu = 1.6 \times 10^4$ (ν is the kinematic viscosity). We recorded the trajectories using a real-time image analysis extension of the 3D-PTV method described in Shnapp *et al.* (2019). The PTV algorithms and the analysis were applied using the OpenPTV consortium (2014) open-source software and the Flowtracks package by Meller & Liberzon (2016). In this work, x is the streamwise direction, y is the horizontal spanwise direction and z is perpendicular to the bottom wall directed upwards, where z = 0corresponds to the bottom wall.

This work is focused on a subset of trajectories that were recorded in a small subvolume of space. The subvolume had a length of $\frac{3}{4}H$, width of $\frac{1}{2}H$, and it was situated at the top of the canopy layer, 0.9 < z/H < 1.1 (this is subvolume b3 in Shnapp *et al.* (2020), and it is shown in figure 1). The root mean square (RMS) of velocity fluctuations was $u' = 0.47 \text{ m s}^{-1}$, the mean dissipation rate was estimated as $\epsilon = 0.25 \text{ m}^2 \text{ s}^{-3}$, the Kolmogorov length scale was $\eta = 0.34 \text{ mm} \approx \frac{1}{290}H$ and the Taylor microscale Reynolds number was $Re_{\lambda} = 440$. Furthermore, the Lagrangian streamwise velocity decorrelation time scale was T = 54 ms, estimated by fitting Lagrangian velocity autocorrelation function. Notably, the



Figure 2. (*a*) Standardized p.d.f.s of Lagrangian temporal velocity increments at $\tau/\tau_{\eta} \in \{0.3, 3, 5, 8, 11\}$, translated vertically; symbols correspond to the empirical canopy flow data, black lines stand for the multifractal model and a Gaussian p.d.f. is shown as a thin grey line at the bottom. (*b*) The flatness of Lagrangian velocity increments plotted against the time lag; the results of the multifractal model shown as a black line and the Gaussian value F = 3 is marked by a dashed line.

decorrelation time scale varied for each velocity component, and the Lagrangian integral time scale T_L is not trivial to define, so T shall be used as a proxy for T_L for simplicity (see Shnapp *et al.* (2020) for a detailed discussion).

3. Results

3.1. Lagrangian velocity increments and small-scale intermittency

In the following section, the focus is put on small-scale intermittency. The Lagrangian temporal velocity increment, defined as

$$\Delta_{\tau} v_i(t_0) \equiv v_i(t_0 + \tau) - v_i(t_0), \tag{3.1}$$

where τ is the time lag, is widely used to study velocity statistics at different scales. Here, we use statistics of $\Delta_{\tau} v_i$ to show the existence of, and to analyse, small-scale intermittency in the canopy flow. Note that assuming stationarity of the flow in the wind tunnel, statistics are reported for different trajectories with different t_0 , namely, we average over t_0 . Furthermore, statistics of $\Delta_{\tau} v_i$ are assumed stationary in the range of τ considered here, due to the local homogeneity that we have shown in Shnapp *et al.* (2020).

The p.d.f.s, $P(\Delta_{\tau} v_x)$, for trajectories from the canopy flow experiment are shown in figure 2(*a*) as symbols for five values of τ . The p.d.f.s were translated vertically for better visualization. The figure shows that despite the average flow velocity and its inhomogeneity, the velocity increments are zero averaged. The figure also shows that as the time lag is reduced the tails of the p.d.f.s become wider, showing that at smaller scales there is a higher probability for observing extreme events. In addition, the flatness coefficient of the velocity differences is plotted in figure 2(*b*) against τ/τ_{η} . The empirical data is shown as symbols, error bars represent the range obtained using bootstrapping with five subsamples of the data, and the Gaussian value of F = 3 is shown as a dashed line. Due to the available volume of the data the flatness at small τ is underestimated in our analysis; and still at small τ the flatness is high, reaching roughly 17, and as the time lag grows it reduces monotonically and reaches down to $F \approx 5$. In the Kolmogorov similarity theory, dimensional analysis predicts that moments of the velocity difference scale with the time lag as $\langle (\Delta_{\tau} v_i)^q \rangle \sim \tau^{q/2}$ (Monin & Yaglom 1972), and so the flatness coefficient should remain constant in the inertial range, $\tau_{\eta} \ll \tau \ll T_L$. Thus, the change of $F(\tau)$ for $\tau \gg \tau_{\eta}$ shows the existence of deviation from the Kolmogorov similarity theory in the canopy flow. As discussed in § 1, this transition of the statistics with τ is a hallmark of turbulent flows that characterizes small-scale intermittency.

Chevillard *et al.* (2003) proposed that the transition from the flat to Gaussian p.d.f. in HIT can be described by the multifractal model, and showed that it was in good agreement with results from two experiments and direct numerical simulation (DNS) at various Re_{λ} . Briefly explained, in the multifractal formalism the velocity increments are specified as

$$\Delta_{\tau} v_i = \mathcal{B}\left(\frac{\tau}{T_L}\right) \Delta_{T_L} v_i, \qquad (3.2)$$

where $\Delta_{T_L} v_i$ is the velocity increments at long times and $\mathcal{B}(\tau/T_L)$ is a random function. Then, $P(\Delta_{\tau} v_i)$ can be calculated by integrating the p.d.f.s of \mathcal{B} and $\Delta_{T_L} v_i$, given a model for \mathcal{B} . Importantly, this work uses the same model for \mathcal{B} that was originally utilized by Chevillard *et al.* (2003) for studies of HIT flows, and it thus assumes the same singularity spectrum for the canopy flow; a full description of the model is given in appendix A. The resulting p.d.f.s that were calculated using the model are shown in figure 2(a) as continuous lines underlying the empirical data. The flatness coefficient that was calculated using the multifractal model is also plotted in figure 2(b) as a continuous line, showing a fair agreement between the empirical data and the model. The fair agreement between the empirical results and the model is important because we used here the same function \mathcal{B} . Indeed, the fact that using the same singularity spectrum we could obtain a close fit for statistics of our data suggests that there exists a similarity between the small-scale dynamics in the canopy flow and HIT, despite the strong inhomogeneity and anisotropy of the canopy flow.

The so-called Lagrangian structure functions are statistical moments of the velocity increments,

$$S_q(\tau) = \langle (\Delta_\tau v_i)^q \rangle. \tag{3.3}$$

As discussed above, in the so-called inertial range, $\tau_{\eta} \ll \tau \ll T_L$, the Kolmogorov similarity theory predicts that $S_q \propto \tau^{\zeta_q}$ with $\zeta_q = q/2$ (Monin & Yaglom 1972). Here, we would like to examine whether this prediction might hold in the canopy flow as well, while noting that for anisotropic flows like ours we may only speak of 'effective' scaling due to effects of anisotropy. Thus, the structure functions for q = 2, 4 and 6 are shown in log-log scales in the inset of figure 3(a). Indeed, no clear scaling region can be found for $\tau > \tau_n$ in the graph, however, this is a common feature that occurs in isotropic flows as well. It was suggested that the lack of scaling in isotropic flows may be a result of finite Reynolds number effects (Toschi & Bodenschatz 2009) and it can hinder theory validations and comparison between different experiments and numerical simulations. A commonly used method to bypass this difficulty is to use the so-called extended self-similarity framework (ESS), in which ζ_q is examined relative to ζ_2 (Toschi & Bodenschatz 2009); the ESS approach can extend the scaling range and it was found to successfully converge various previous experimental and numerical results (Arnèodo et al. 2008). Thus, in the main panel of figure 3(*a*), we examine ζ_q/ζ_2 by plotting S_q against S_2 in log–log scales for q = 4 and 6. The figure shows a narrow range $\tau_{\eta} < \tau \lesssim 4.5 \tau_{\eta}$ in which a scaling exists for the canopy flow experiment. Notably, the separation of scales in the canopy experiment was $T/\tau_{\eta} \approx 6$, which is very low as compared with homogeneous flows with similar Re_{λ} , due to the so-called rapid decorrelation that was explored by Shnapp et al. (2020), and this limited severely the extent of the scaling range of S_q/S_2 . Yet, in the existing range figure 3(a) gives $\zeta_4/\zeta_2 \approx 1.51$ and $\zeta_6/\zeta_2 \approx 1.81$. These values are in remarkable agreement with previous



Figure 3. (a) The inset shows Lagrangian structure functions, $S_q(\tau)$, for q = 2, 4 and 6; the main figure is an ESS plot that shows S_4 and S_6 against S_2 to probe relative scaling. (b) Lagrangian autocorrelation function of temporal velocity increments with $\tau = \tau_{\eta}$, shown for the three velocity components and for the magnitude of the full velocity vector. The inset is a 3-D representation of a convoluted trajectory in a box of size $(0.2H)^3$.

experimental results from HIT flows; for example, Mordant, Lévêque & Pinton (2004) found $\zeta_4/\zeta_2 = 1.54 \pm 0.06$ and $\zeta_6/\zeta_2 = 1.8 \pm 0.2$ for the $Re_{\lambda} = 570$ experiment (cf. table 4 there). The agreement we observe here is important, because it supports the argument of local homogeneity in canopy flows at small scales.

Let us briefly consider dynamical scenarios for small-scale Lagrangian intermittency. Results from HIT DNS by Biferale et al. (2005), Bec et al. (2006) and Bentkamp et al. (2019) suggested that small-scale intermittency is a result of encounters between particles and intense vortex filaments; indeed, Wilczek, Jenko & Friedrich (2008) showed that the characteristic transition of the increments' p.d.f.s can be captured by a heuristic flow model of superimposed constitutive vortices. Similarly, Liberzon et al. (2012) showed that acceleration-vorticity-strain alignment in a quasi-homogeneous flow is associated with intense energy flux. Here, we can show hints suggesting that similar scenarios occur in the canopy flow as well. As shown by Mordant et al. (2002, 2004), while Lagrangian acceleration components decorrelate on time scales of the order $\sim \tau_n$, the magnitude of acceleration decorrelates on much longer time scales. Here, figure 3(b)suggests that the same is true in our canopy flow data. It shows four autocorrelation functions: three for the increments of each of the velocity components (x, y and z) and one for increments of the magnitude of the velocity vector, taking the time lag $\tau = \tau_{\eta}$ (the velocity difference can be used as a proxy for the acceleration because, as Voth, Satyanarayan & Bodenschatz (1998) and Shnapp et al. (2020) showed, at such small time-lags the acceleration is still correlated, so approximately $\Delta_{\tau_{\eta}} v_i \approx \frac{\partial v_i}{\partial t} \tau_{\eta}$). While the three components' velocity increments became decorrelated ($\rho = 0$) at roughly $\tau \approx 2\tau_{\eta}$, the velocity magnitude difference retained correlation with itself over the whole range of the measurement, with the minimum value of around $\rho \approx 0.4$. This difference between the components' and the magnitude's autocorrelations agrees with the vortex trapping picture. In addition to that, for $\tau \gtrsim 2\tau_{\eta}$ the components' increments were anti-correlated, which, as shown by Wilczek et al. (2008), can result from trajectories rotation around vortex filaments' cores; this too supports the picture of vortex trapping. Thus, figure 3(b)supports the notion that small-scale Lagrangian intermittency in the canopy flow is related to the encounter of trajectories with sparse and intense vortex filaments, similar to the HIT case. The inset of figure 3(b) visualizes a convoluted trajectory, which is a possible instance of such a trapping scenario.

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Figure 4. (*a*) Normalized histogram for the number of trajectories in our dataset with each quadrant Q_i . (*b*) Joint p.d.f. of the streamwise and vertical velocity components averaged over the velocity decorrelation time scale, T.

3.2. Conditional statistics and large-scale intermittency

In the following section we use conditional statistics in order to detect large-scale intermittency. Consider the velocity of a certain Lagrangian trajectory between the times t_0 and $t_0 + \tau$: $v_{t_0,\tau} \equiv \{v(t) | t_0 \leq t < t_0 + \tau\}$. The Lagrangian average of a function in this section of time shall be denoted with a tilde symbol as

$$\widetilde{f(\boldsymbol{v})}_{t_0,\tau} \equiv \frac{1}{\tau} \int_{t_0}^{t_0+\tau} f(\boldsymbol{v}_{t_0,\tau}) \,\mathrm{d}t.$$
(3.4)

Note that such averages are properties of individual trajectories over periods of duration τ . Also, since we assume stationarity of the flow we present statistics for different trajectories, namely averaged over t_0 . We denote fluctuations of the trajectory averaged velocity with respect to the Eulerian mean velocity as $\tilde{v}'_{\tau} \equiv \tilde{v}_{\tau} - U$. Now, using (3.4) and in analogy to the Eulerian quadrant analysis (Antonia 1981; Shaw, Tavangar & Ward 1983; Raupach, Coppin & Legg 1986; Zhu, van Hout & Katz 2007), we define the Lagrangian quadrant of a trajectory using signs of the components of \tilde{v}'_{τ} on the *x* and *z* plane as follows:

$$Q_{i} \equiv \begin{cases} 1, & \text{if } \widetilde{v}_{xT}' > U_{x} \text{ and } \widetilde{v}_{zT}' > U_{z}, \\ 2, & \text{if } \widetilde{v}_{xT}' \leqslant U_{x} \text{ and } \widetilde{v}_{zT}' > U_{z}, \\ 3, & \text{if } \widetilde{v}_{xT}' \leqslant U_{x} \text{ and } \widetilde{v}_{zT}' \leqslant U_{z}, \\ 4, & \text{if } \widetilde{v}_{xT}' > U_{x} \text{ and } \widetilde{v}_{zT}' \leqslant U_{z}, \end{cases}$$

$$(3.5)$$

where we use the averaging time $\tau = T$, the Lagrangian velocity decorrelation time scale (Shnapp *et al.* 2020). Figure 4(*a*) shows a normalized histogram for the trajectories being associated with the four quadrant states. It shows that Q_2 trajectories were the most common, followed by Q_4 trajectories and then Q_1 and Q_3 trajectories, which is in qualitative agreement with the duration fractions reported by Yue *et al.* (2007) and Zhu *et al.* (2007). It is also interesting to see that the time-averaged Lagrangian velocity fluctuation components are correlated, similar to the Eulerian turbulent velocity components that make up the Reynolds stress. This is shown in figure 4(*b*) through the elliptical shape of the joint p.d.f. that is elongated in the direction of the negative diagonal; the correlation coefficient was -0.26. Below, conditional statistics based on Q_i are used to probe large-scale intermittency effects on the Lagrangian statistics.



Figure 5. Lagrangian statistics condition with (3.5). (*a*) The p.d.f.s of the activity A_T for four groups of trajectories divided according to their quadrant. (*b*) Second-order Lagrangian structure function for trajectories from different velocity quadrants.

The trajectories in the canopy flow experiment were observed to be associated with more/less strong changes of their kinetic energy $(e \equiv \frac{1}{2} |\boldsymbol{v}|^2)$ when trajectories were conditioned based on the value of Q_i . To demonstrate this, let us denote the following property:

$$A_{\tau} \equiv \left[e - E_{\tau} \right]^2_{\tau}^{1/2}, \tag{3.6}$$

where $E_{\tau} = \tilde{e}_{\tau}$ is the average kinetic energy of a trajectory. Thus, A_{τ} is the RMS of the kinetic energy increments that were discussed by Xu *et al.* (2014) along the path of a Lagrangian trajectory during a time τ . It is a non-negative scalar that quantifies the amplitude of kinetic energy changes undergone by a trajectory. Loosely speaking, it can be interpreted to show how active a trajectory is for fixed durations. In figure 5(*a*) we show p.d.f.s of A_T conditioned on Q_i , where note that again we use $\tau = T$. It is seen that A_T was typically the highest for trajectories with Q_4 or Q_1 , and that it was the lowest for trajectories with Q_2 . Also, the average of A_T for trajectories with Q_4 was more than two times higher than the average over trajectories at Q_2 , but only 20 % higher than the average over Q_1 trajectories. Notably, the p.d.f.s of A_T were roughly log-normal.

Figure 5(*a*) reveals anisotropy in the kinetic energy increments of Lagrangian trajectories, since statistics of A_T depended on the direction of trajectory's velocity fluctuations. While it is expected that statistics of A_T will depend on the magnitude of velocity even in HIT, a dependence on direction reveals a symmetry breaking that can only persist in inhomogeneous or anisotropic flows. Furthermore, since A_{τ} measures Lagrangian fluctuations of the kinetic energy, higher values of A_{τ} result from stronger forces that act on particles. Correspondingly, A_T was typically higher for both Q_4 and Q_1 which are associated with higher streamwise velocity (relative to U_x). This suggests that the changes in statistics of A_T are due to increased/decreased levels of the canopy drag that fluctuated due to large-scale flow structures in the shear layer and the boundary layer above the canopy. This is in qualitative agreement with Keylock *et al.* (2020) who recently associated streamwise velocity and intermittency in a canopy flow.

The central role of the energy cascade in the understanding of turbulent flows makes it important to examine how large-scale intermittency is reflected across the different scales. To this end, similar to Blum *et al.* (2010, 2011), we use conditioned structure functions. Specifically, S_q as defined in (3.3) is now calculated over groups of trajectories with the same Q_i . Figure 5(b) shows the conditioned S_2 , plotted on a log-log scale. The structure functions for different Q_i have nearly identical shapes, but they are translated vertically with respect to one another. In fact, the structure functions appear in the figure according to the average levels of A_T observed in figure 5(a): S_2 is highest for Q_4 ; then Q_1 ; Q_3 ; and the lowest is Q_2 . In addition to that, since the figure is in log-log scales, the identical shapes mean that the time scaling of structure functions relating to different quadrants is the same: $\zeta_2 \approx 2$ for $\tau \leq \tau_{\eta}$ and it reduces below 1 by the end of our measurement range. In particular, ζ_2 was almost independent of Q_i .

The important observation from figure 5(b) is that when conditioning on Q_i the changes in the examined statistics occurred homogeneously across the scales. However, results from other flows suggest that this is not a universal feature. For example, Sreenivasan & Dhruva (1998) and Blum et al. (2010, 2011) showed that conditioning samples on a representative large-scale velocity affected the scaling of Eulerian structure functions only in some flows while in other flows it did not. While a rigorous explanation of why this occurred in our flow is beyond the scope of the present work, we can suggest phenomenological reasoning. In canopy flows, the turbulence is severely obstructed by the canopy obstacles and the forcing acts mostly due to the interaction with fixed boundaries. This is different from free flows, in which turbulence production intrinsically depends on correlations between the flow and the forcing, as shown recently by Bos & Zamansky (2019). This consideration suggests that energy input occurred mostly on scales larger than our measurement volume, so canopy drag fluctuations did not significantly alter the structure functions on the scales available in figure 5(b). This observation is important for two reasons. First, it implies that changes in statistics when conditioning on Q_i occur due to variations in 'turbulence parameters', namely this is indeed large-scale intermittency. Second, it is important for Lagrangian near-field dispersion models since it suggests that temporal fluctuations in canopy drag may be treated by varying the simulation's parameters over long time scales, e.g. as discussed by Pope & Chen (1990), Pope (1991), Aylor (1990) and Duman et al. (2014, 2016).

4. Discussion and conclusions

To conclude, this work presents observations of both small-scale intermittency and large-scale intermittency of Lagrangian statistics in a canopy flow by using the results of a recent wind tunnel experiment. This is the first experimental observation of Lagrangian intermittency in a canopy flow, and thus, it presented a unique opportunity to probe these two different types of intermittency in parallel. Our results demonstrate the importance of direct Lagrangian investigations of inhomogeneous and anisotropic turbulent flows.

The Lagrangian small-scale intermittency was manifested by deviations of the velocity increment's statistics from self-similarity and, in particular, their flatness increased strongly when τ was decreased. Furthermore, a marked similarity was observed between our results for the canopy flow and previous observations from HIT. Specifically, using the Lagrangian multifractal model and the ESS framework, we found remarkable quantitative agreement with Chevillard *et al.* (2003) and Mordant *et al.* (2004). Lastly, the long correlation of acceleration magnitude and the short correlation of acceleration components suggests that the source for small-scale intermittency is, similar to HIT, rooted in encounters of particles with vortex filaments (Biferale *et al.* 2005; Bec *et al.* 2006; Wilczek *et al.* 2008; Bentkamp *et al.* 2019). These results strongly support the picture suggested by Arnèodo *et al.* (2008) of universal Lagrangian intermittency in turbulence, and it also suggests its extension to certain highly turbulent inhomogeneous and anisotropic flows.

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This observed similarity may have been due to the dominance of the isotropic dissipation terms over contributions from the flow's inhomogeneity to the particle's dynamics, as we reported in Shnapp *et al.* (2020). In this case, the main conclusion is that even in the presence of marked inhomogeneity and anisotropy, the HIT picture may still be relevant at small-scales if the turbulence energy flux is sufficiently high.

It was also observed that when conditioned on the direction of the time-averaged velocity fluctuation, Lagrangian trajectories had significantly different statistics for the RMS of kinetic energy increments. It was typically much higher (lower) for trajectories whose streamwise velocity component was higher (lower) than the mean. Correspondingly, the second-order Lagrangian structure functions were higher (lower) for these groups of trajectories. This suggests that fluctuations of the canopy drag force affect the activity of Lagrangian trajectories and, therefore, this observation is a manifestation of large-scale intermittency. Furthermore, it was observed that the large-scale intermittency did not affect the scaling of the Lagrangian structure functions, namely that the effect of conditional statistics was felt homogeneously across the different scales. This observation is important for the treatment of large-scale intermittency in Lagrangian dispersion models.

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Appendix A

In the main text, the multifractal model was used to support the argument that small-scale intermittency in the canopy flow reflects processes that are characteristic of fully developed turbulence. The specific formulation of the model we used follows the development by Chevillard *et al.* (2003). We only slightly modified it in order to fit the canopy data. According to Chevillard *et al.* (2003), the Lagrangian velocity differences are given by

$$\Delta_{\tau} v_i = \mathcal{B}\left(\frac{\tau}{T_L}\right) \Delta_{T_L} v_i,\tag{A1}$$

where \mathcal{B} is a random function, and their p.d.f. can be calculated as

$$P(\Delta_{\tau}v_i) = \int_{-1/2}^{+\infty} \frac{\mathcal{P}\left(h, \frac{\tau}{T_L}, Re, \mathcal{D}(h)\right)}{\mathcal{B}\left(h, \frac{\tau}{T}, Re\right)} \mathcal{G}\left(\frac{\Delta_{\tau}v_i}{\mathcal{B}\left(h, \frac{\tau}{T_L}, Re\right)}\right) dh.$$
(A2)

The function \mathcal{B} and its p.d.f. \mathcal{P} were calculated using the exact same specification as in Chevillard *et al.* (2003), and similarly, the p.d.f. of the increments of $\Delta_{T_L} v_i$ was assumed to be Gaussian. However, as we discussed in Shnapp *et al.* (2020), while the separation of scales T/τ_{η} in the HIT case is a function only of the Reynolds number $T/\tau_{\eta} = f(Re)$, in canopy flows it depends also on other macroscopic parameters of the flow, such as the geometry or the arrangement of the obstacles. Therefore, to fit the model to the canopy

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flow data we required an additional parameter, denoted ϑ , that adjusts the separation of scales to the measured values. Thus, we obtain the following formulation:

$$\mathcal{B}\left(h,\frac{\tau}{T_L}, Re, \vartheta\right) = \frac{\left(\frac{\tau}{T_L}\frac{1}{\vartheta}\right)^h}{\left[1 + \left(\frac{\tau}{\tau_\eta(h)}\right)^{-\delta}\right]^{(1-h)/\delta}}$$
(A3)

and

$$\mathcal{P}\left(h,\frac{\tau}{T_L}, Re, \mathcal{D}(h), \vartheta\right) = \frac{\left(\frac{\tau}{T_L}\frac{1}{\vartheta}\right)^{1-\mathcal{D}(h)}}{\left[1 + \left(\frac{\tau}{\tau_{\eta}(h)}\right)^{-\delta}\right]^{(\mathcal{D}(h)-1)/\delta}}.$$
 (A4)

In addition, the so-called singularity spectrum \mathcal{D} and the local (fluctuating) dissipation time scale were also chosen following Chevillard *et al.* (2003) as

$$\mathcal{D}(h) = 1 - \frac{(h - c_1)^2}{2\left(c_1 - \frac{1}{2}\right)},\tag{A5}$$

$$\tau_n = T_L R e^{-1/(2h+1)}.$$
 (A6)

The free parameters of the model are thus the Reynolds number, Re, and the integral time scale, T, that were given in § 2 of the paper, and three additional free parameters, δ , c_1 and ϑ that govern the details of the transition from dissipation to the inertial regimes. The three parameters were fitted to the empirical data by minimizing the difference between the flatness coefficient at fixed τ values. The values that were obtained and that were used to plot figure 3 are $\delta = 0.6$, $c_1 = 0.593$ and $\vartheta = 3.5$.

REFERENCES

- ANTONIA, R.A. 1981 Conditional sampling in turbulence measurement. Annu. Rev. Fluid Mech. 13 (1), 131–156.
- ARNÈODO, A., et al. 2008 Universal intermittent properties of particle trajectories in highly turbulent flows. *Phys. Rev. Lett.* **100**, 254504.
- AYLOR, D.E. 1990 The role of intermittent wind in the dispersal of fungal pathogens. *Annu. Rev. Phytopathol.* **28** (1), 73–92.
- BATCHELOR, G.K., TOWNSEND, A.A. & JEFFREYS, H. 1949 The nature of turbulent motion at large wave-numbers. *Proc. R. Soc. Lond.* A **199** (1057), 238–255.

BEC, J., BIFERALE, L., CENCINI, M., LANOTTE, A.S. & TOSCHI, F. 2006 Effects of vortex filaments on the velocity of tracers and heavy particles in turbulence. *Phys. Fluids* **18** (8), 081702.

- BENTKAMP, L., LALESCU, C.C. & WILCZEK, M. 2019 Persistent accelerations disentangle Lagrangian turbulence. *Nat. Commun.* **10**, 3550.
- BENZI, R., BIFERALE, L., FISHER, R., LAMB, D.Q. & TOSCHI, F. 2010 Inertial range Eulerian and Lagrangian statistics from numerical simulations of isotropic turbulence. J. Fluid Mech. 653, 221–244.
- BIFERALE, L., BOFFETTA, G., CELANI, A., LANOTTE, A. & TOSCHI, F. 2005 Particle trapping in three-dimensional fully developed turbulence. *Phys. Fluids* **17** (2), 021701.
- BLUM, D.B., BEWLEY, G.P., BODENSCHATZ, E., GIBERT, M., GYLFASON, Á., MYDLARSKI, L., VOTH, G.A., XU, H. & YEUNG, P.K. 2011 Signatures of non-universal large scales in conditional structure functions from various turbulent flows. *New J. Phys.* 13 (11), 113020.
- BLUM, D.B., KUNWAR, S.B., JOHNSON, J. & VOTH, G.A. 2010 Effects of nonuniversal large scales on conditional structure functions in turbulence. *Phys. Fluids* **22** (1), 015107.

BOS, W.J.T. & ZAMANSKY, R. 2019 Power fluctuations in turbulence. Phys. Rev. Lett. 122, 124504.

- CHEVILLARD, L., ROUX, S.G., LEVÊQUE, E., MORDANT, N., PINTON, J.-F. & ARNEODO, A. 2003 Lagrangian velocity statistics in turbulent flows: effects of dissipation. *Phys. Rev. Lett.* **91**, 214502.
- CORRSIN, S. 1943 Investigation of flow in an axially symmetrical heated jet of air. *Tech. Rep.* NACA-ACR-3L23. NACA Wartime Reports.
- DUMAN, T., KATUL, G.G., SIQUEIRA, M.B. & CASSIANI, M. 2014 A velocity-dissipation lagrangian stochastic model for turbulent dispersion in atmospheric boundary-layer and canopy flows. *Boundary-Layer Meteorol.* 152, 1–18.
- DUMAN, T., TRAKHTENBROT, A., POGGI, D., CASSIANI, M. & KATUL, G.G. 2016 Dissipation intermittency increases long-distance dispersal of heavy particles in the canopy sublayer. *Boundary-Layer Meteorol.* 159 (1), 41–68.
- ELSINGA, G.E., ISHIHARA, T. & HUNT, J.C.R. 2020 Extreme dissipation and intermittency in turbulence at very high Reynolds numbers. *Proc. R. Soc. Lond.* A **476** (2243), 20200591.
- FERACO, F., MARINO, R., PUMIR, A., PRIMAVERA, L., MININNI, P.D., POUQUET, A. & ROSENBERG, D. 2018 Vertical drafts and mixing in stratified turbulence: sharp transition with Froude number. *Europhys. Lett.* **123** (4), 44002.
- FINNIGAN, J. 2000 Turbulence in plant canopies. Annu. Rev. Fluid Mech. 32, 519-571.
- FINNIGAN, J.J. 1979 Turbulence in waving wheat. Boundary-Layer Meteorol. 16 (2), 213–236.
- FRISCH, U. 1995 Turbulence: The Legacy of A. N. Kolmogorov. Cambridge University Press.
- GAO, W., SHAW, R.H. & PAW U, K.T. 1989 Observation of Organized Structure in Turbulent Flow within and above a Forest Canopy, pp. 349–377. Springer.
- HUANG, Y., BIFERALE, L., CALZAVARINI, E., SUN, C. & TOSCHI, F. 2013 Lagrangian single-particle turbulent statistics through the Hilbert–Huang transform. *Phys. Rev. E* 87, 041003.
- KAILASNATH, P., SREENIVASAN, K.R. & STOLOVITZKY, G. 1992 Probability density of velocity increments in turbulent flows. *Phys. Rev. Lett.* 68, 2766–2769.
- KEYLOCK, C.J., GHISALBERTI, M., KATUL, G.G. & NEPF, H.M. 2020 A joint velocity-intermittency analysis reveals similarity in the vertical structure of atmospheric and hydrospheric canopy turbulence. *Environ. Fluid Mech.* **20**, 77–101.
- LIBERZON, A., LÜTHI, B., HOLZNER, M., OTT, S., BERG, J. & MANN, J. 2012 On the structure of acceleration in turbulence. *Physica* D 241 (3), 208–215.
- LOUKA, P., BELCHER, S.E. & HARRISON, R.G. 2000 Coupling between air flow in streets and the well-developed boundary layer aloft. *Atmos. Environ.* **34** (16), 2613–2621.
- MELLER, Y. & LIBERZON, A. 2016 Particle data management software for 3d particle tracking velocimetry and related applications the flowtracks package. J. Open Res. Softw. 4 (1), e23.
- MONIN, A.S. & YAGLOM, A.M. 1972 Statistical Fluid Mechanics. Dover Publications Inc.
- MORDANT, N., DELOUR, J., LÉVEQUE, E., ARNÉODO, A. & PINTON, J.-F. 2002 Long time correlations in Lagrangian dynamics: a key to intermittency in turbulence. *Phys. Rev. Lett.* **89**, 254502.
- MORDANT, N., LÉVÊQUE, E. & PINTON, J.F. 2004 Experimental and numerical study of the Lagrangian dynamics of high Reynolds turbulence. New J. Phys. 6, 116–116.
- MUCHINSKI, A., FREHLICH, R. & BALSLEY, B. 2004 Small-scale and large-scale intermittency in the nocturnal boundary layer and the residual layer. J. Fluid Mech. 515, 319–351.
- OPENPTV CONSORTIUM 2014 Open source particle tracking velocimetry. http://www.openptv.net/
- POPE, S.B. 1991 Application of the velocity-dissipation probability density function model to inhomogeneous turbulent flows. *Phys. Fluids* A **3** (8), 1947–1957.
- POPE, S.B. & CHEN, Y.L. 1990 The velocity-dissipation probability density function model for turbulent flows. *Phys. Fluids* A **2** (8), 1437–1449.
- RAUPACH, M.R., COPPIN, P.A. & LEGG, B.J. 1986 Experiments on scalar dispersion within a model plant canopy. Part 1: the turbulence structure. *Boundary-Layer Meteorol.* **35**, 21–52.
- REYNOLDS, A.M. 1998 On the formulation of Lagrangian stochastic models of scalar dispersion within plant canopies. *Boundary-Layer Meteorol.* **86** (2), 333–344.
- SHAW, R.H., TAVANGAR, J. & WARD, D.P. 1983 Structure of the Reynolds stress in a canopy layer. J. Clim. Appl. Meteorol. 22 (11), 1922–1931.
- SHE, Z. & LEVEQUE, E. 1994 Universal scaling laws in fully developed turbulence. *Phys. Rev. Lett.* **72**, 336–339.
- SHNAPP, R., BOHBOT-RAVIV, Y., LIBERZON, A. & FATTAL, E. 2020 Turbulence-obstacle interactions in the Lagrangian framework: applications for stochastic modeling in canopy flows. *Phys. Rev. Fluids* 5, 094601.
- SHNAPP, R., SHAPIRA, E., PERI, D., BOHBOT-RAVIV, Y., FATTAL, E. & LIBERZON, A. 2019 Extended 3D-PTV for direct measurements of Lagrangian statistics of canopy turbulence in a wind tunnel. *Sci. Rep.* 9, 7405.

- SREENIVASAN, K.R. & DHRUVA, B. 1998 Is there scaling in high-Reynolds-number turbulence?. Prog. Theor. Phys. Supp. 130, 103–120.
- TOSCHI, F. & BODENSCHATZ, E. 2009 Lagrangian properties of particles in turbulence. Annu. Rev. Fluid Mech. 41 (1), 375–404.
- TSINOBER, A. 2009 An Informal Conceptual Introduction to Turbulence. Springer.
- VIGGIANO, B., FRIEDRICH, J., VOLK, R., BOURGOIN, M., CAL, R.B. & CHEVILLARD, L. 2020 Modelling Lagrangian velocity and acceleration in turbulent flows as infinitely differentiable stochastic processes. *J. Fluid Mech.* 900, A27.
- VOTH, G.A., SATYANARAYAN, K. & BODENSCHATZ, E. 1998 Lagrangian acceleration measurements at large Reynolds numbers. *Phys. Fluids* 10 (9), 2268–2280.
- WATTEAUX, R., SARDINA, G., BRANDT, L. & IUDICONE, D. 2019 On the time scales and structure of Lagrangian intermittency in homogeneous isotropic turbulence. J. Fluid Mech. 867, 438–481.
- WILCZEK, M., JENKO, F. & FRIEDRICH, R. 2008 Lagrangian particle statistics in turbulent flows from a simple vortex model. *Phys. Rev.* E 77, 056301.
- WILCZEK, M., XU, H., OUELLETTE, N.T., FRIEDRICH, R. & BODENSCHATZ, E. 2013 Generation of Lagrangian intermittency in turbulence by a self-similar mechanism. *New J. Phys.* **15** (5), 055015.
- WILSON, J.D. & SAWFORD, B.L. 1996 Review of Lagrangian stochastic models for trajectories in the turbulent atmosphere. *Boundary-Layer Meteorol.* 78, 191–210.
- WYGNANSKI, I.J. & CHAMPAGNE, F.H. 1973 On transition in a pipe. Part 1. The origin of puffs and slugs and the flow in a turbulent slug. J. Fluid Mech. 59 (2), 281–335.
- XU, H., PUMIR, A., FALKOVICH, G., BODENSCHATZ, E., SHATS, M., XIA, H., FRANCOIS, N. & BOFFETTA, G. 2014 Flight-crash events in turbulence. *Proc. Natl Acad. Sci. USA* **111** (21), 7558–7563.
- YUE, W., MENEVEAU, C., PARLANGE, M.B., ZHU, W., VAN HOUT, R. & KATZ, J. 2007 A comparative quadrant analysis of turbulence in a plant canopy. *Water Resour. Res.* 43 (5), W05422.
- ZHU, W., VAN HOUT, R. & KATZ, J. 2007 On the flow structure and turbulence during sweep and ejection events in a wind-tunnel model canopy. *Boundary-Layer Meteorol.* 124 (2), 205–233.